

火星衛星フォボスとデイモスの 起源と進化の現状理解

Phobos

Ryuki Hyodo (ELSI, Tokyo Tech),

Thanks to S. Charnoz (IPGP), H. Genda (ELSI, Tokyo-Tech),
K. Kurosawa (Chiba Institut.), T. Nakamura (Tohoku Uni.)

and
MMX science team

Related papers:

Rosenblatt, Charnoz, ... [Hyodo](#), ... (2016), *Nature Geo.*

[Hyodo](#) et al. (2017a), *ApJ*

[Hyodo](#) et al. (2017b), *ApJ*

Pignatale, Charnoz, ... [Hyodo](#), ... (2018), *ApJ*

[Hyodo](#) & Genda (2018), *ApJL*

[Hyodo](#) et al. (2018), *ApJ*

[Hyodo](#) et al. (2018) in *planetary and space science*

Deimos



本日の内容

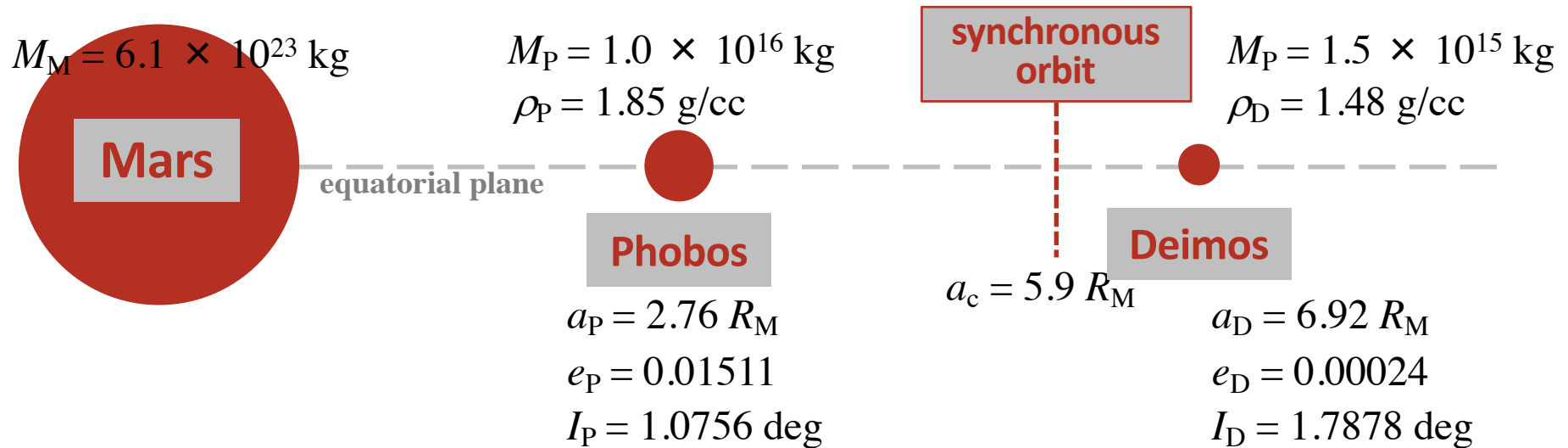
1. Dynamical origin and evolution of Martian moons
2. Physical & chemical properties of Martian moons
3. Recent delivery of Martian material to Martian moons



探査へ

Dynamical origin and evolution of Martian moons

Phobos and Deimos



- ▶ **Orbits:** very **circular** ($e \sim 0$) and parallel to the Mars' **equatorial plane** ($I \sim 0$).
e.g., Rosenblatt (2011) A&A Rev.
- ▶ **Tide:** Deimos is migrating outward because of tidal interaction with Mars
Phobos is migrating inward, and will break apart or hit on Mars in 30 Myrs.
Black & Mittal (2015) Nature Geo.
- ▶ **Spectra:** **featureless** & **very dark** (similar to D-type asteroids)
e.g., Murchie (1999) JGR

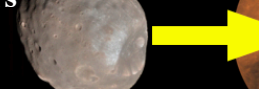
現在の軌道

Deimos: $7R_{\text{Mars}}$



潮汐進化

Phobos: $2.76R_{\text{Mars}}$



潮汐進化
(2 cm/年)

共回転半径: $6R_{\text{Mars}}$

ロッシュ半径: $3R_{\text{Mars}}$

約40億年前の軌道

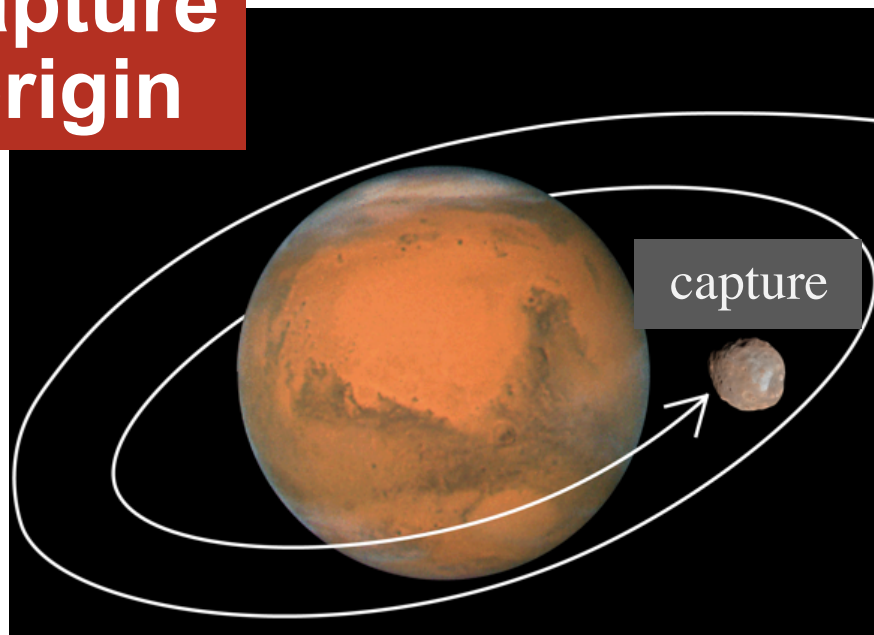
Deimos: $7R_{\text{Mars}}$

Phobos: $5.7R_{\text{Mars}}$

Two Leading Hypothesis

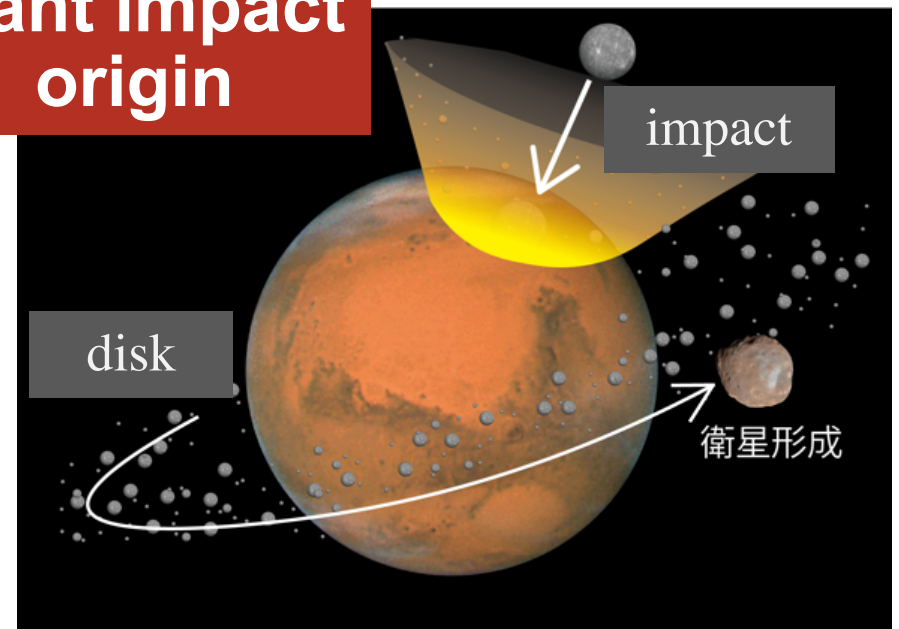
MMX is a JAXA's sample return mission to Martian moons
Primary goal of *MMX* mission is to solve moons' origin

capture
origin



supported by spectral features

giant impact
origin



supported by orbital elements

Dark & Featureless: D-type?

Circular & Equatorial

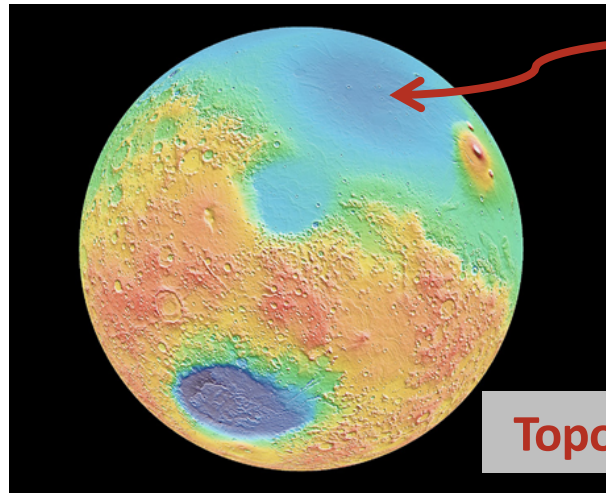
→ This talk

Impact Origin?

Giant impact hypothesis for Moon



- ▶ If this happens on Mars, satellites with $e \sim 0$ and $I \sim 0$ may be naturally formed.
- ▶ Phobos and Deimos are small.
→ small impactor is sufficient

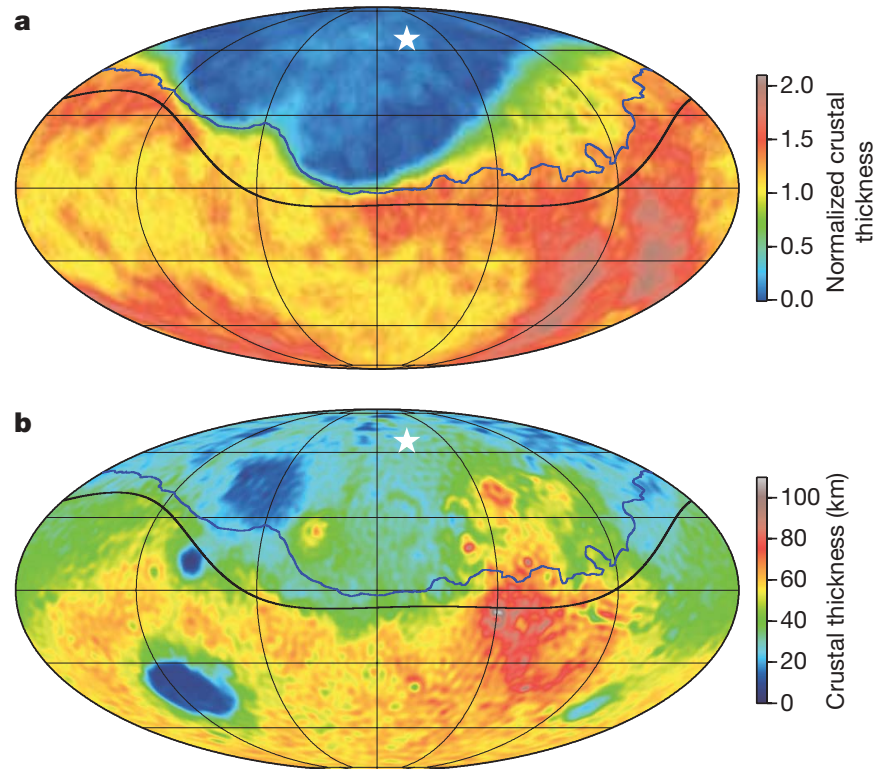


Topography of Mars

Craddock (2011) proposed the impact which made northern lowlands on Mars (Borealis basin) may create Martian moons.

Marinova et al. (2008) showed the impact condition to form the Borealis basin
-> we use the same impact conditions to form Martian moons

Borealis Basin-forming Impact



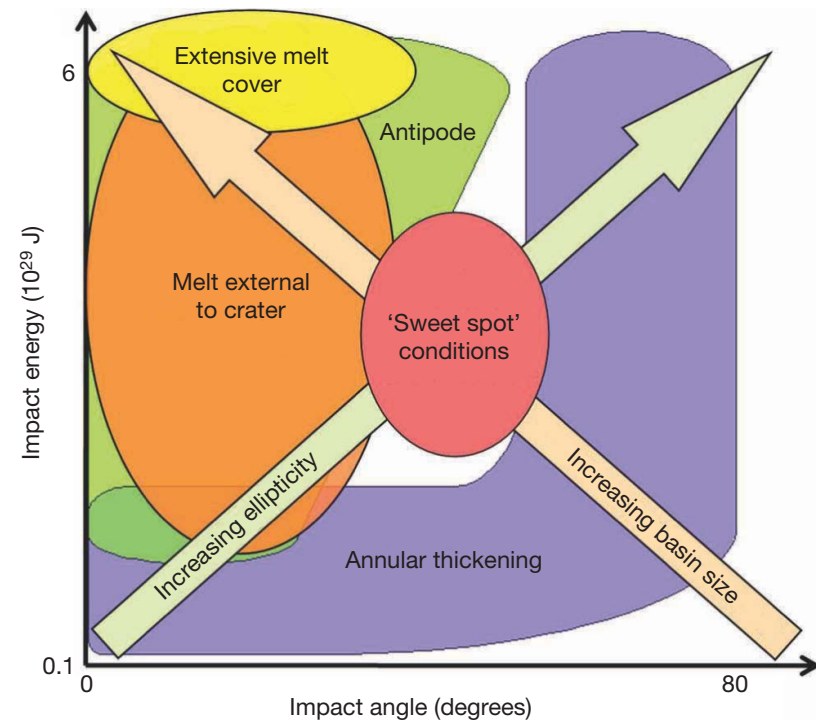
Marinova et al. (2008) Nature

Sweet Spot

Impactor mass : 3% of M_{Mars}

Impact velocity : ~ 6 km/s

Impact angle: ~ 45 deg



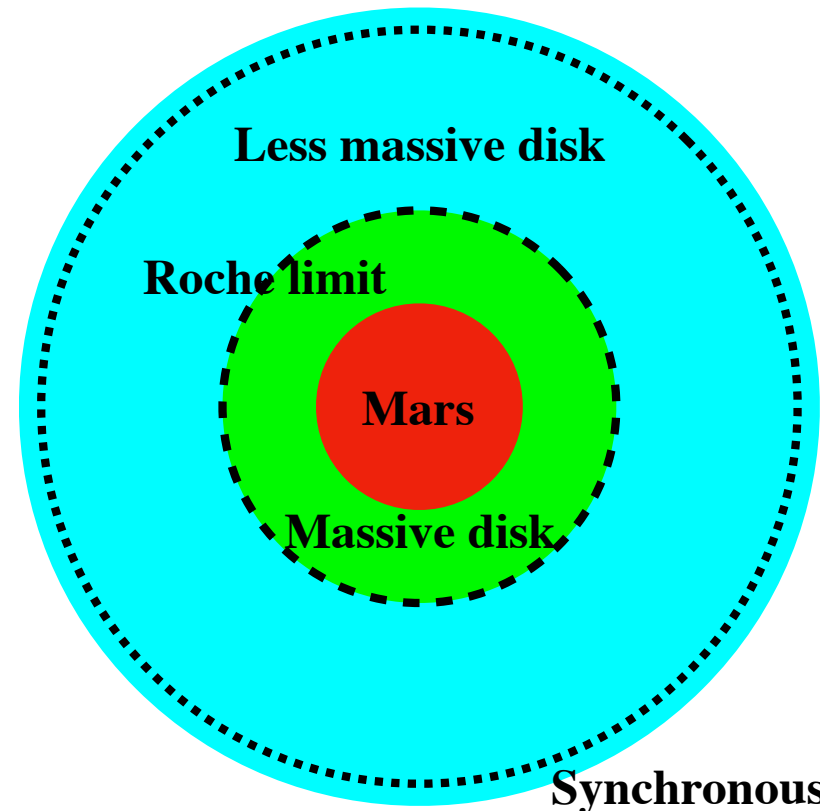
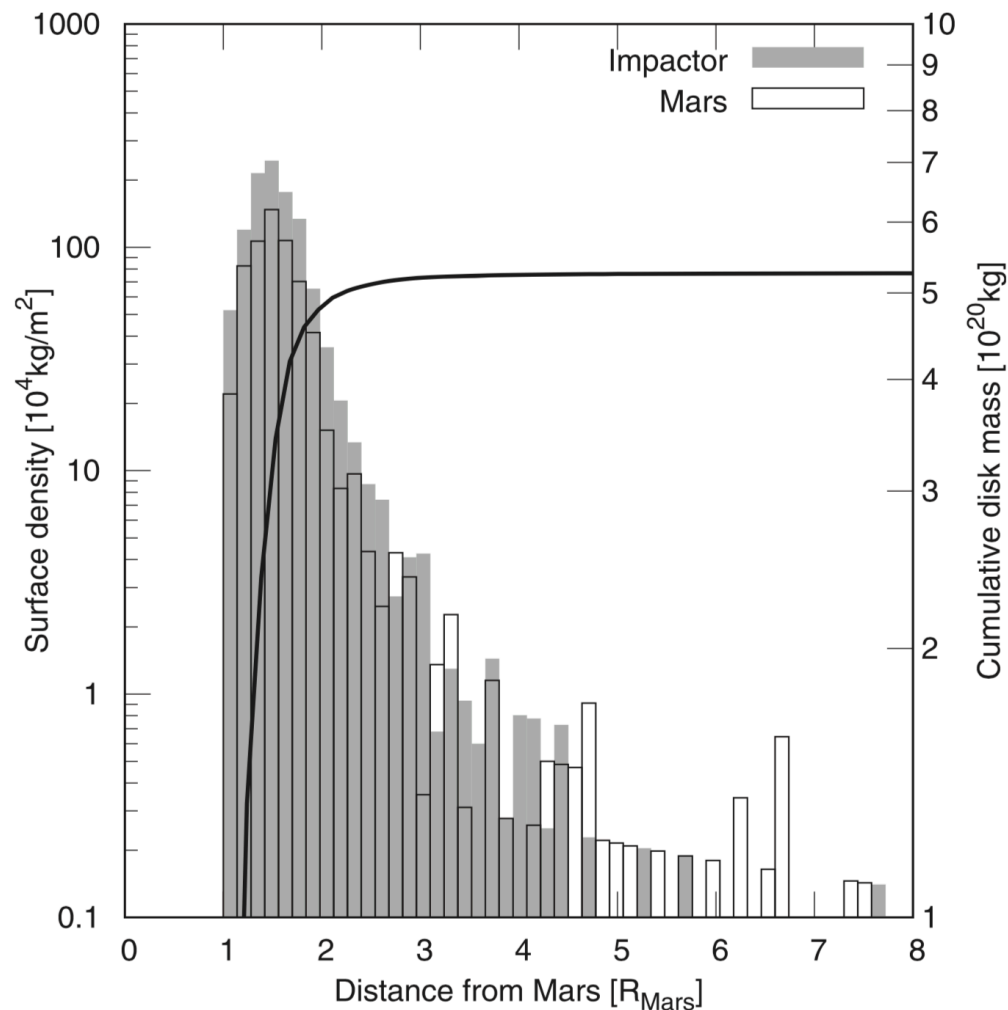
Inner Massive and Outer Light Disk

Impact simulation (SPH method)

Smoothed Particle Hydrodynamics

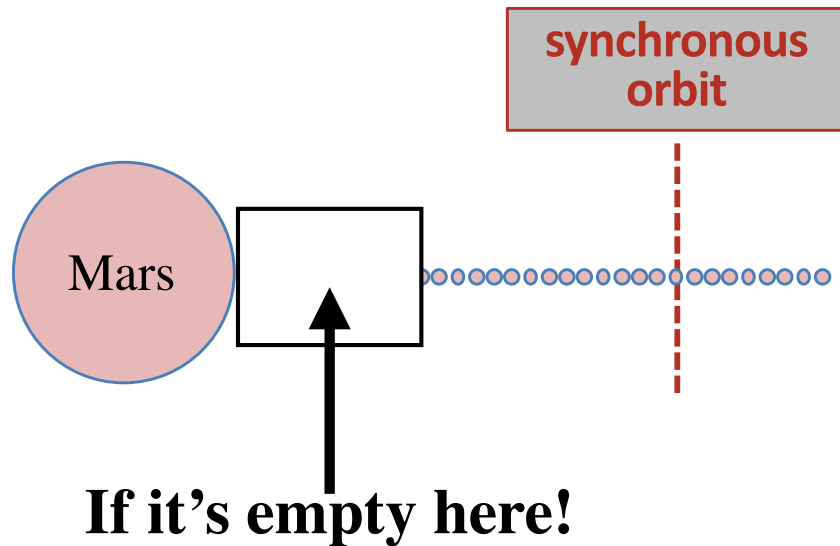
Formed disk is relatively massive.

$\sim 10^{20}$ kg : 10,000 \times Phobos

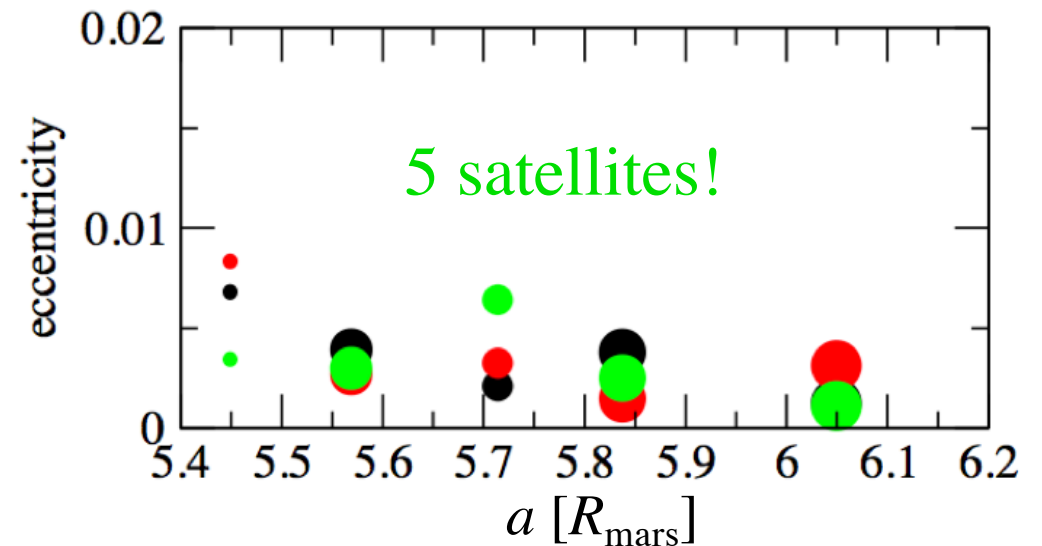


Hyodo, et al. (2017) ApJ

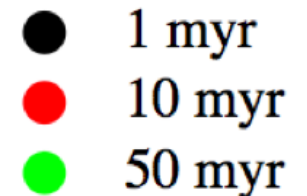
Accretion in Outer Light Disk



An example result of N-body simulation



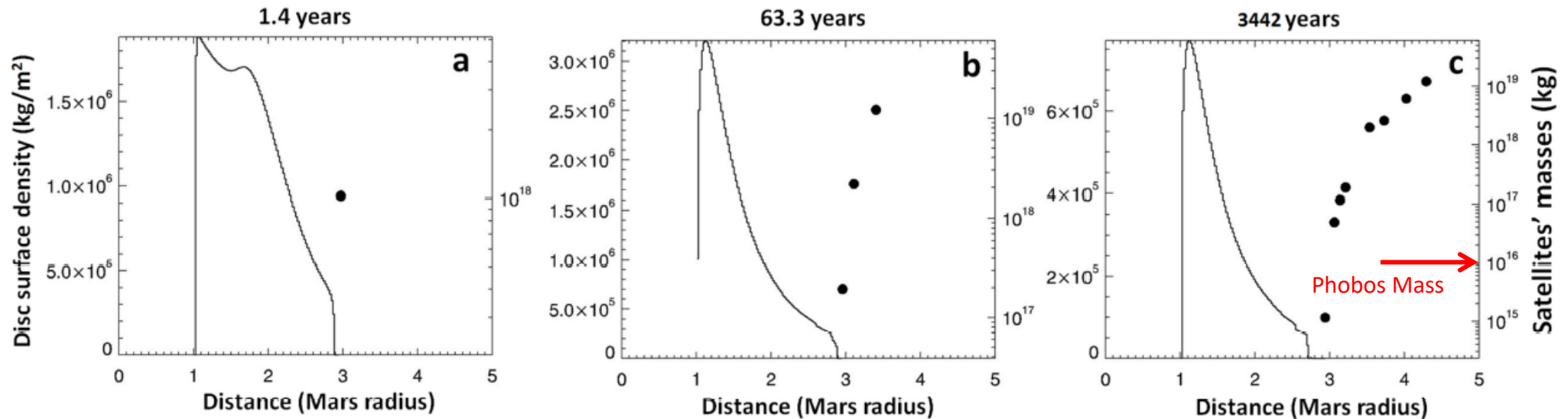
- ✓ More than 2 satellites (typically ~ 5) are formed, because the system is NOT enough dynamically excited.



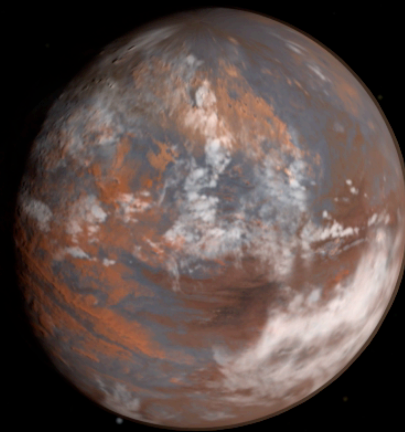
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We need a trick which can enhance or force the accretion of satellites in the outer disk.

Formation of Big Moons from Inner Massive Disk

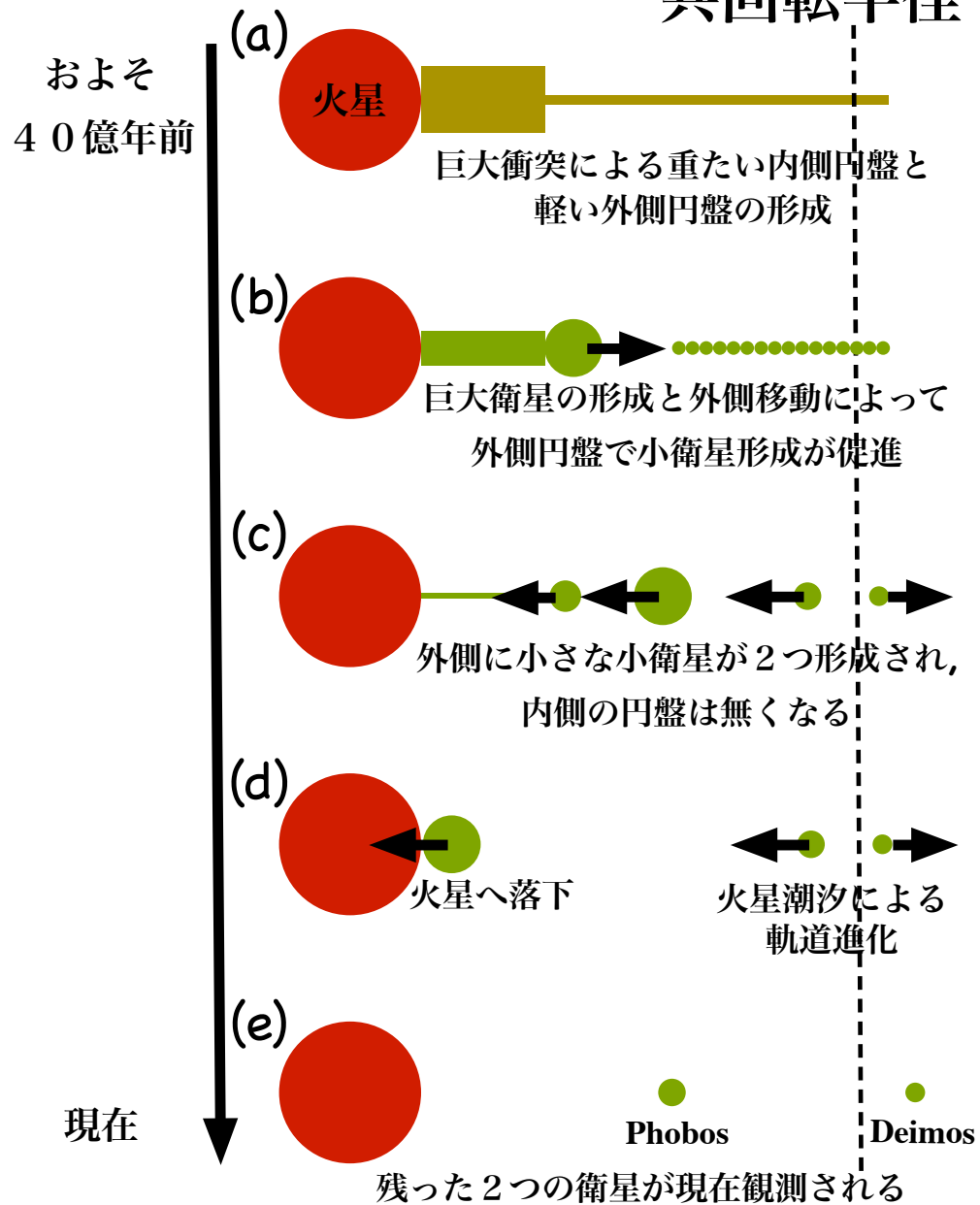


- ✓ From massive inner disk ($10,000 \times$ Phobos), a huge satellite ($\sim 1000 \times$ Phobos) can be formed.
- ✓ This huge satellite moves outward up to $4.4R_{\text{mars}}$ due to the gravitational interaction with the disk.

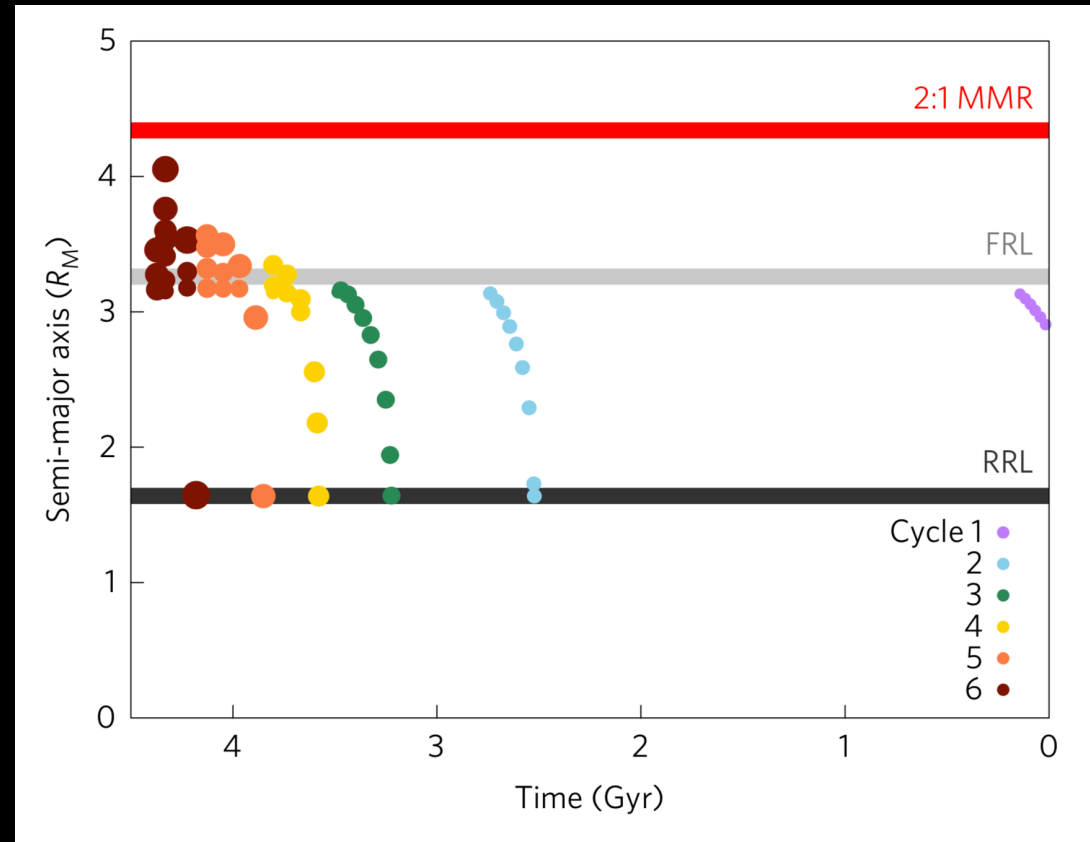
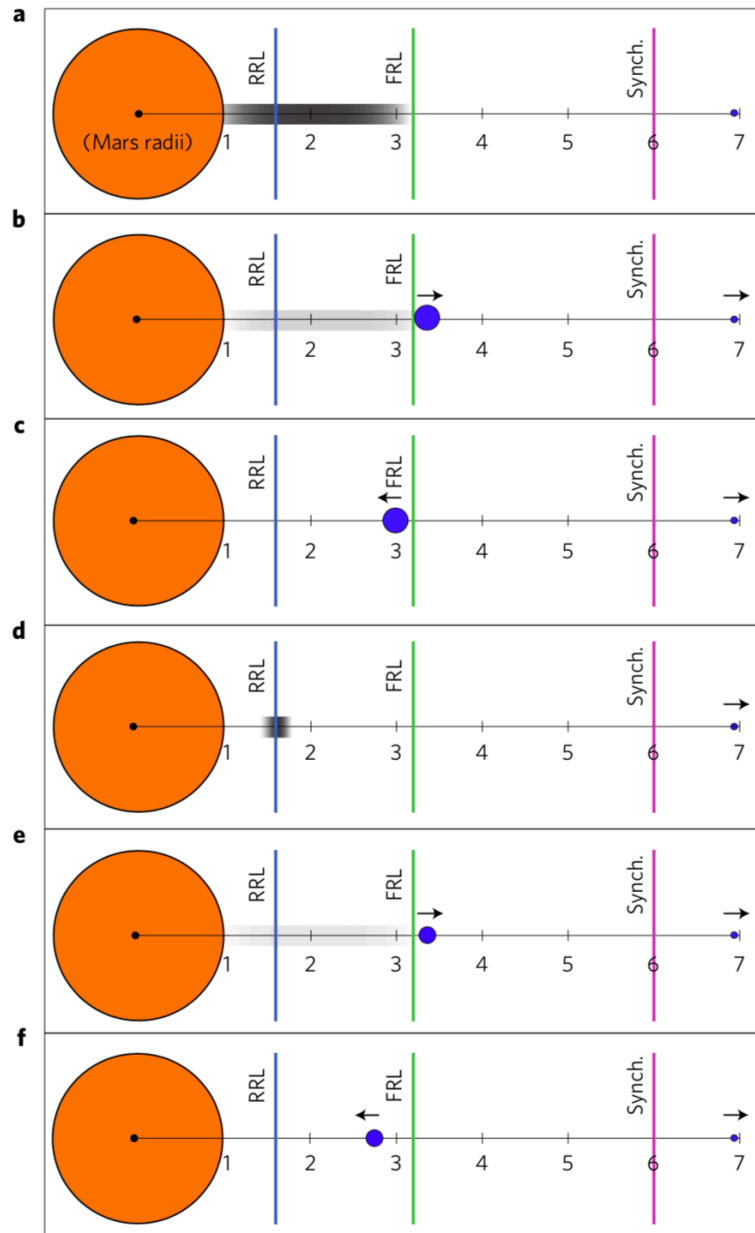


4 billion years ago...

共回転半径

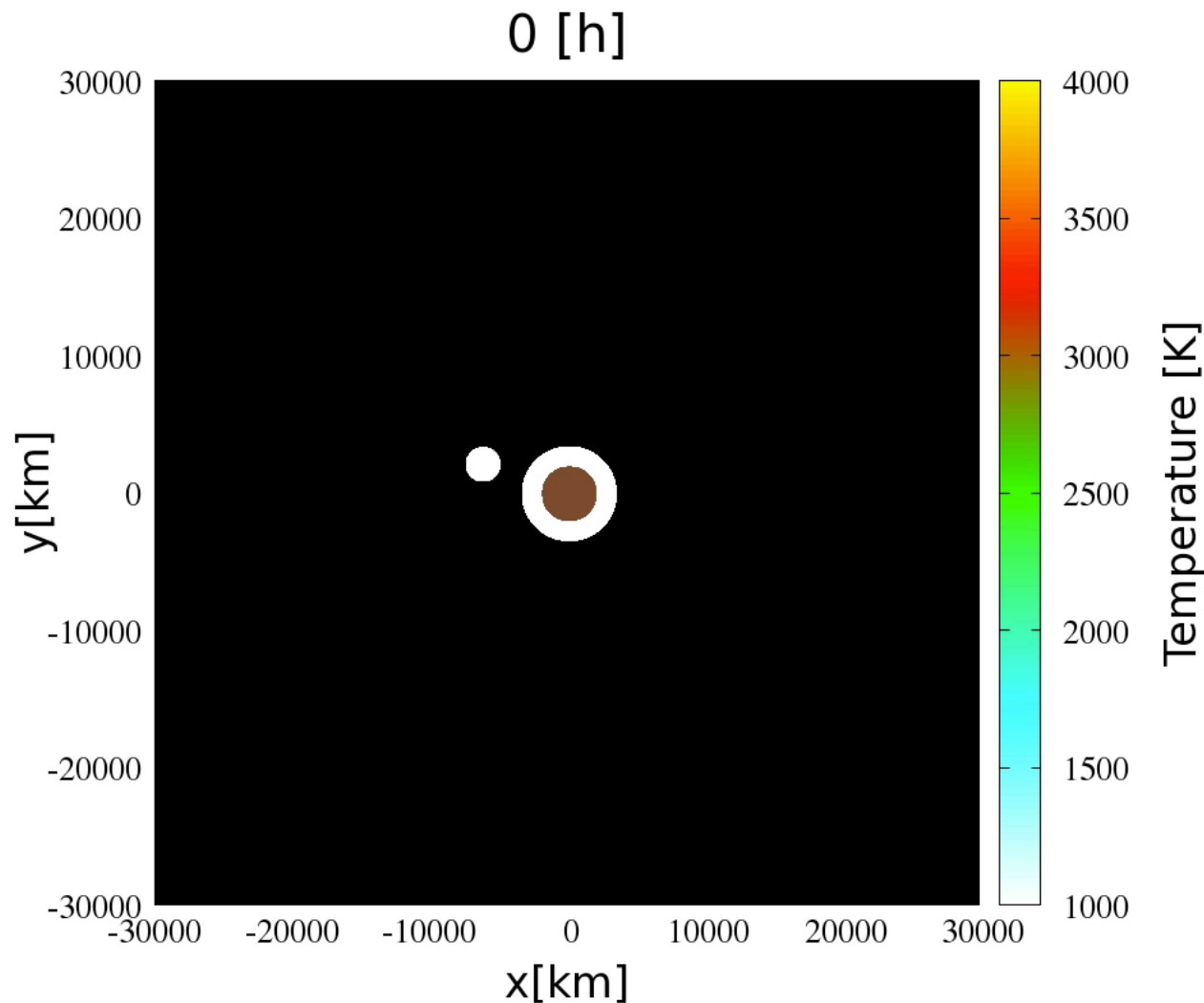


Creations and Destructions of Phobos?



Physical & chemical properties of Martian moons

High-Resolution SPH Simulations



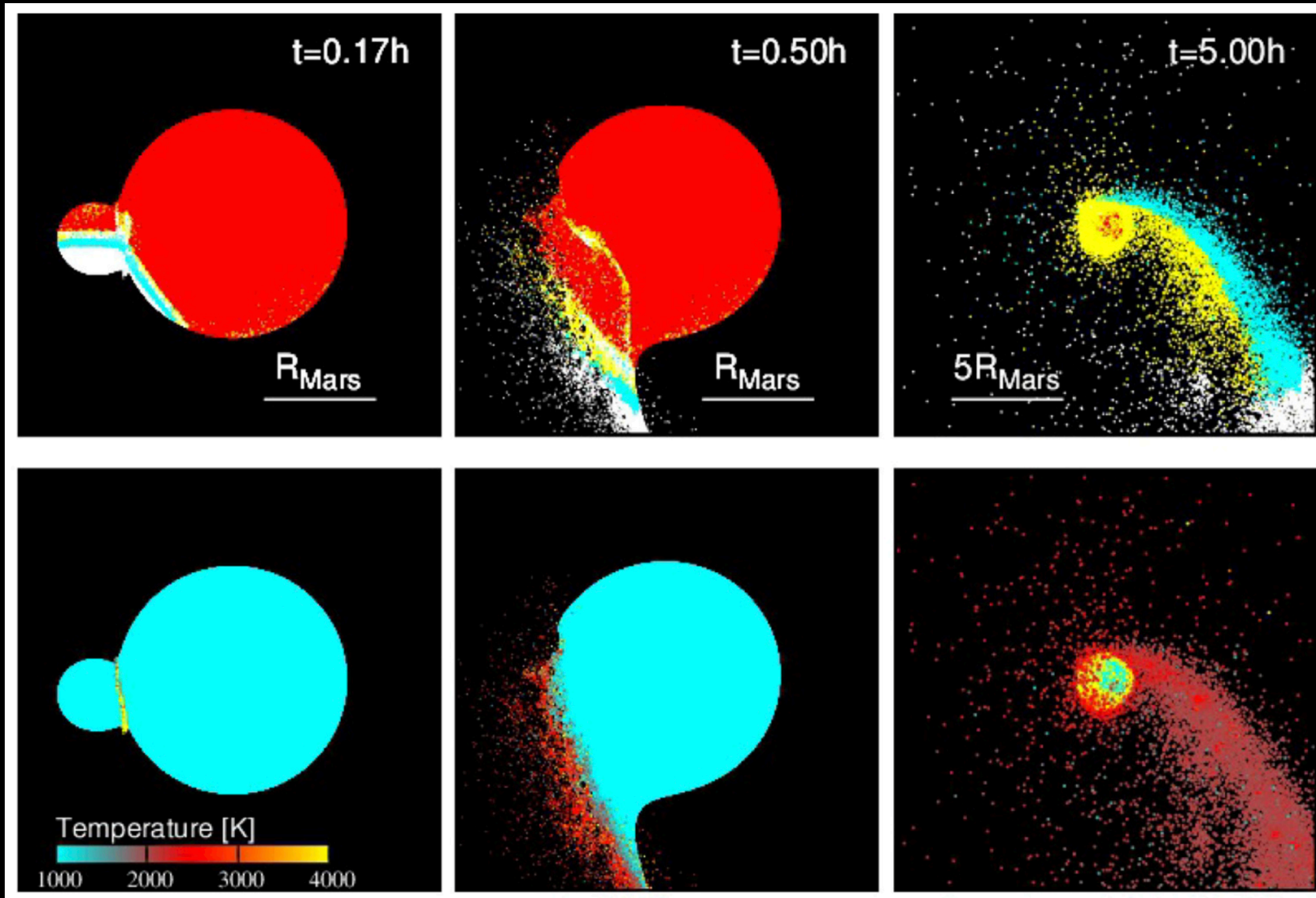
Conditions:

- $N_{\text{SPH}}=3 \times 10^6$
- M-ANEOS
- $\theta_{\text{imp}}=45\text{deg}$
- $V_{\text{imp}}=6\text{km/s}$
- $m_{\text{imp}}=0.03M_{\text{Mars}}$

Outcome:

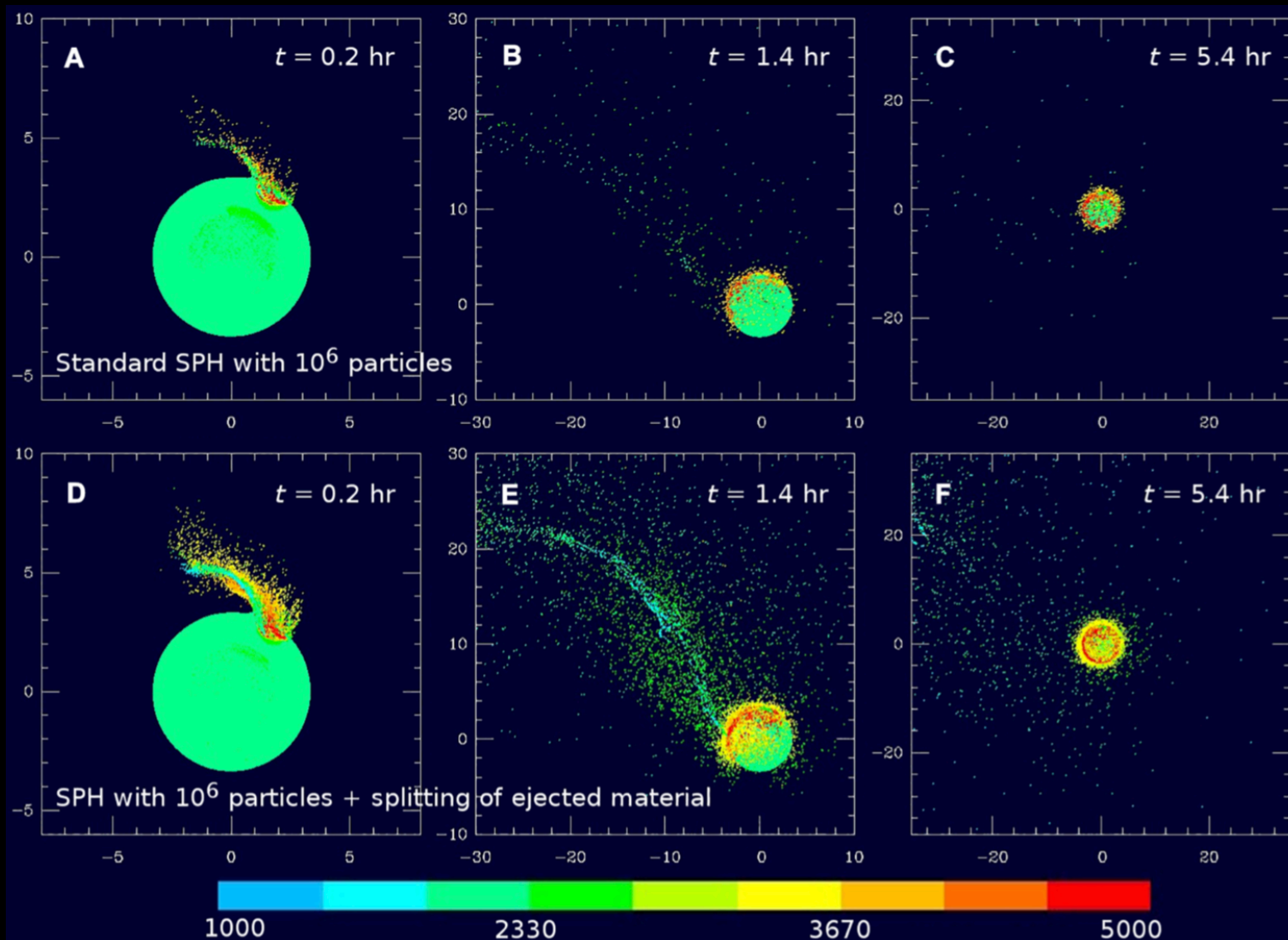
- $M_{\text{disk}}=5 \times 10^{20}\text{kg}$

Debris just after the Ginat Impact



Giant impact by Vesta-to-Ceres impactor?

Physical, thermodynamical & compositional properties are similar to Hyodo et al. 2017

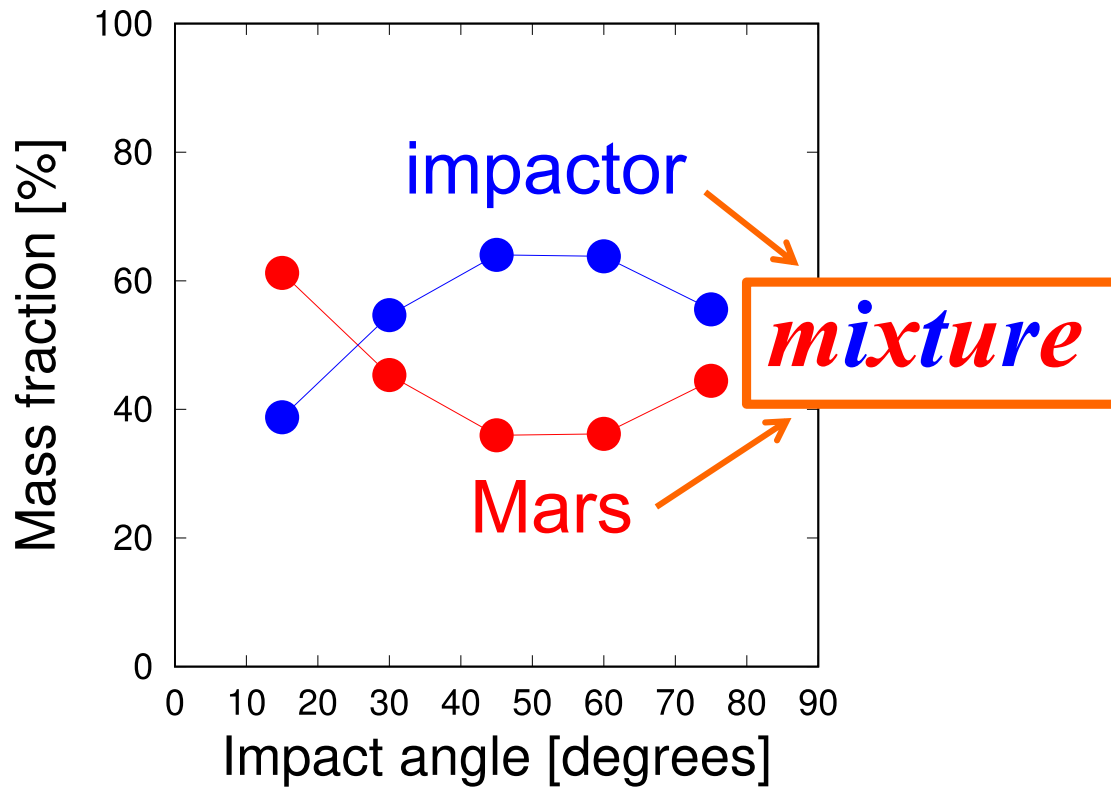


- $\theta_{\text{imp}} = 45 \text{ degs}$
- $V_{\text{imp}} = 7 \text{ km/s}$
- $m_{\text{imp}} = 5 \times 10^{-4} M_{\text{Mars}}$
- $M_{\text{disk}} = 5 \times 10^{18} \text{ kg}$

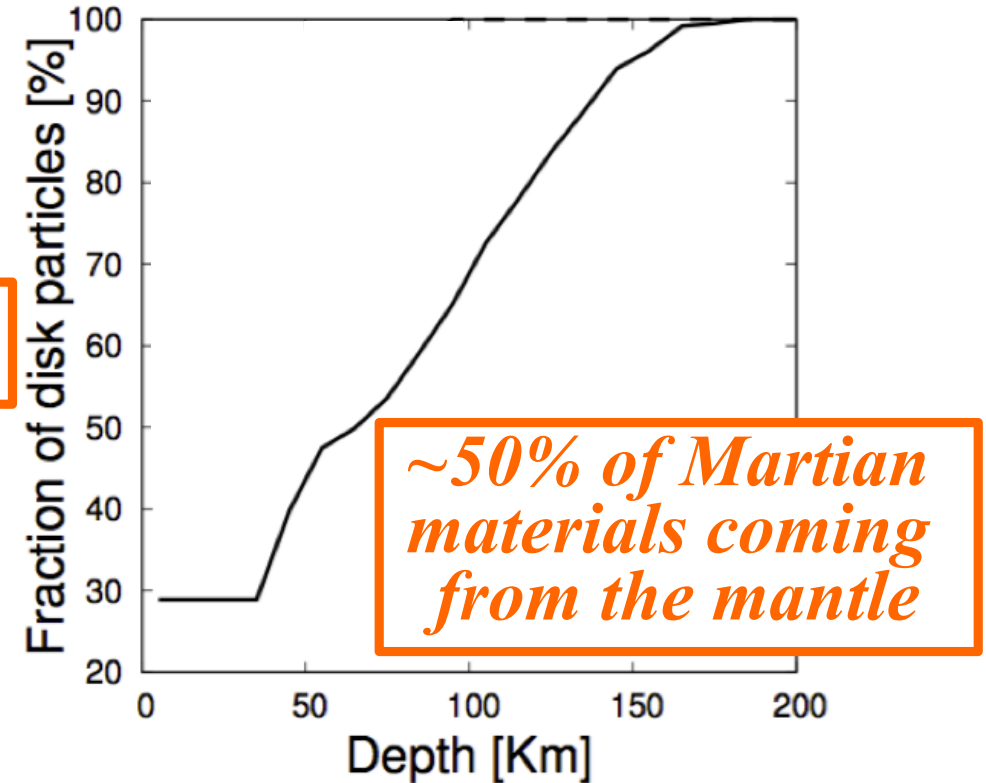
Building Blocks

Phobos and Deimos contain Martian mantle materials

disk composition

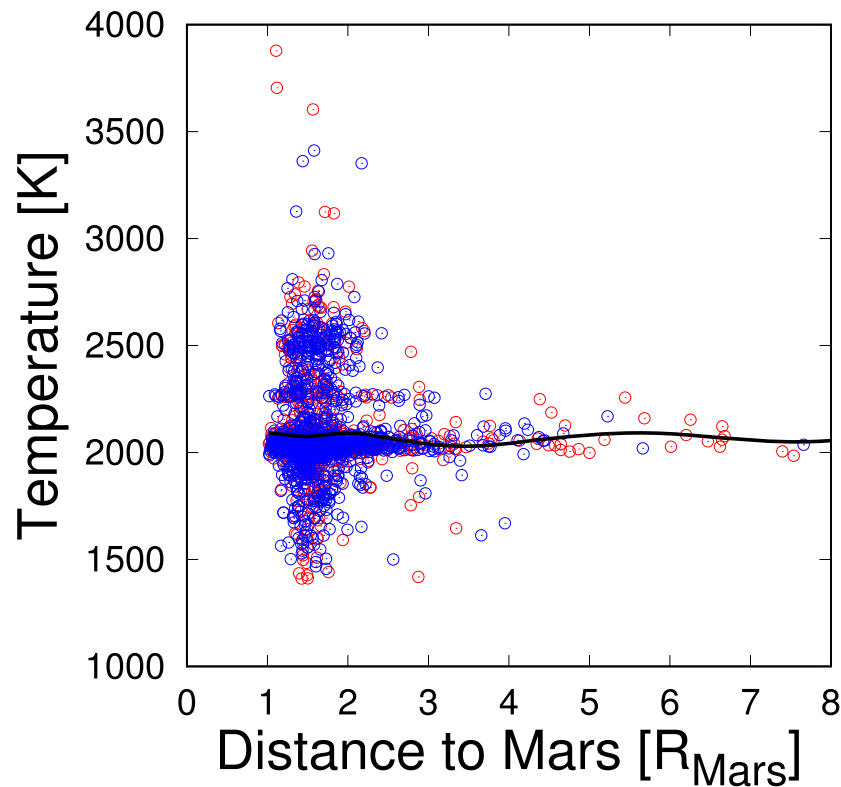


ejected depth



Thermal and Physical Aspects

disk's temperature



disk materials:

almost fully molten

low vapor fraction ($< 5\%$)

particle sizes: *Melosh & Vickery (1991) Nature*

melt fragmentation: ~ 1 m

vapor condensation: $\sim 0.1 \mu\text{m}$

very dark (FeS, C)

featureless

Pignatale, ..., Hyodo (2018) ApJ

Ronnet et al. (2016) ApJ

space-weathered Anorthosite

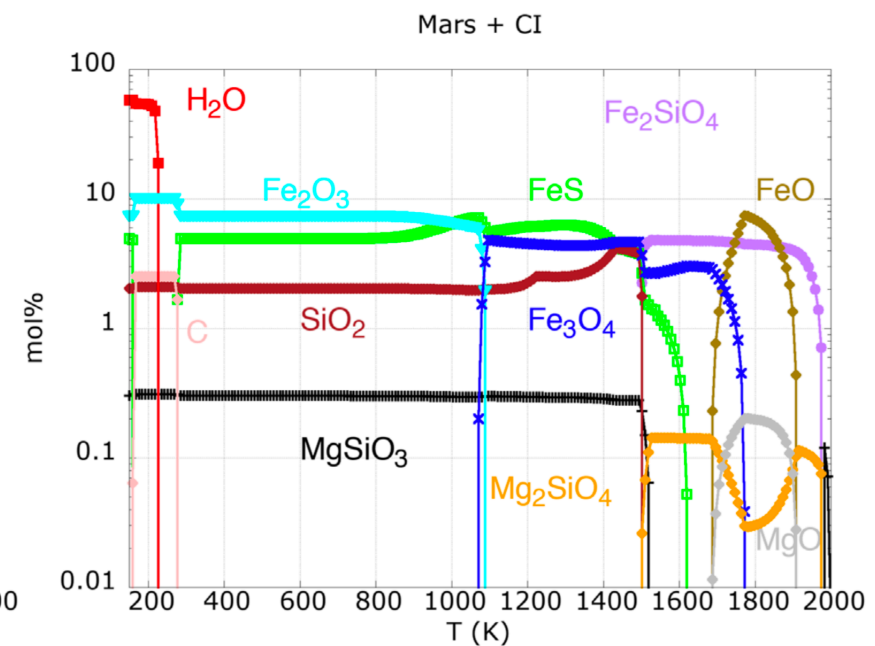
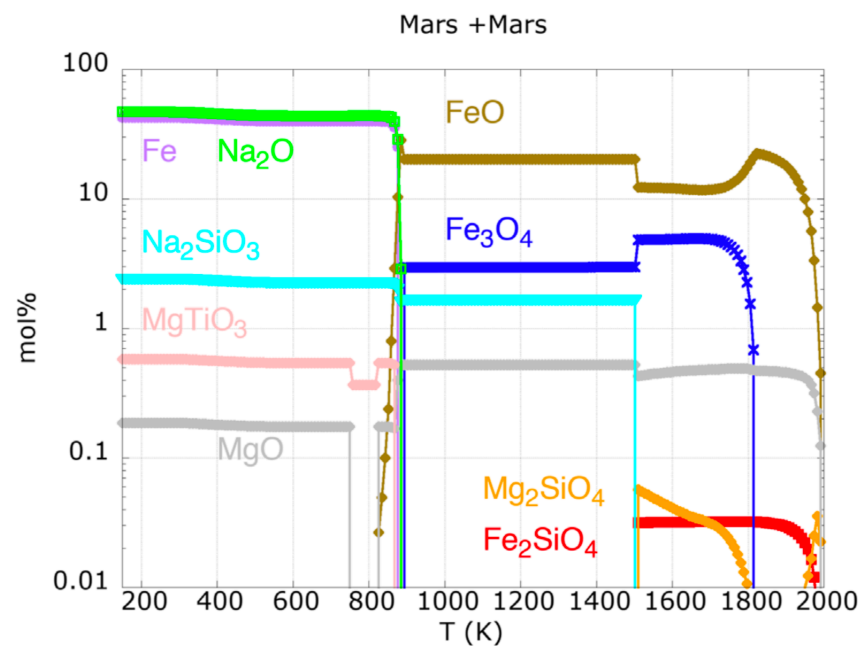
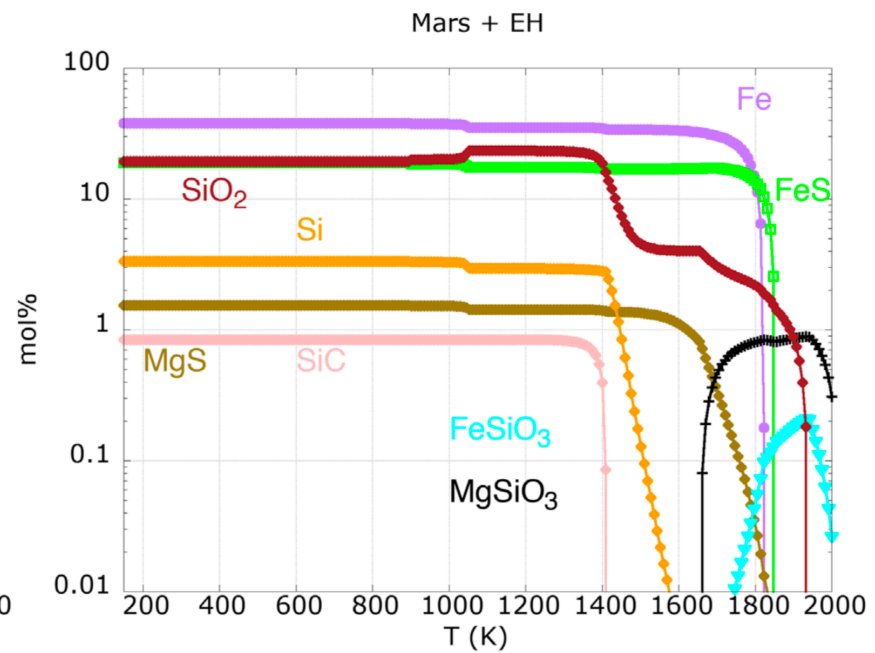
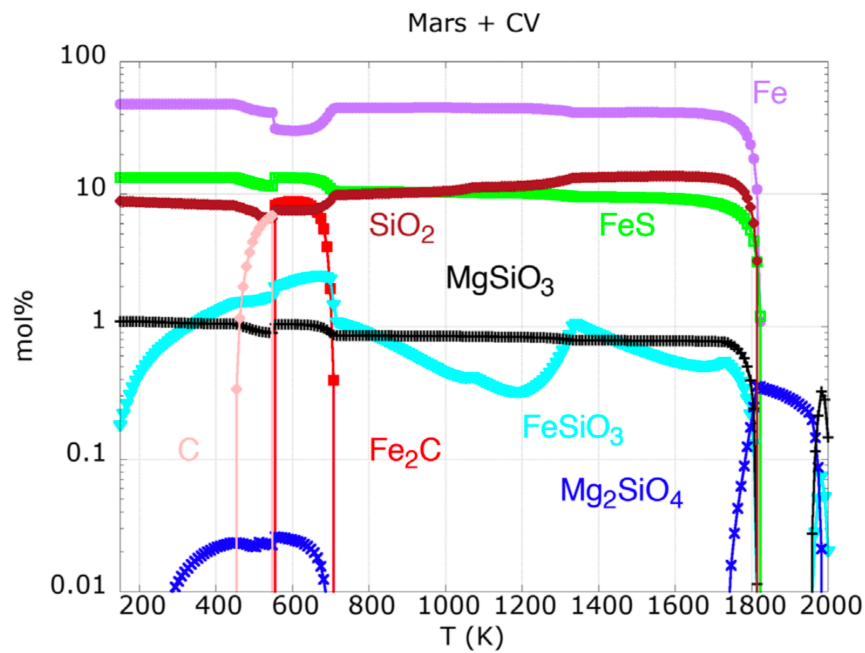
Yamamoto et al. (2018), GRL

disk evolution:

further fragmentation

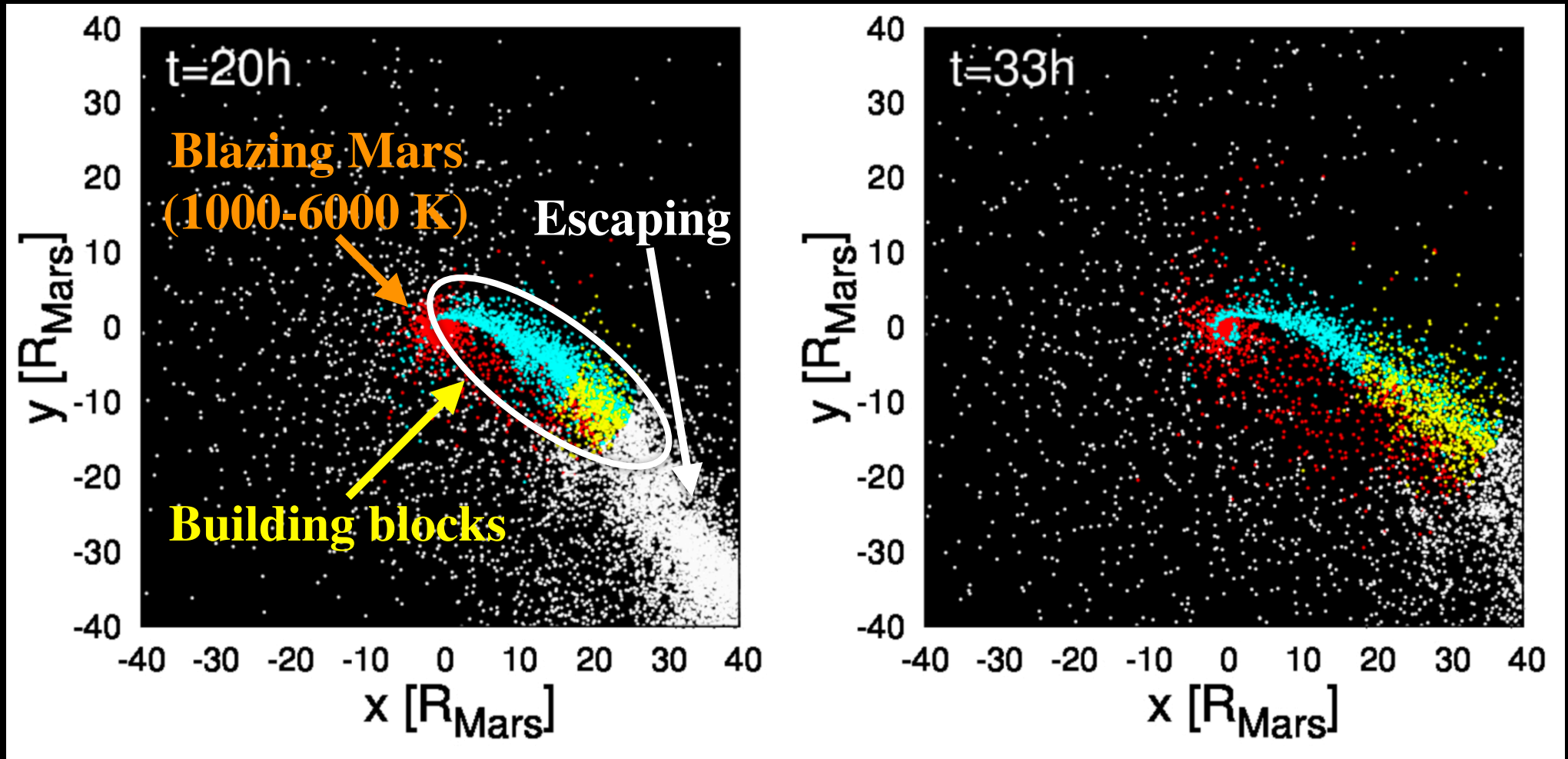
from ~ 1 m to $100 \mu\text{m}$

Dust Chemistry depends on impactor



Volatile Loss from Building Blocks

Hydrodynamic Escape & Radiation Pressure may remove volatiles?

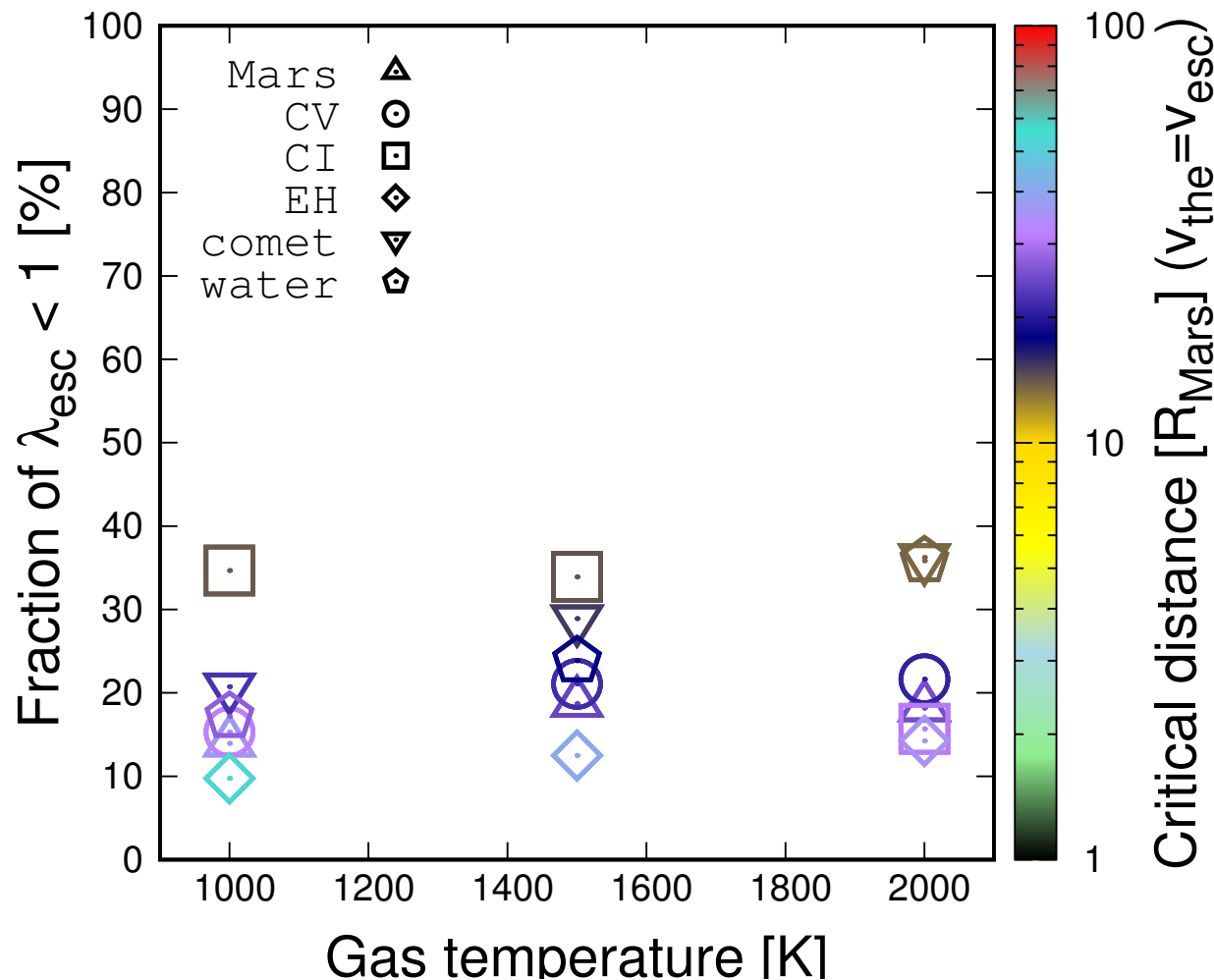


Outer parts of the building blocks: Opacity $\tau < 1$ (yellow points)

Volatile Loss from Disk

$$\lambda_{esc} = \frac{GM_{Mars}m_{vap}}{kT_{vap}r}$$

M_{Mars} : Mars mass, m_{vap} : mean molecular weight of vapor,
 k : the Boltzmann constant, T_{vap} : vapor temperature, r : distance to Mars



Hydrodynamic Escape

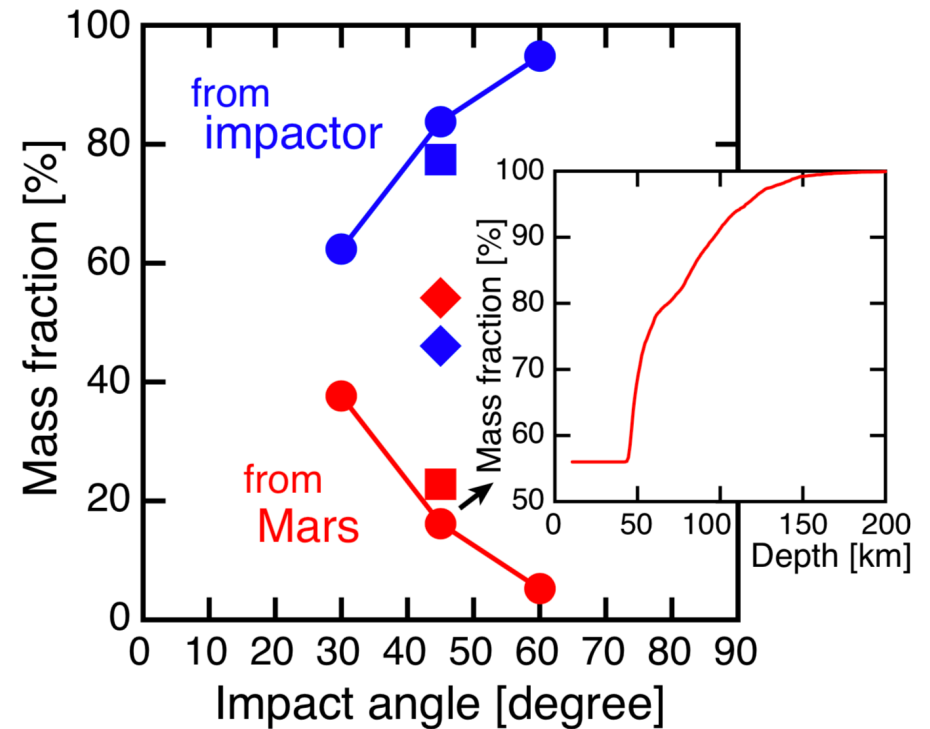
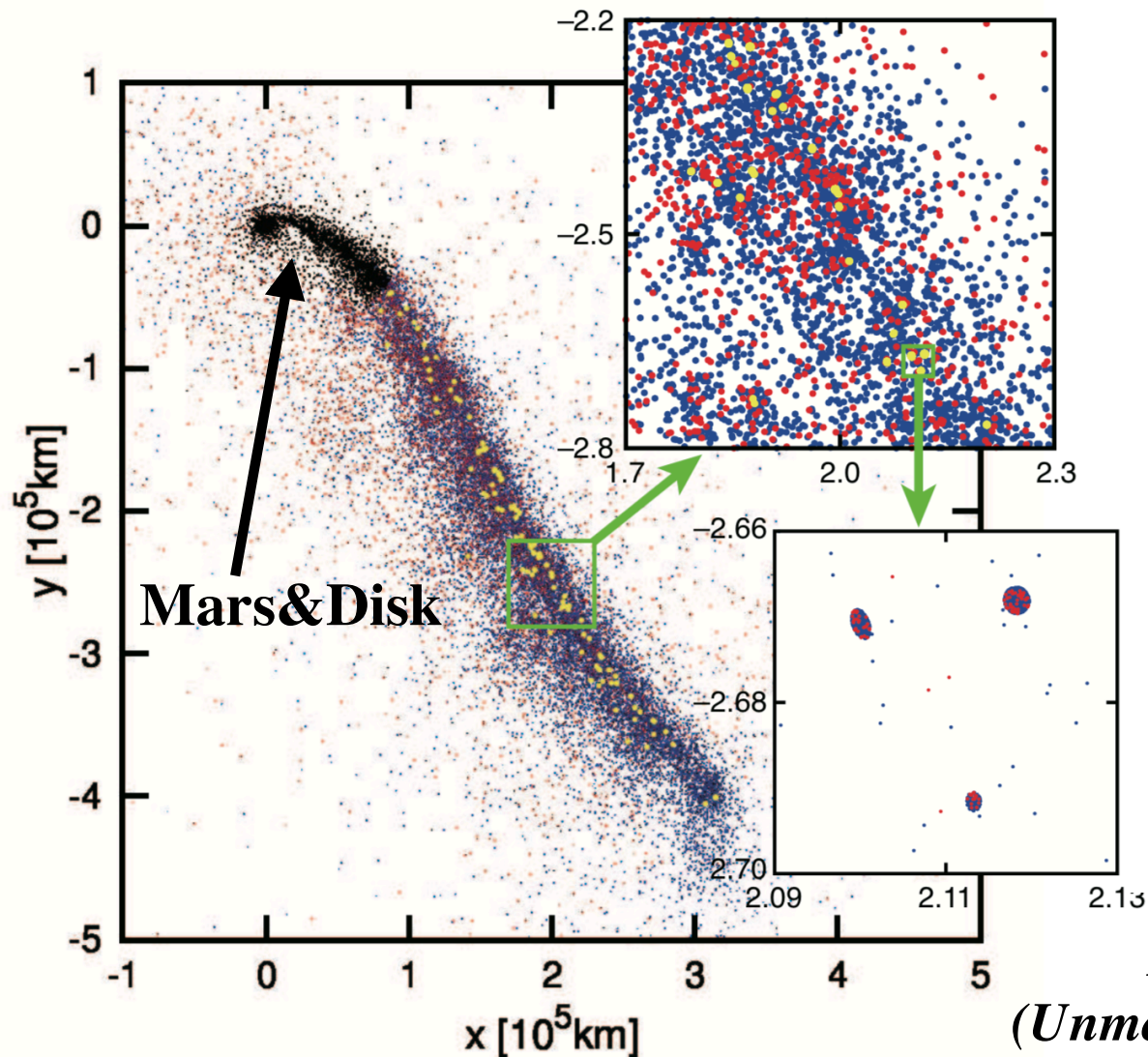
~ 30% of vapor can escape
 → ~ 70% of volatile elements can survive

Radiation Pressure

Moderately volatile elements (condensation $T > 1000$ K) are selectively removed.

[Hyodo, et al. ApJ submitted](#)

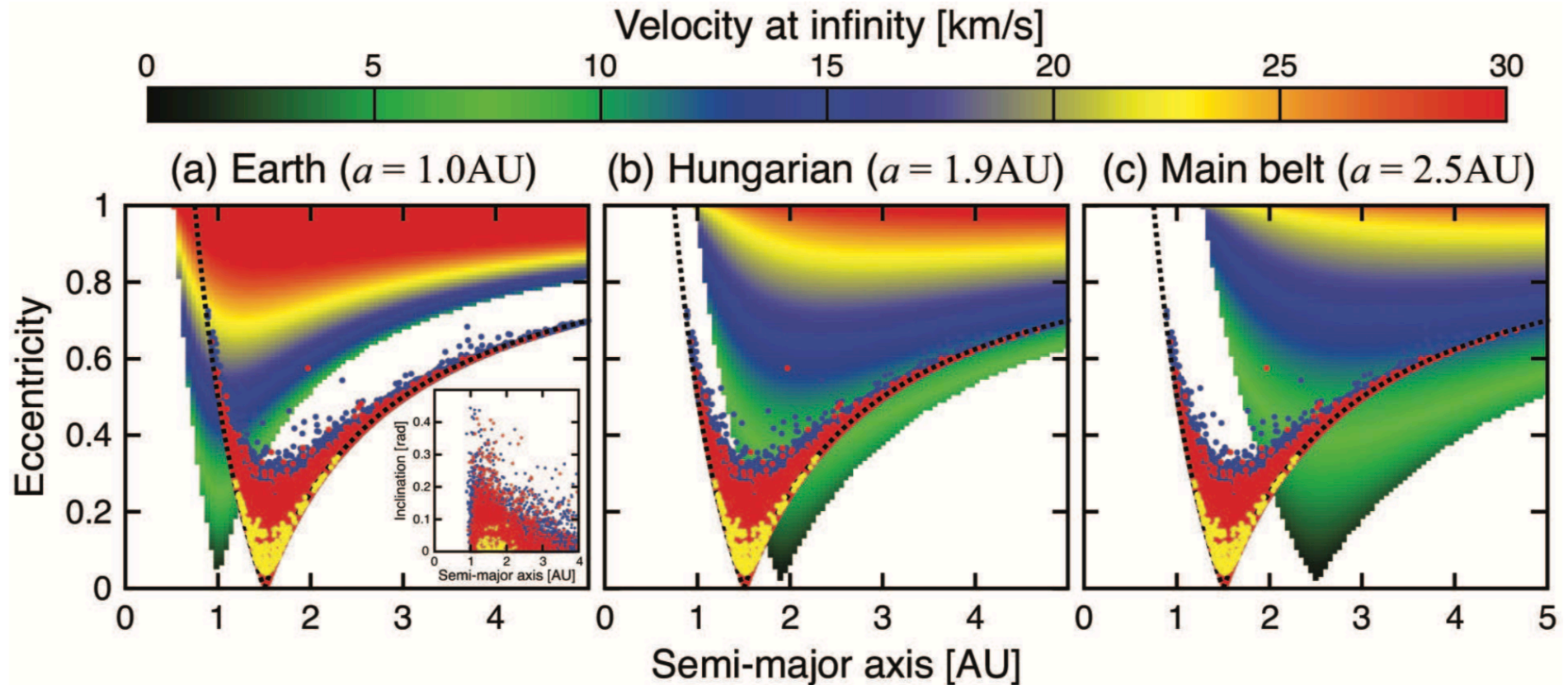
Distribution of the Debris within the Inner Solar System



*Debris mass $M_{debris} \sim 2\%$ of M_{Mars}
(Unmelted Martian Mantle $\sim 0.02\%$ of M_{Mars})*

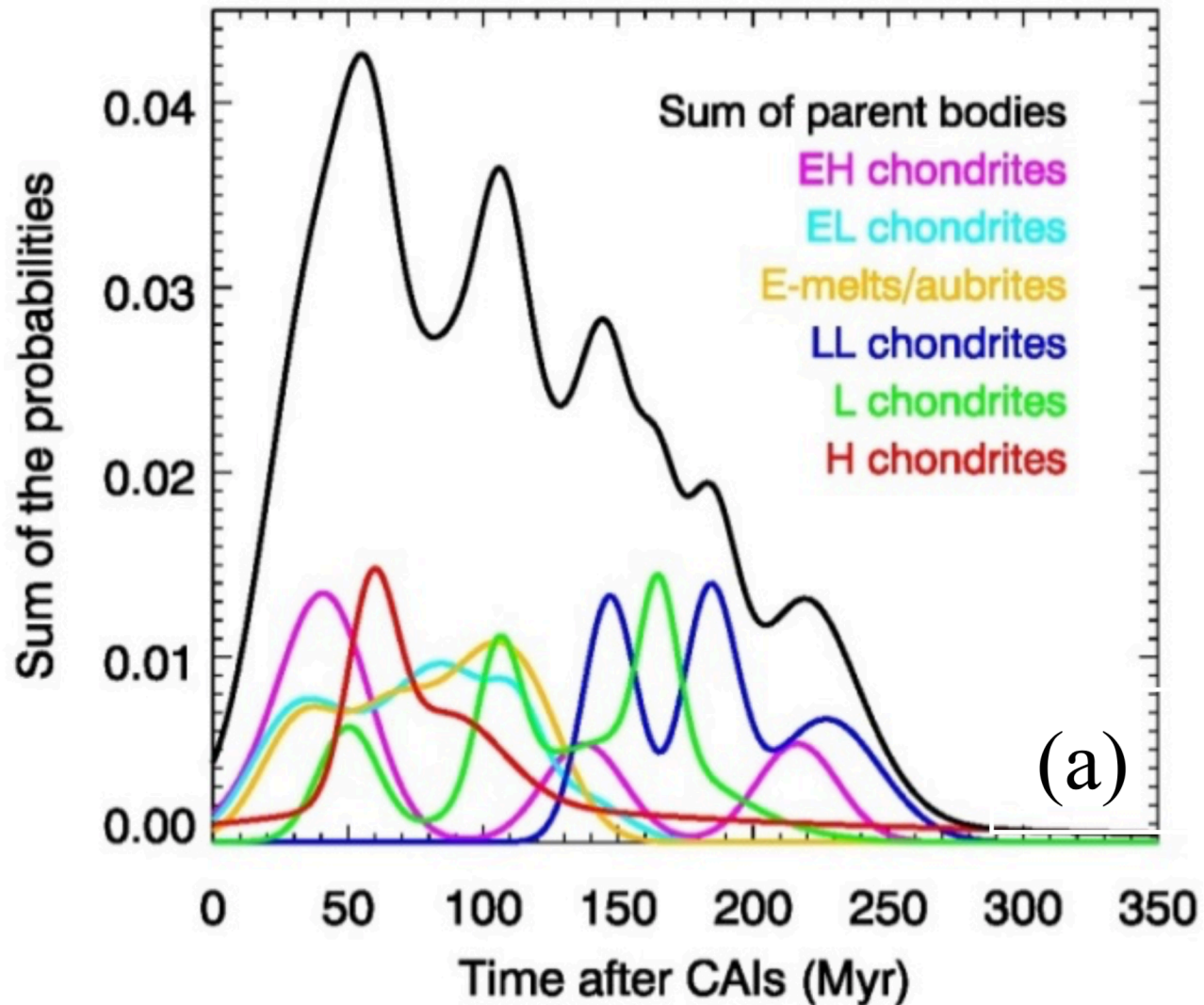
Hyodo, & Genda ApJL 2018

Distribution of Impact Debris within the inner Solar System



- Impacting asteroids would reset ^{40}Ar - ^{39}Ar age and/or cause impact melts
- Unmelted Martian mantle debris ($\sim 0.02\%$ of M_{Mars}) can be the source of
 - ☆ *Martian Trojan (Olivin-rich)*
 - ☆ *Rare A-type asteroids (Olivin-rich) in Hungarian and Main belt*

^{40}Ar - ^{39}Ar Resetting Age Distribution



Take Home Message

A giant impact on Mars can

- ▶ produce Phobos and Deimos (*Rosenblatt, ... Hyodo et al. 2016*)
- ▶ create the Borealis basin (*Marinova et al. 2008, Hyodo, et al. 2017b*)
- ▶ produce the current Mar's spin period (*Hyodo, et al. 2017b*)
- ▶ distribute debris as Martian Trojan & A-type asteroids (*Hyodo, & Genda 2018, ApJL*)

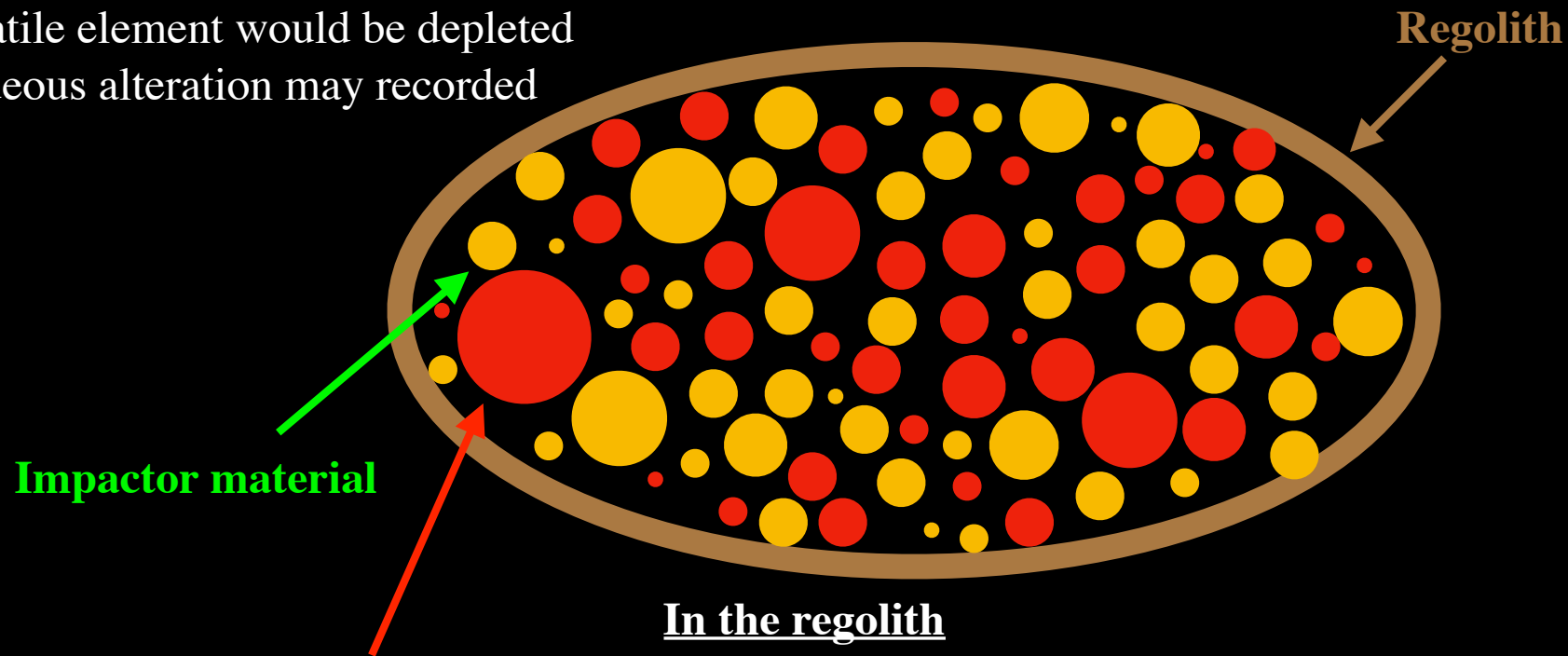
Building blocks of Phobos and Deimos:

- ▶ Mixture of impactor's and Martian materials
crust & mantle (*Hyodo, et al. 2017a*)
- ▶ Mixture of 0.1 μm and 100 μm – 1 m sized particles
Featureless & maybe dark (FeS, C) (*Pignatale, ... Hyodo et al. 2018*)
- ▶ Volatile loss is limited (*Hyodo, et al. ApJ submitted*)

Summary (Expected Phobos & Deimos)

Bulk composition (rubble-pile object):

- ~50% Martian material (at the time of impact: ~4 Gyr ago)
 - ◆ Martian crust and mantle (up to 150 km in depth)
- ~50% impactor material
- Particle size: 100 μ m - 10 m
- Volatile elements would be depleted
- Aqueous alteration may be recorded



Martian material

In the regolith

- Dust condensed from the vapor produced by giant impact
- Recently delivered material
 - ◆ Martian surface material
 - * ~150 ppm delivered within recent 10 Myr
 - * ~1500 ppm delivered within recent 500 Myr
 - ◆ Impactor material