

# Analyzing the Instabilities in the Venus Atmosphere Using Bred Vectors

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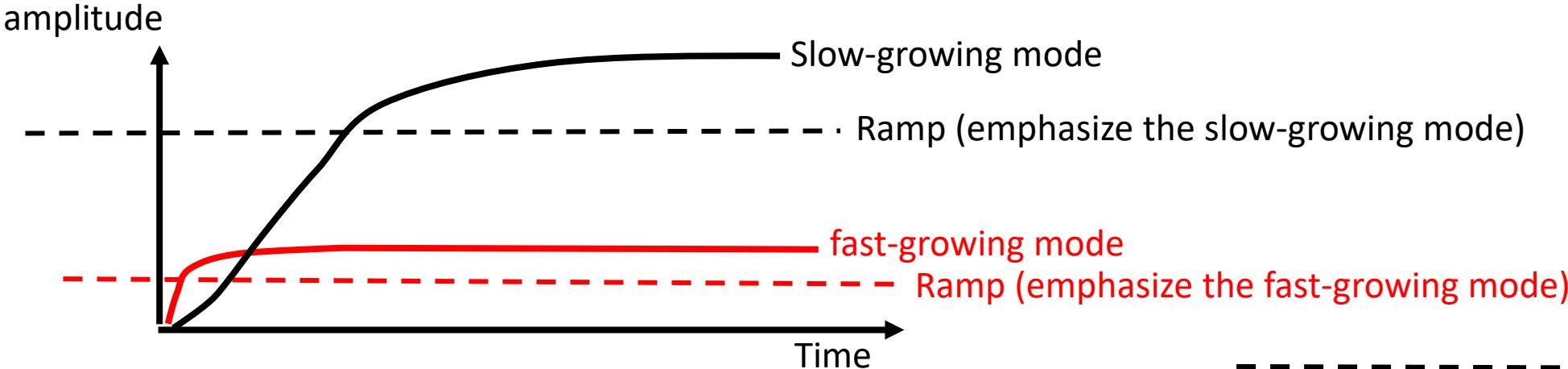
**Workshop on Venus and other related atmospheres**

# Overview

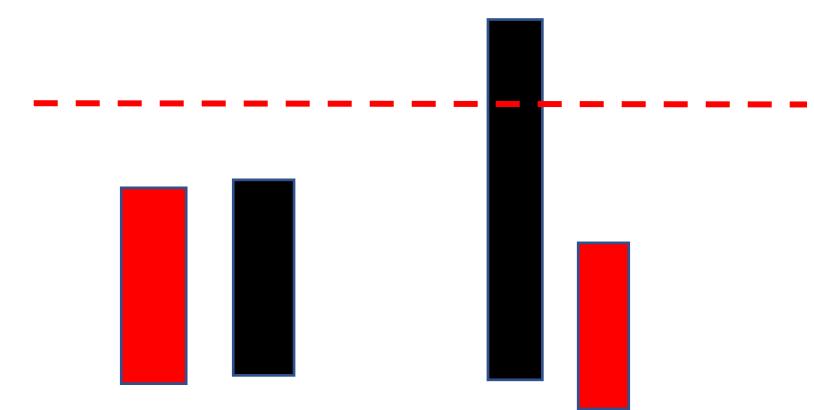
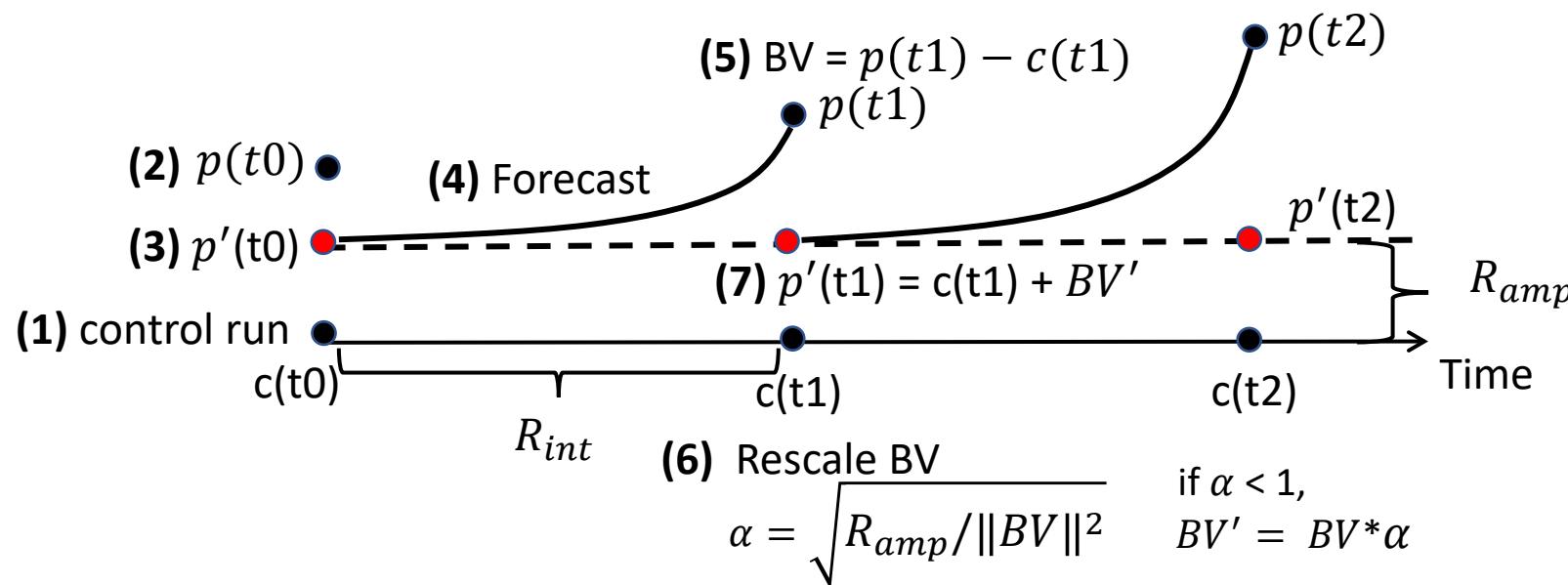
- What is Bred vector?
- Overview of instabilities by previous studies
- Bred vector of the Venus atmosphere
  - ❖ compare deviation from the zonal mean and Bred Vectors
  - ❖ Bred Vector identify different growing modes
  - ❖ baroclinic instability impact on thermal tide
- Predictability

## Breeding cycle and Bred Vector (BV)

(a)



(b)

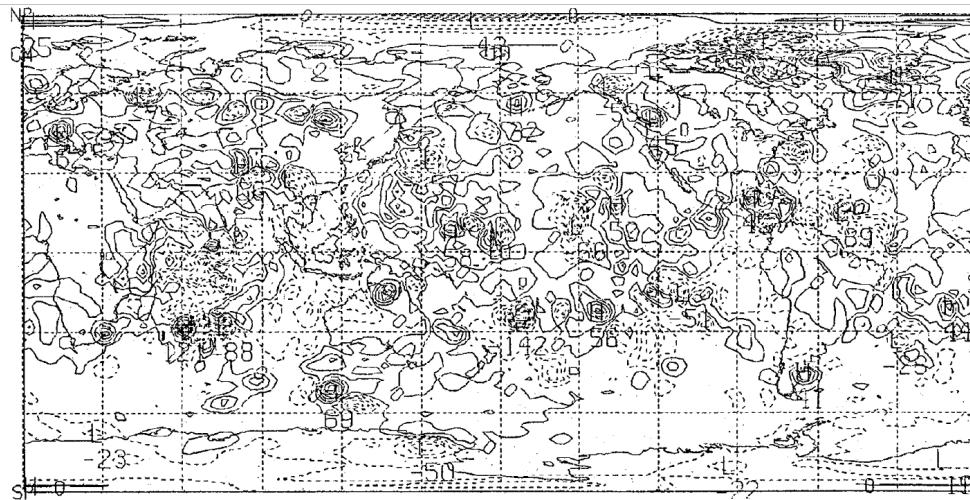


## Applications of BVs

- Generating ensemble perturbations (before EnKF method is applied in Meteorology)
- Analyzing the **unstable** and growing modes of the dynamical system
- Estimating the effective local **dimension** (Patil et al. 2001) (prepare for LETKF)
- Studied **energy conversion** (Hoffman 2009; Greybush 2013)
- Applied to **Earth** (Toth and Kalnay 1993, 1997) and **Mars** atmospheres (Newman 2004; Greybush 2013)

500-hPa streamfunction perturbations

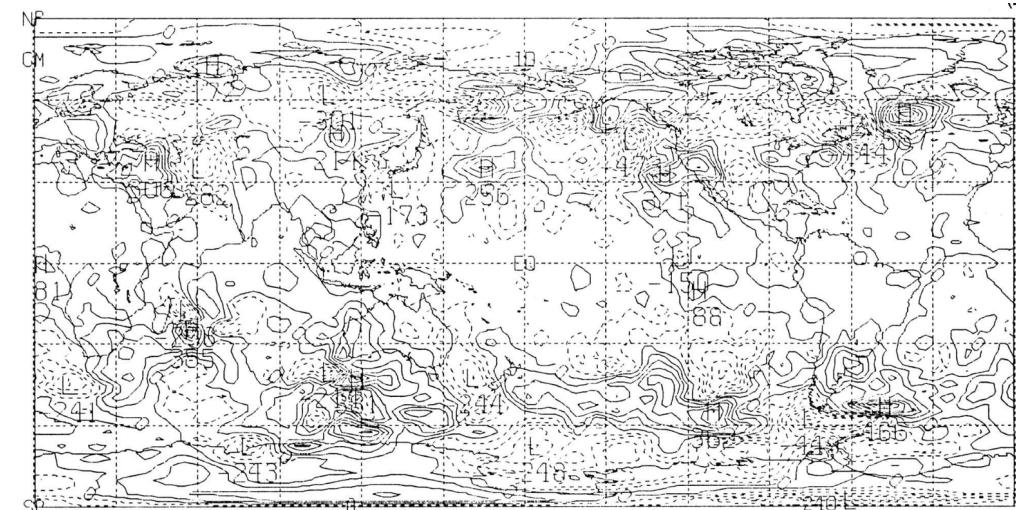
Convective modes



Toth and Kalnay 1997

Size of the perturbation: 0.015% total rms variance

Baroclinic modes



Toth and Kalnay 1993

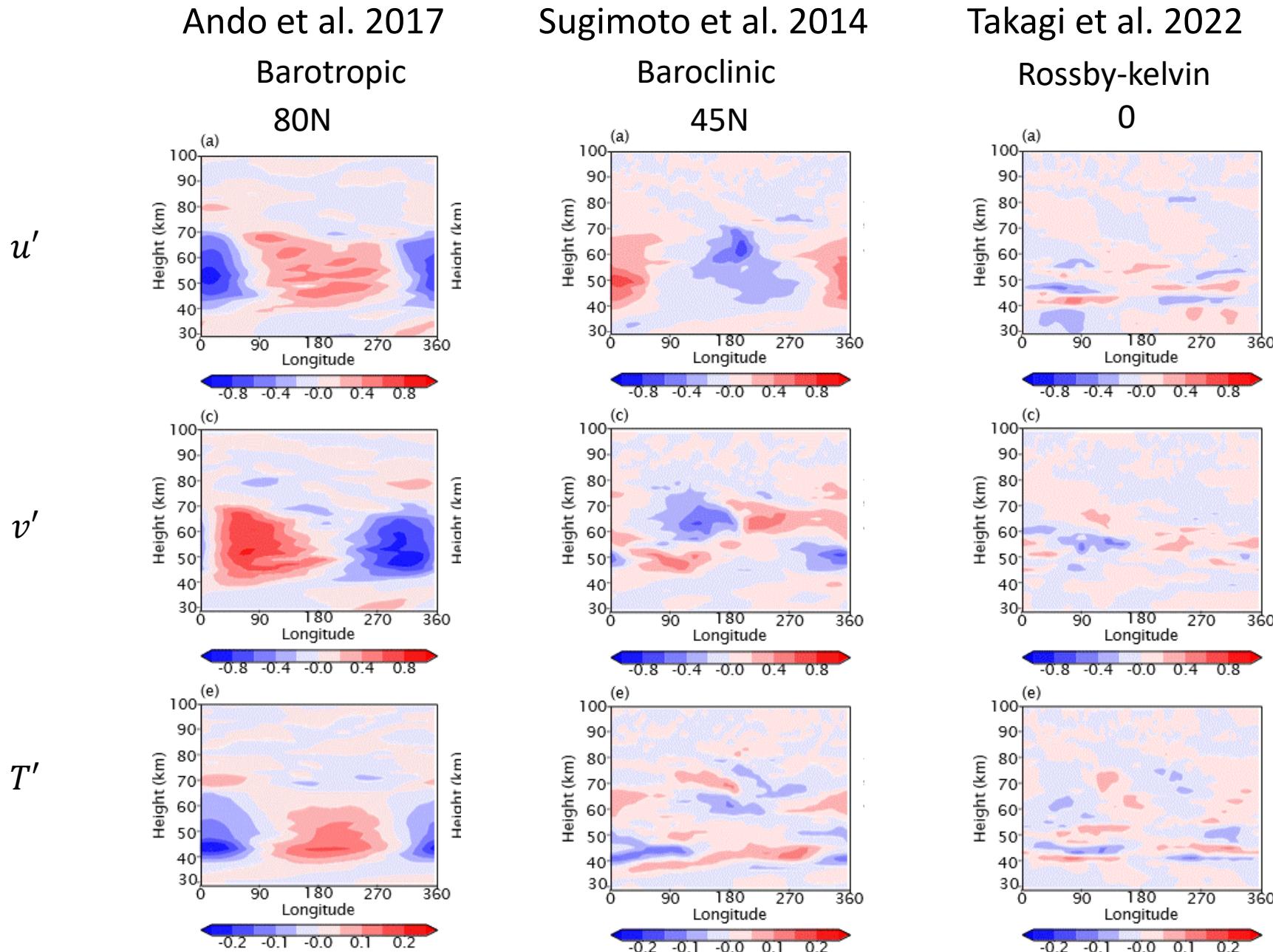
Size of the perturbation: >1% total rms variance

## Experimental setting

- AFES-Venus (VAFES) (Sugimoto et al., 2014 JGR, GRL)
  - ✓ 3-D Primitive equation on sphere (hydro static balance) without moist processes
  - ✓ Resolution: T42L60 ( $128 \times 64 \times 60$ )
  - ✓ Horizontal hyper-viscosity: 0.1 Earth days for  $1/e$
  - ✓ Vertical eddy viscosity:  $0.15 \text{ m}^2\text{s}^{-1}$
  - ✓ Rayleigh friction: lowest and above 80 km (sponge layer except for zonal flow)
  - ✓ No topography and planetary boundary layer
- Solar heating
  - ✓ Zonal ( $Q_z$ ) and diurnal ( $Q_t$ ) component of realistic heating; Based on Tomasko et al. (1980) and Crisp (1986)
- Infrared radiative process
  - ✓ Simplified by Newtonian cooling:  $dT/dt = -\kappa (T - T_{ref}(z))$     $\kappa$ : based on Crisp (1986)    $T_{ref}(z)$ : horizontally uniform field

Cases	Solar heating
$Q_z$	Zonal mean component only
$Q_t$	Including diurnal variation

# Deviation from the zonal mean



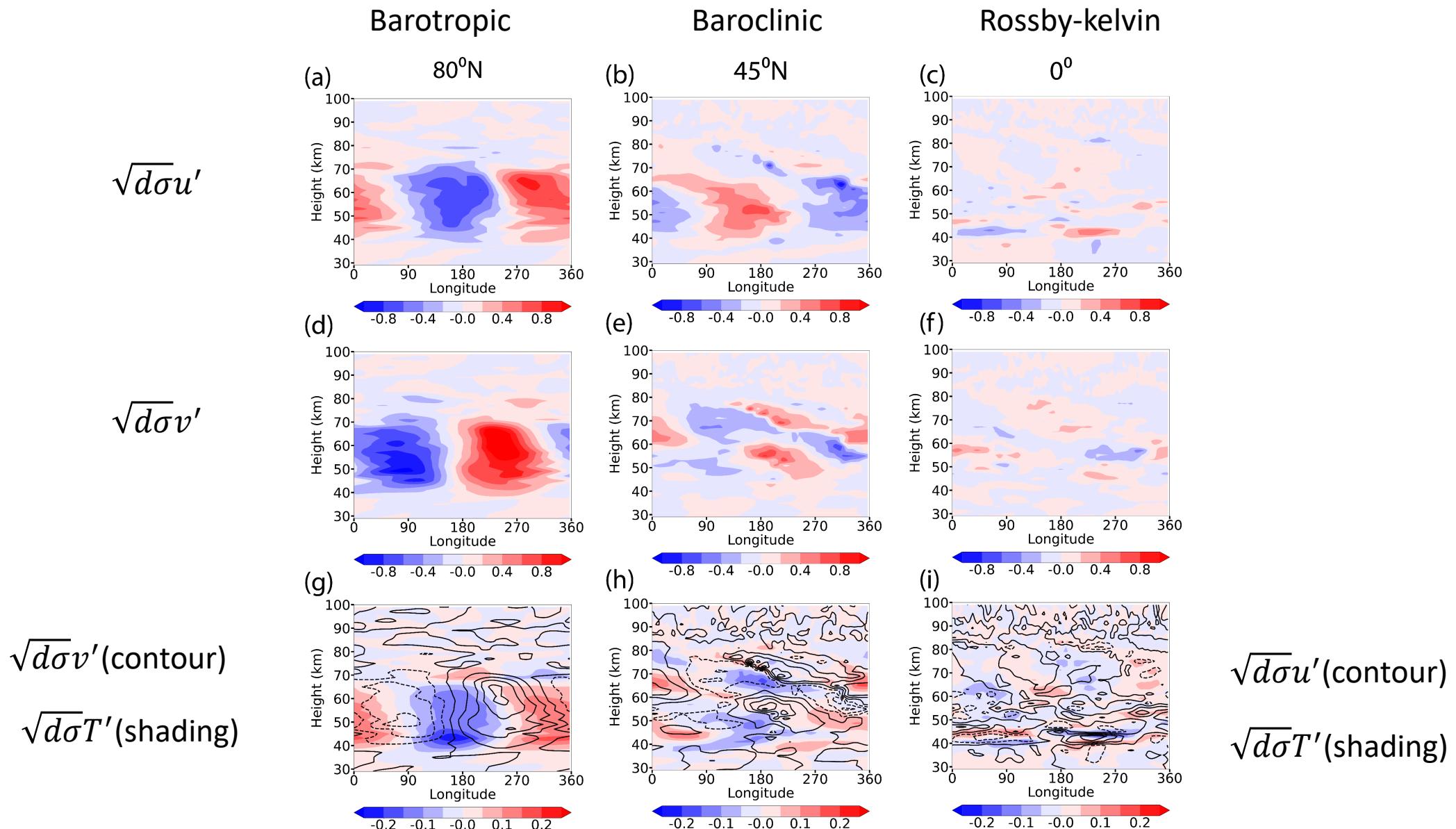
# Deviation from the zonal mean

at 0005/06/03/18

Ando et al. 2017

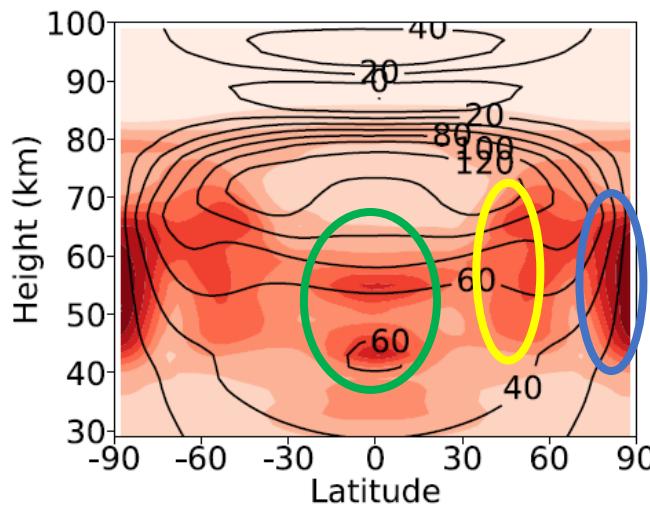
Sugimoto et al. 2014

Takagi et al. 2022

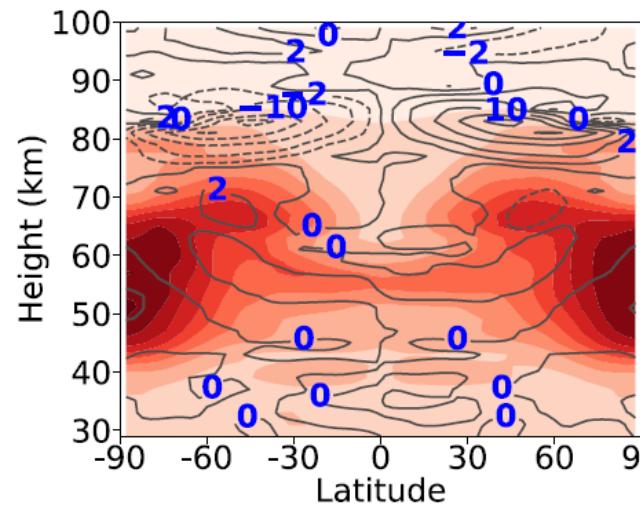


# average fields and the magnitude of the deviation

(a)



(b)



$$\sqrt{d\sigma} \sqrt{v'^2}$$

Barotropic



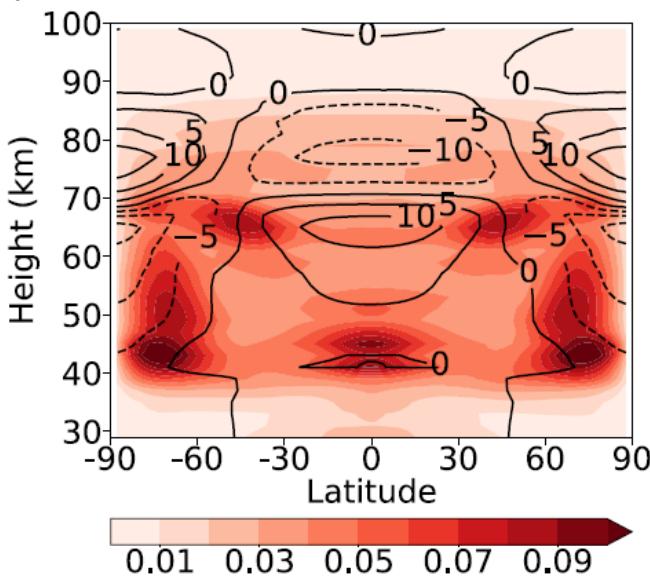
Baroclinic



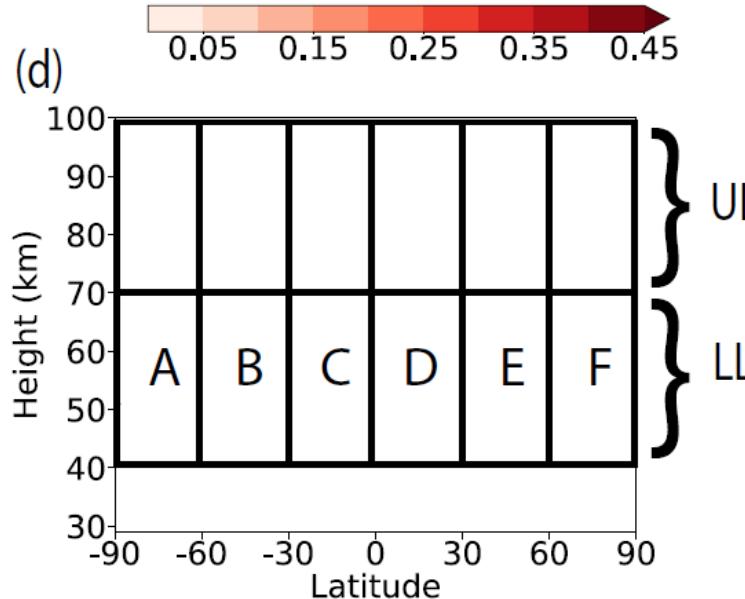
Rossby-kelvin



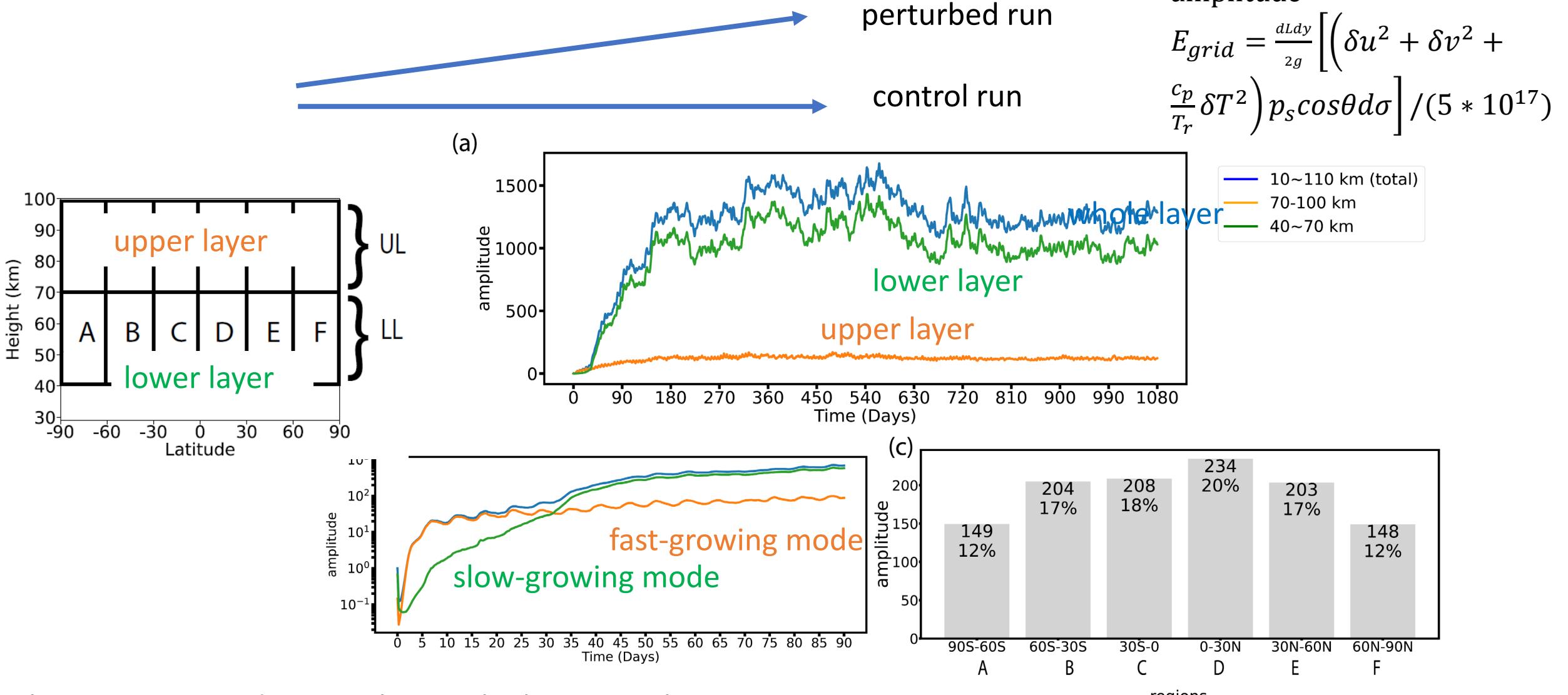
(c)



(d)



## Identical twin experiments



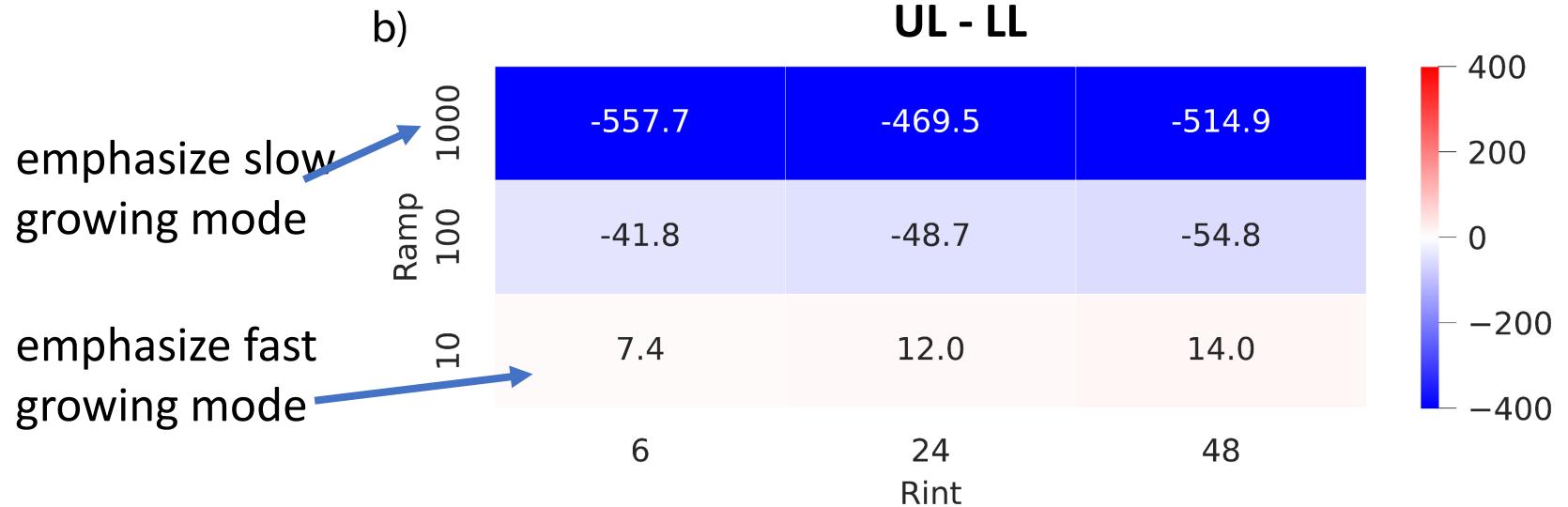
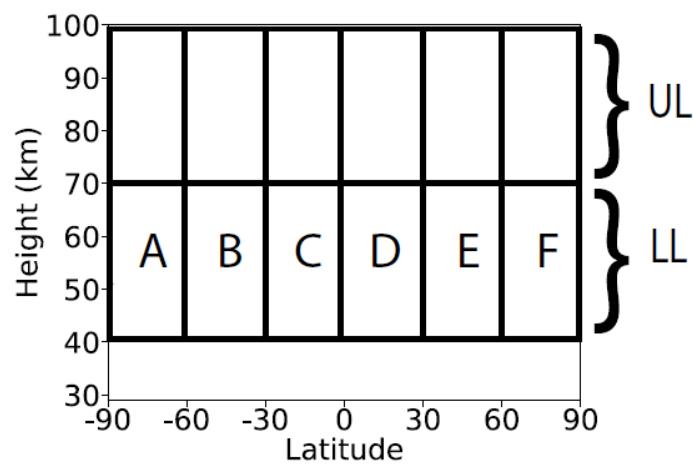
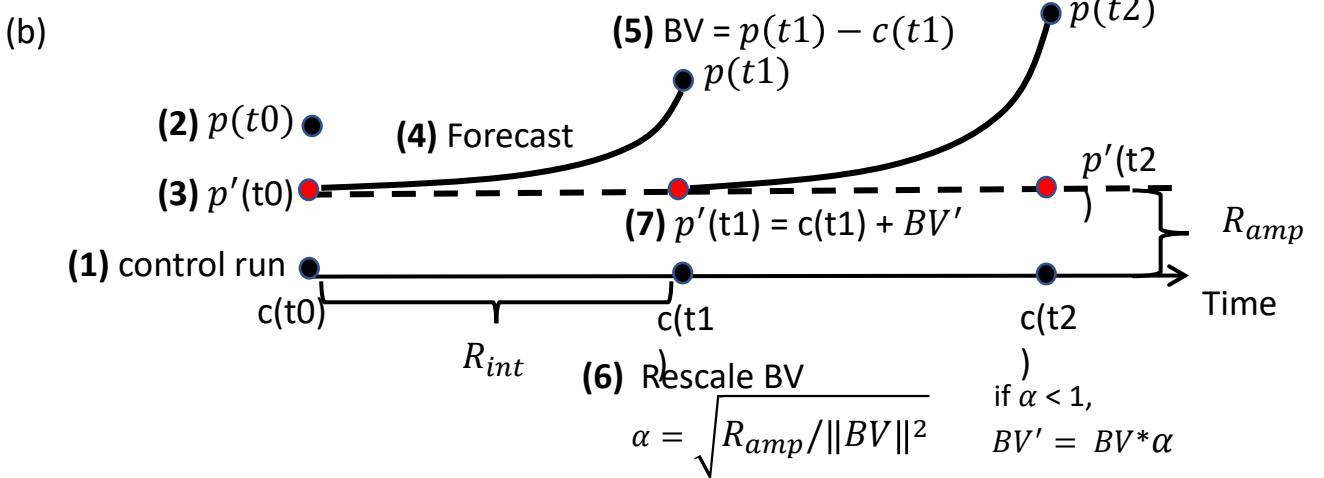
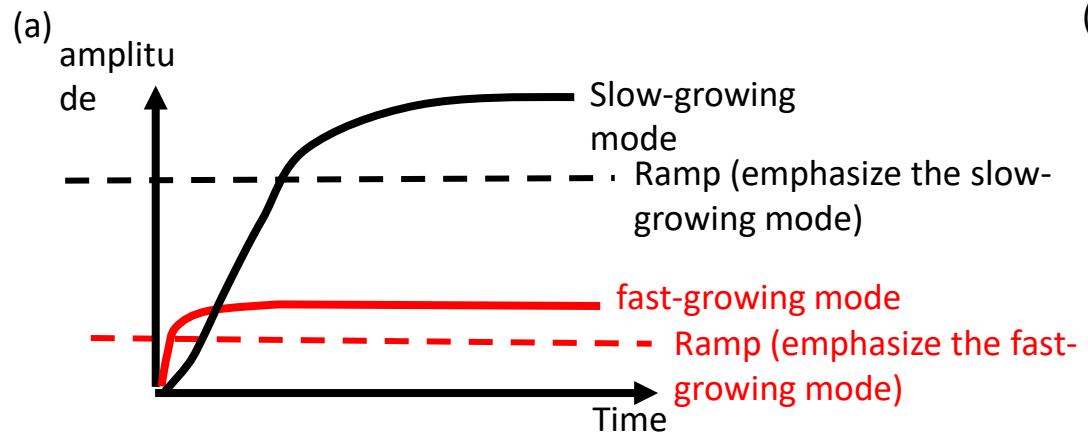
slow-growing mode grow slow at the beginning but  
saturate at higher amplitude

amplitude in the lower latitudes are  
comparable to those in mid-latitudes

use total energy to measure the  
amplitude

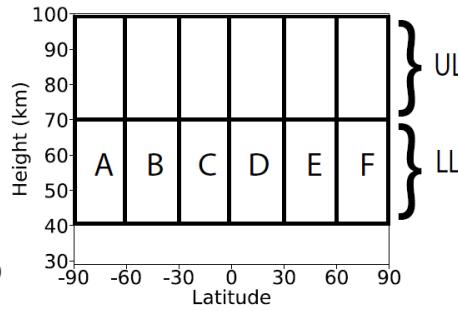
$$E_{grid} = \frac{dLdy}{2g} \left[ \left( \delta u^2 + \delta v^2 + \frac{c_p}{T_r} \delta T^2 \right) p_s \cos\theta d\sigma \right] / (5 * 10^{17})$$

# control BV by different rescaling amplitudes and rescaling intervals

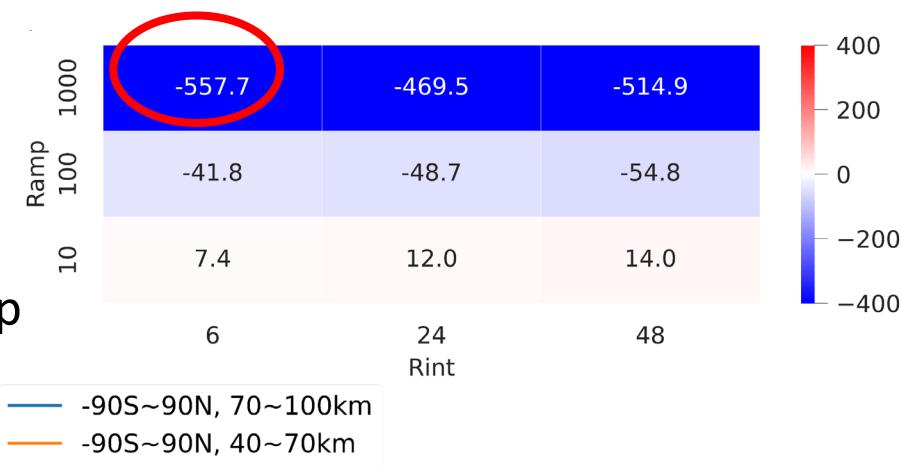
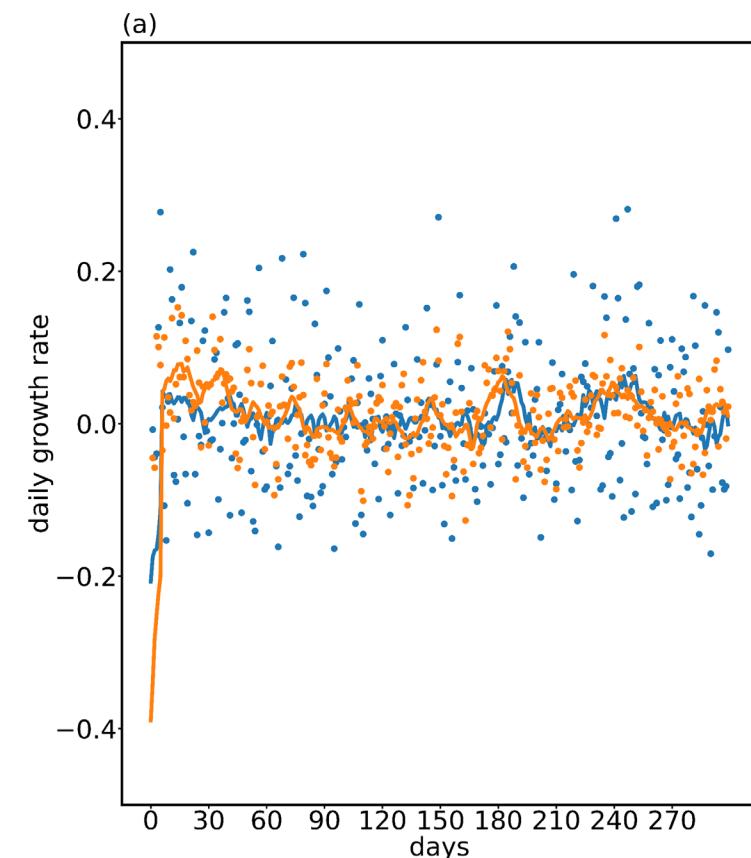


# emphasize slow growing mode

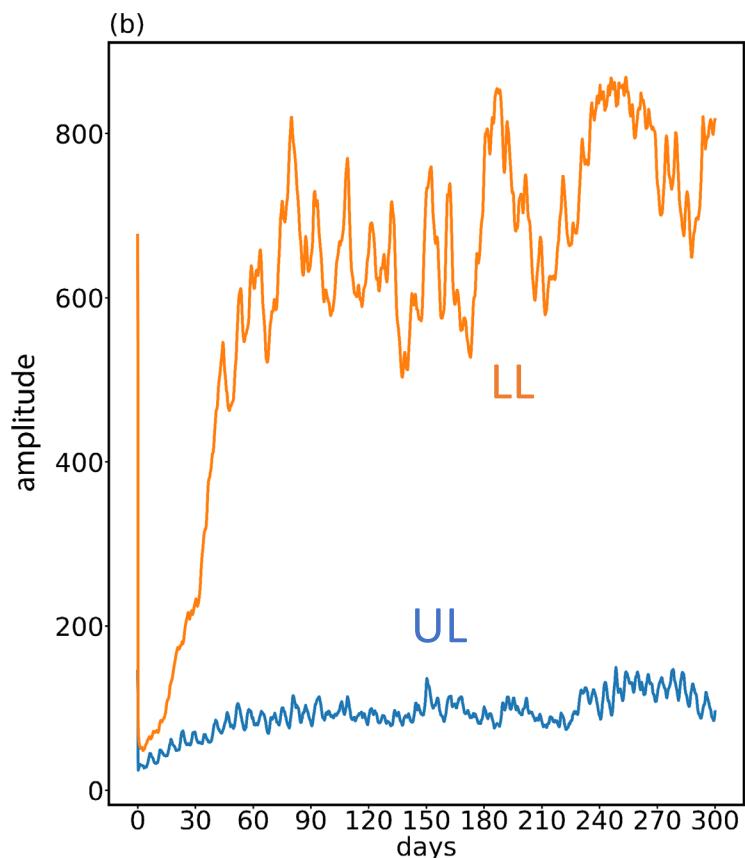
Ramp = 1000



growth rate

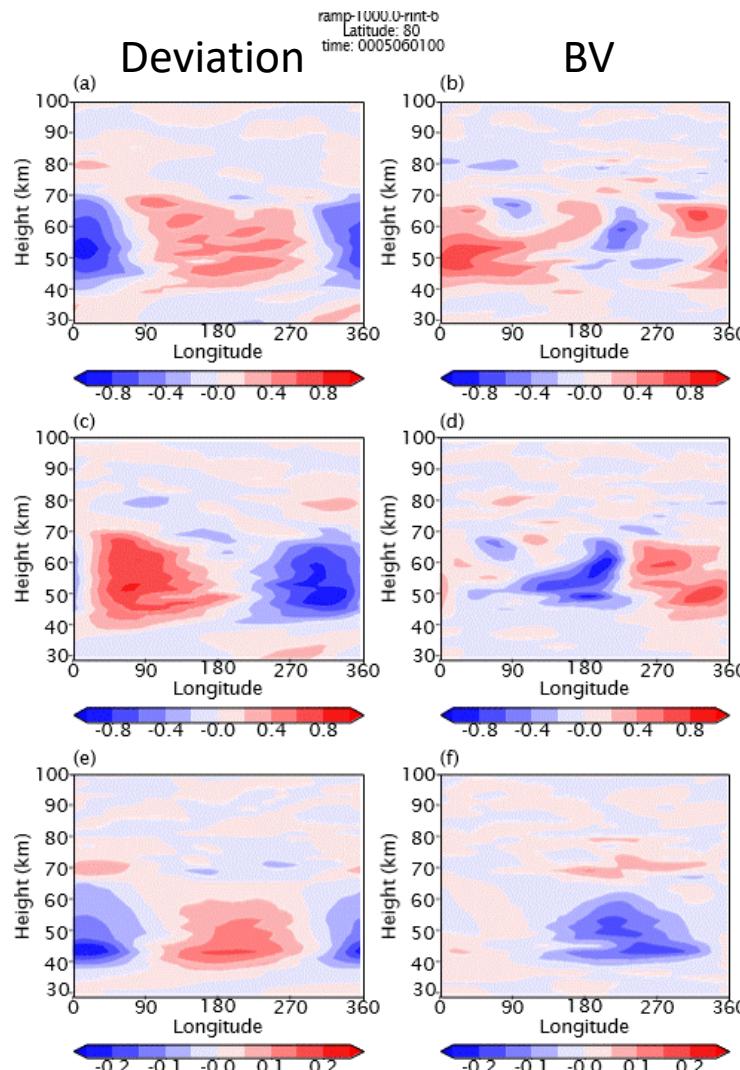


slow growing mode dominate

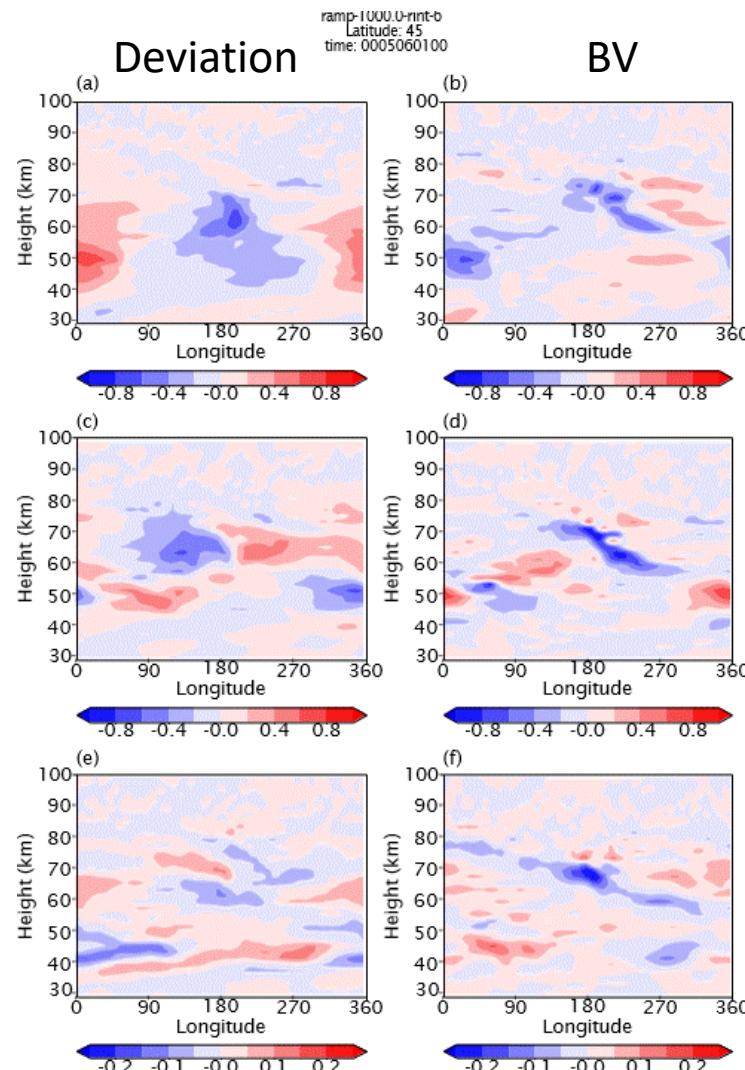


# compare deviation and BV

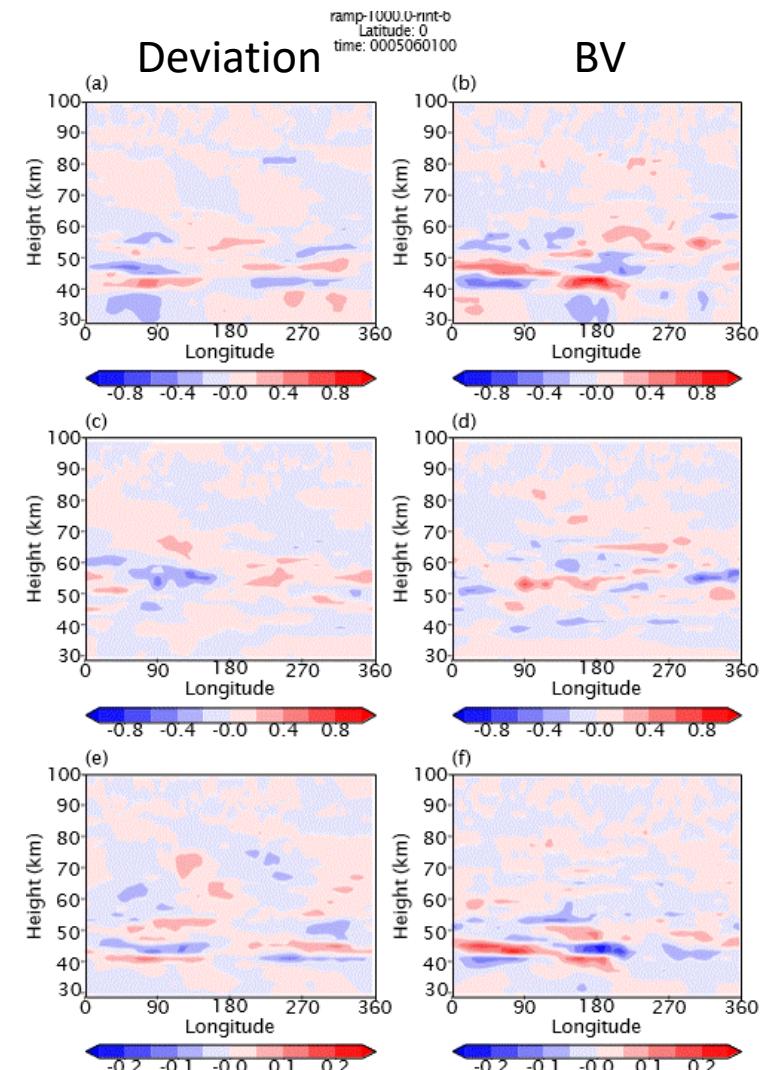
80N  
Barotropic



45N  
Baroclinic

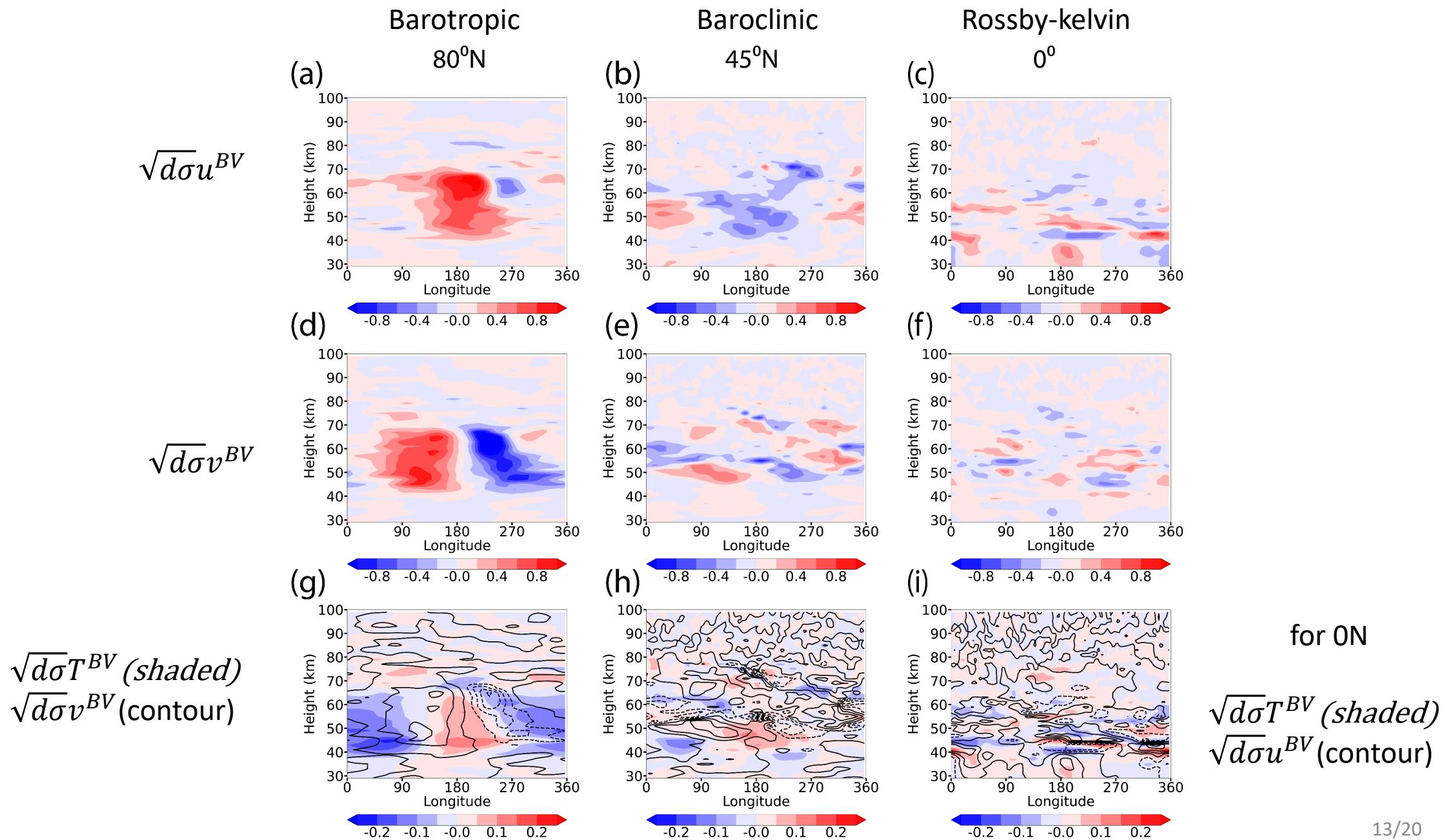


0  
Rossby-kelvin



- the structures are similar
- the zonal phases are different

## Bred Vectors at different latitudes



# Composite mean at different latitudes

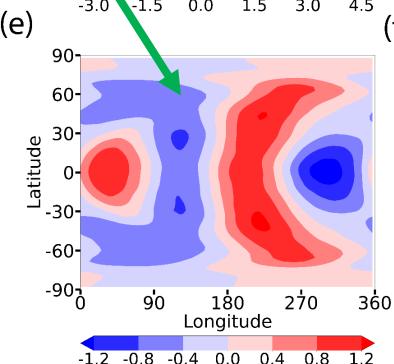
thermal tide increase  
baroclinic instability

$$\sqrt{d\sigma} \bar{u}'^{\text{comp}}$$

temperature  
gradient is  
increased

$$\bar{T}'^{\text{comp}}$$

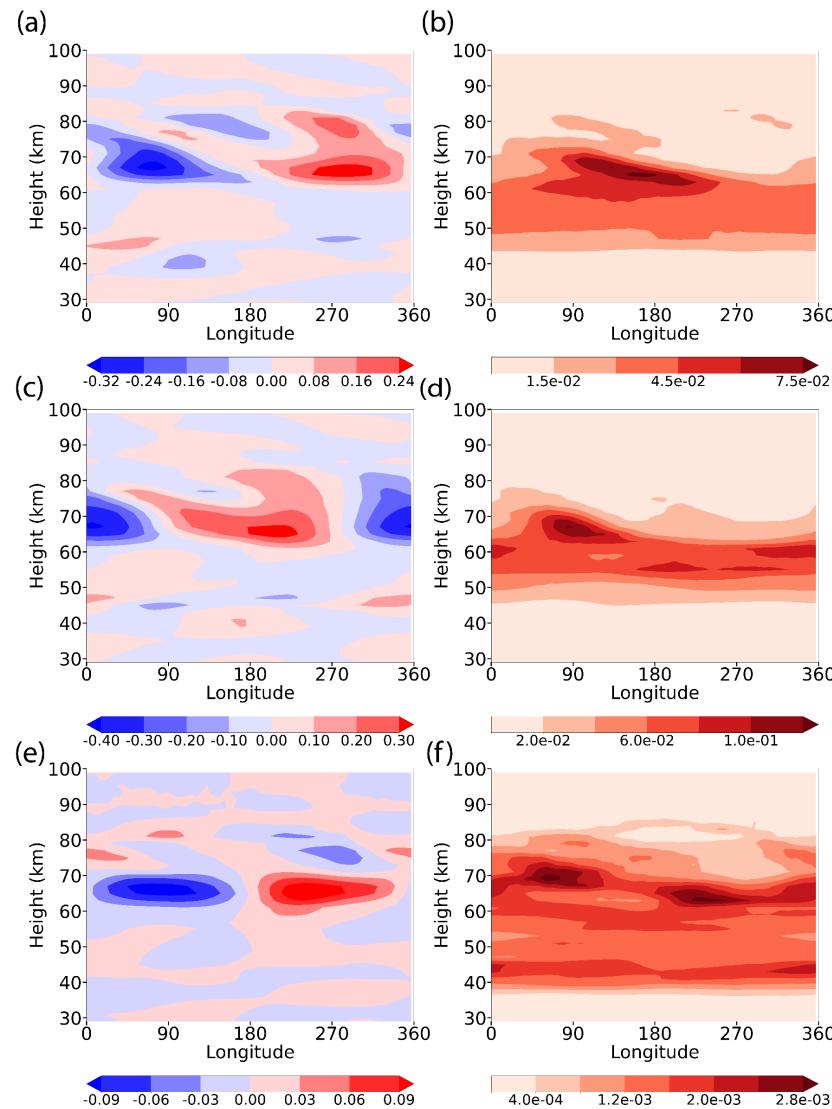
60 km



$$\sqrt{d\sigma} \bar{T}'^{\text{comp}}$$

control

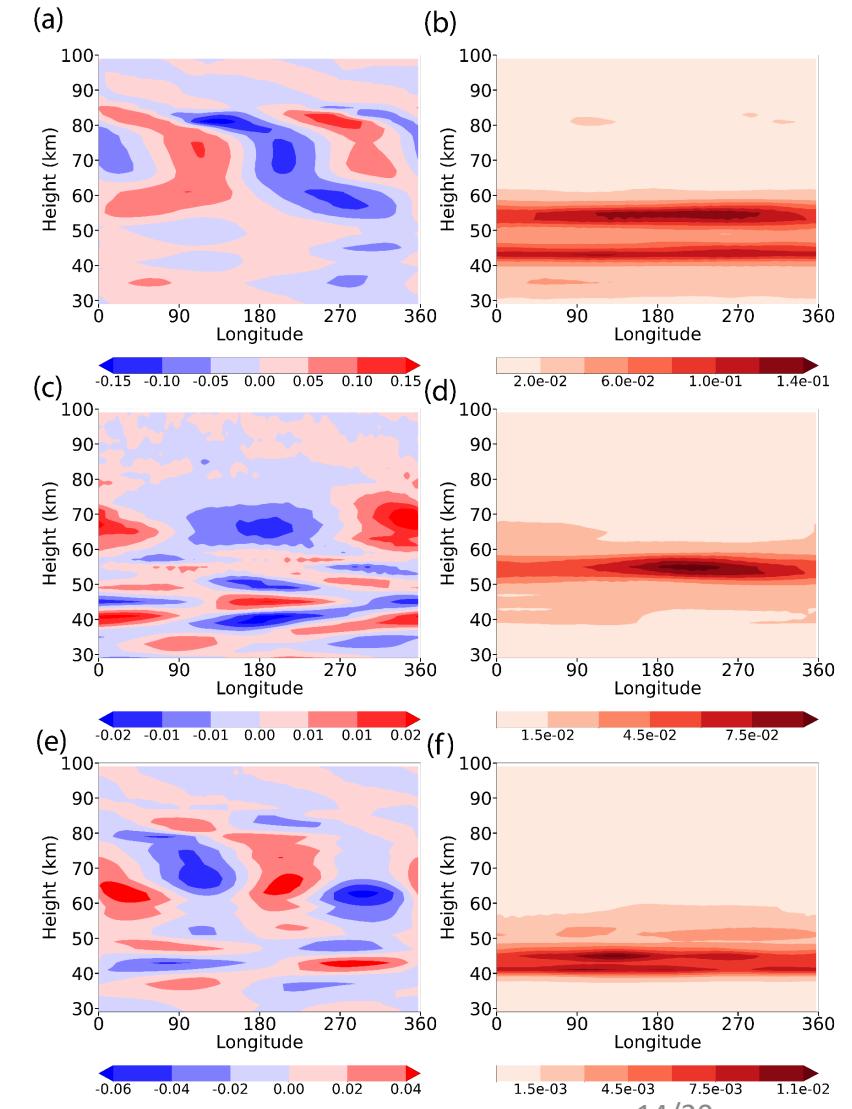
45 N



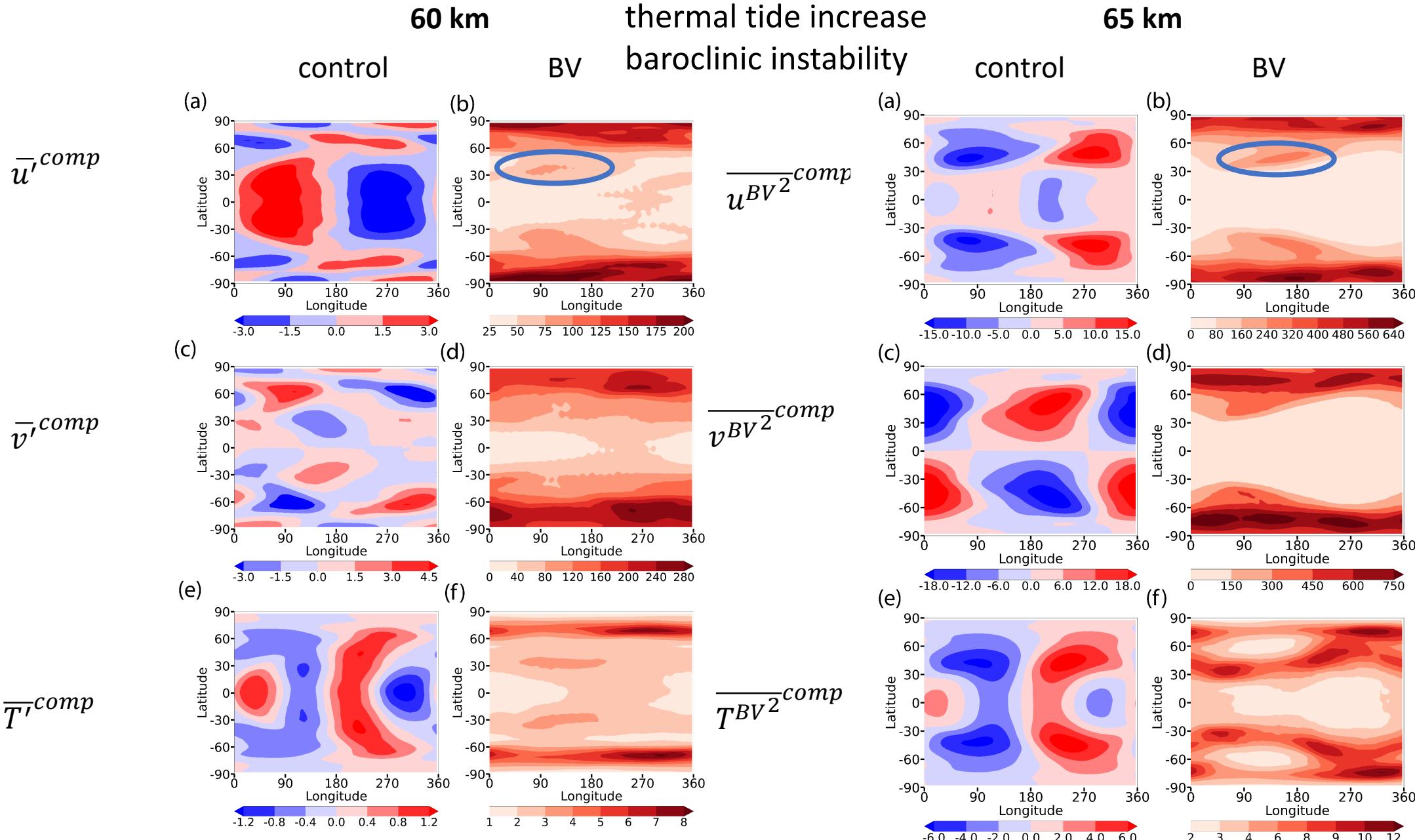
control

0 N

BV



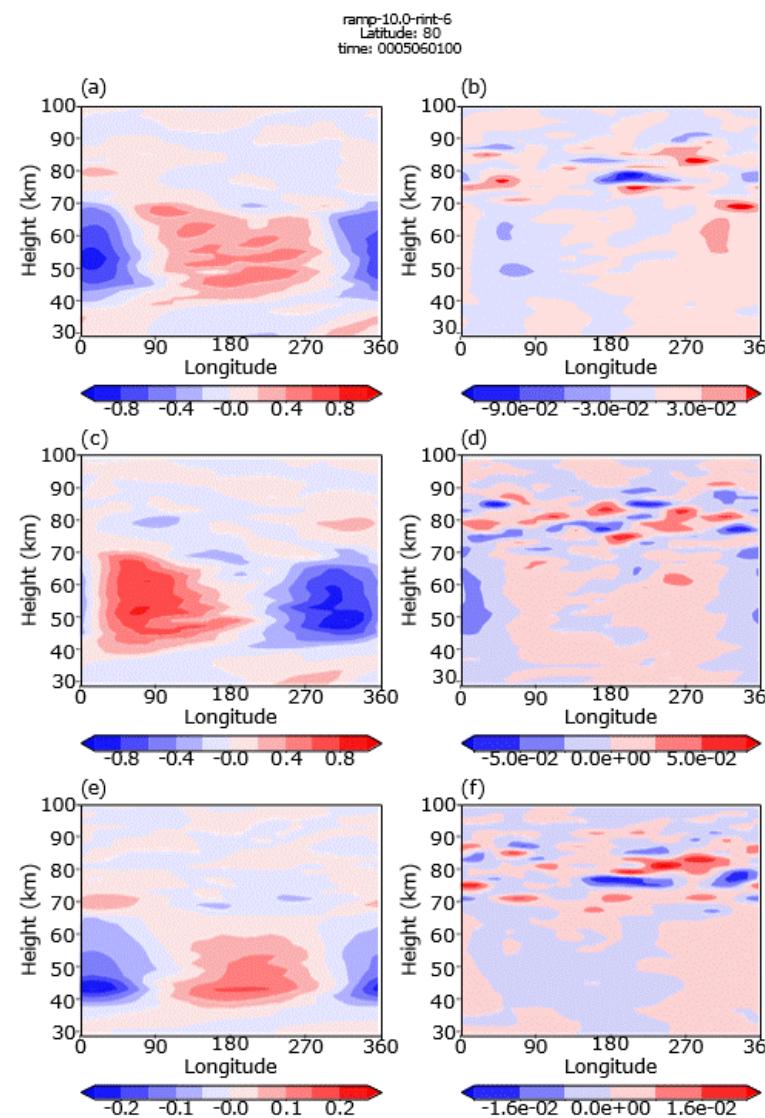
# Composite mean at different height



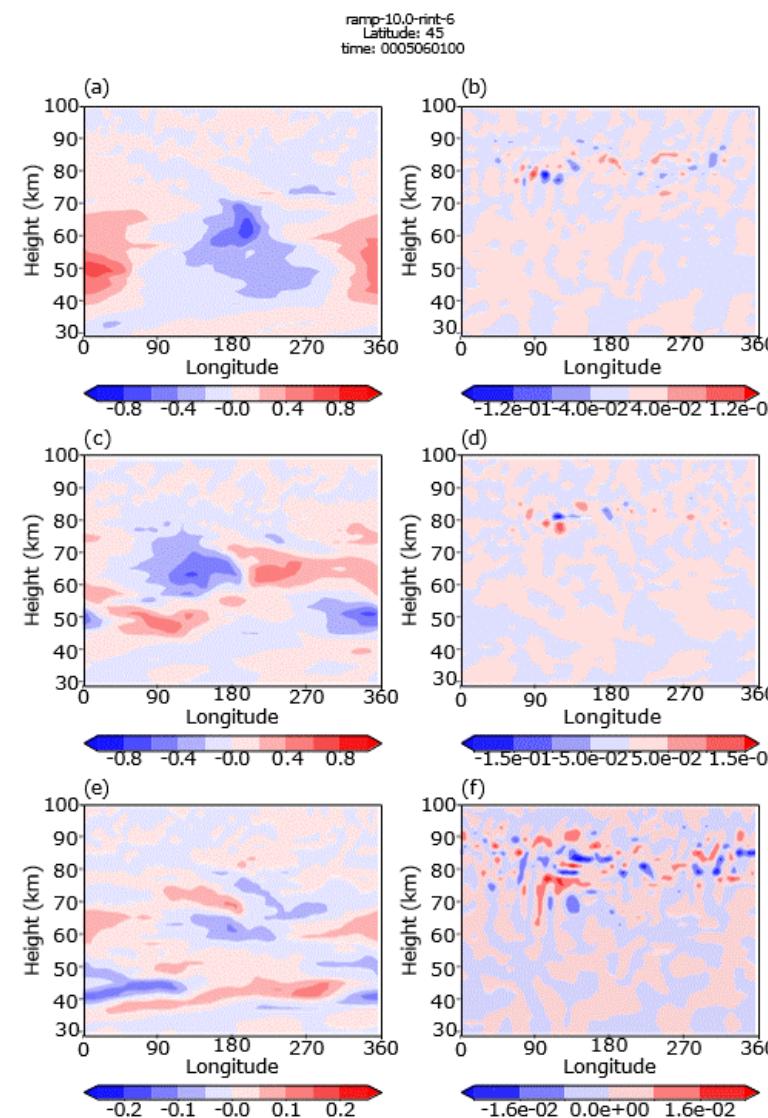
# emphasize fast growing mode

Ramp10-Rint06

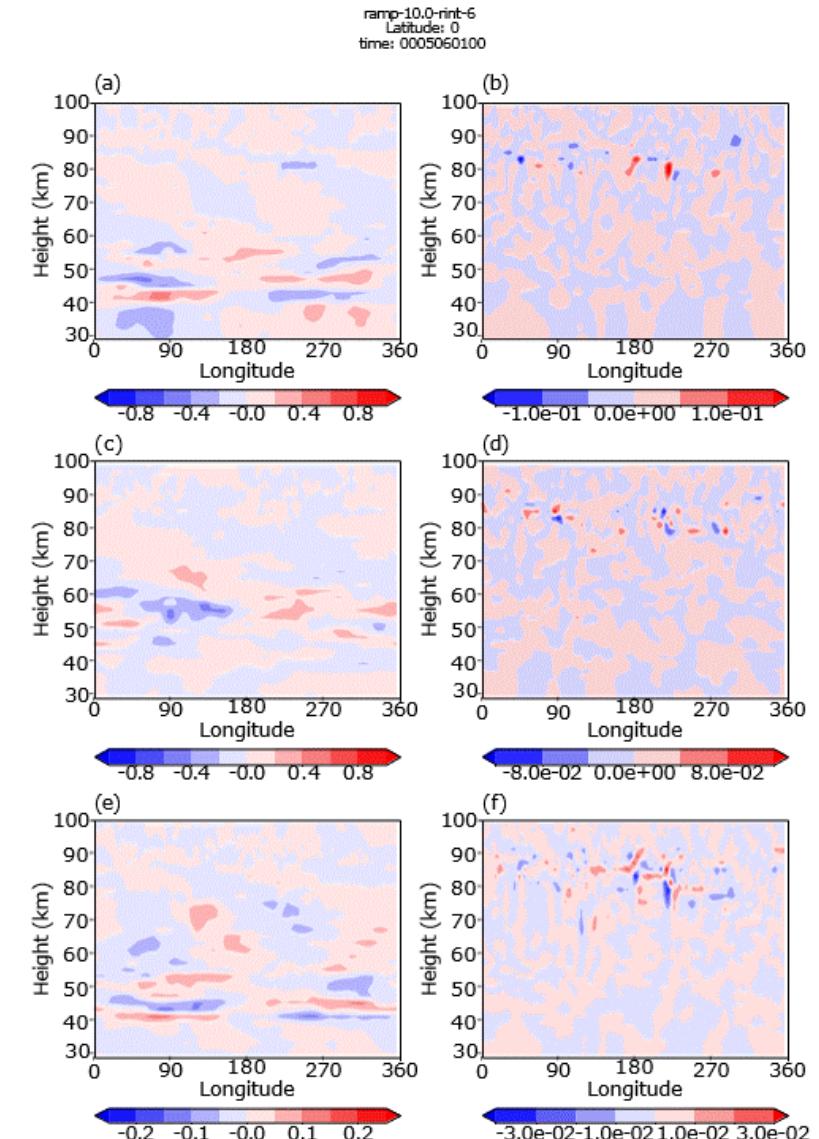
80N



45N



0



# gravity waves generation and breaking

80 km

control

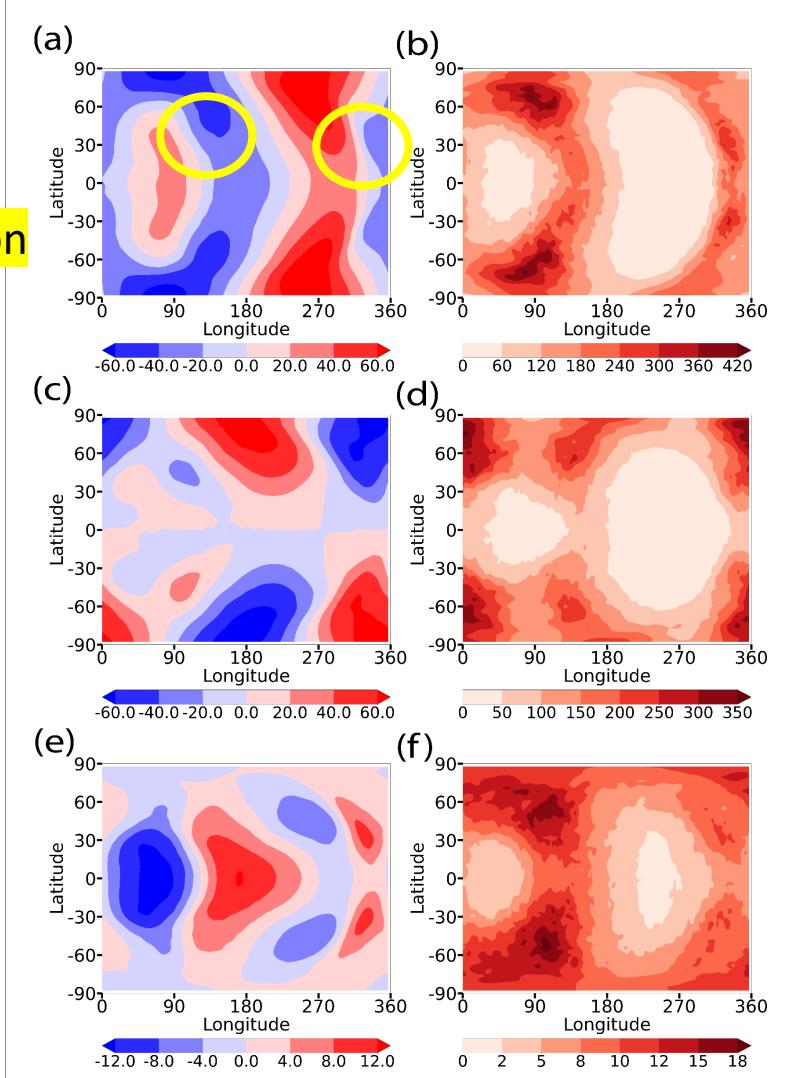
BV

$\bar{u}'^{\text{comp}}$

Jet exit region

$\bar{v}'^{\text{comp}}$

$\bar{T}'^{\text{comp}}$



$\bar{u}^{BV^2 \text{ comp}}$

$\bar{u}^{BV^2 \text{ comp}}$

$\bar{u}^{BV^2 \text{ comp}}$

(a)

45°N

(b)

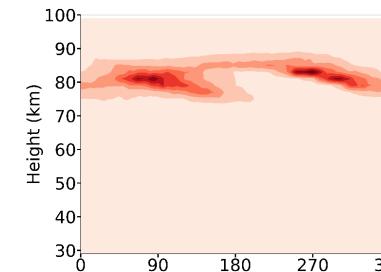
0°

(c)

(d)

(e)

(f)

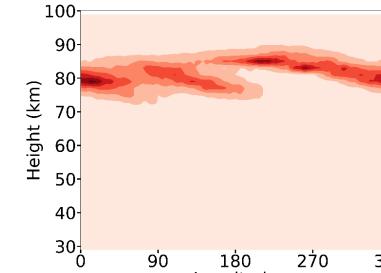


BV

(b)

(d)

(f)

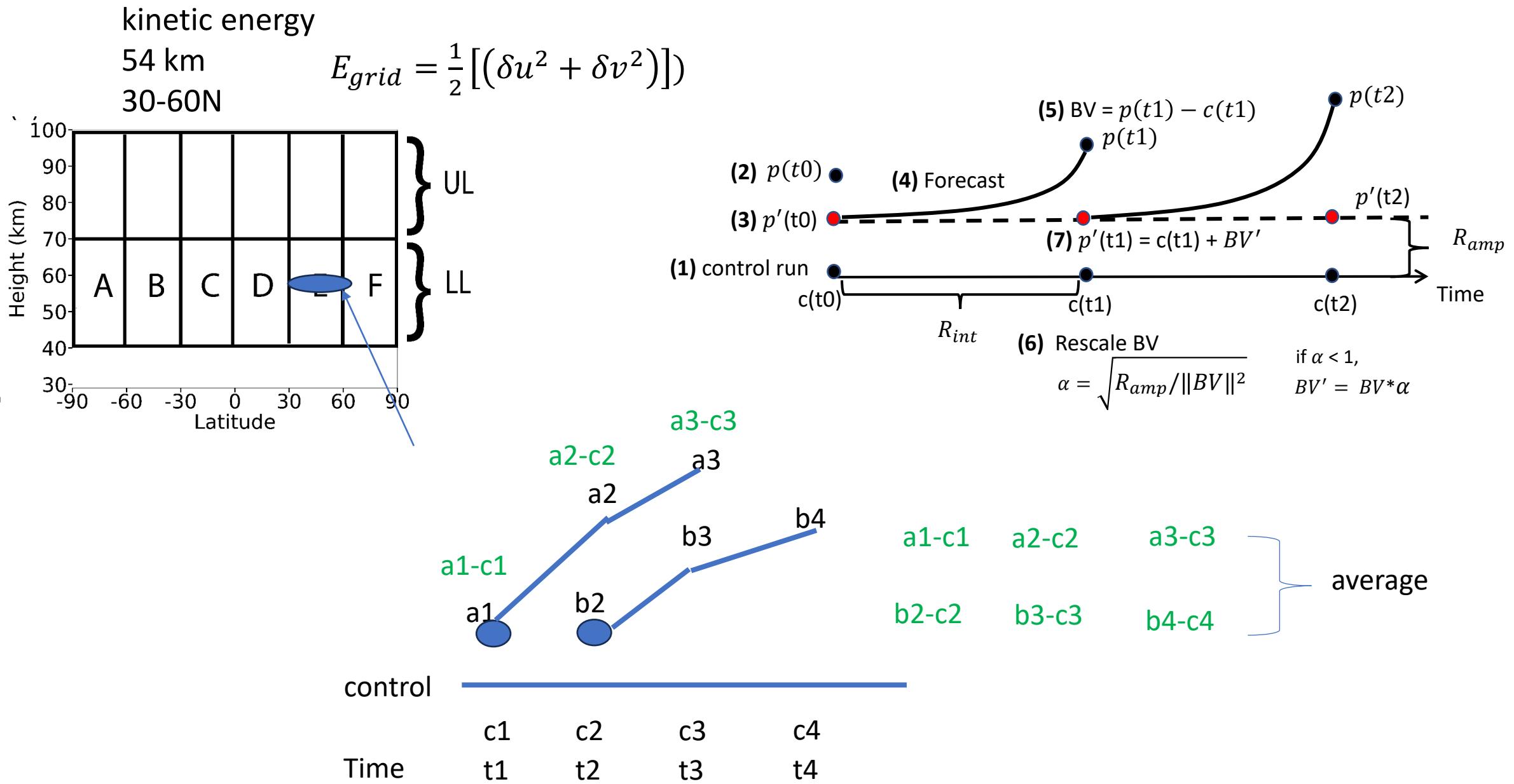


$d\sigma \bar{u}^{BV^2 \text{ comp}}$

$d\sigma \bar{u}^{BV^2 \text{ comp}}$

$d\sigma \bar{u}^{BV^2 \text{ comp}}$

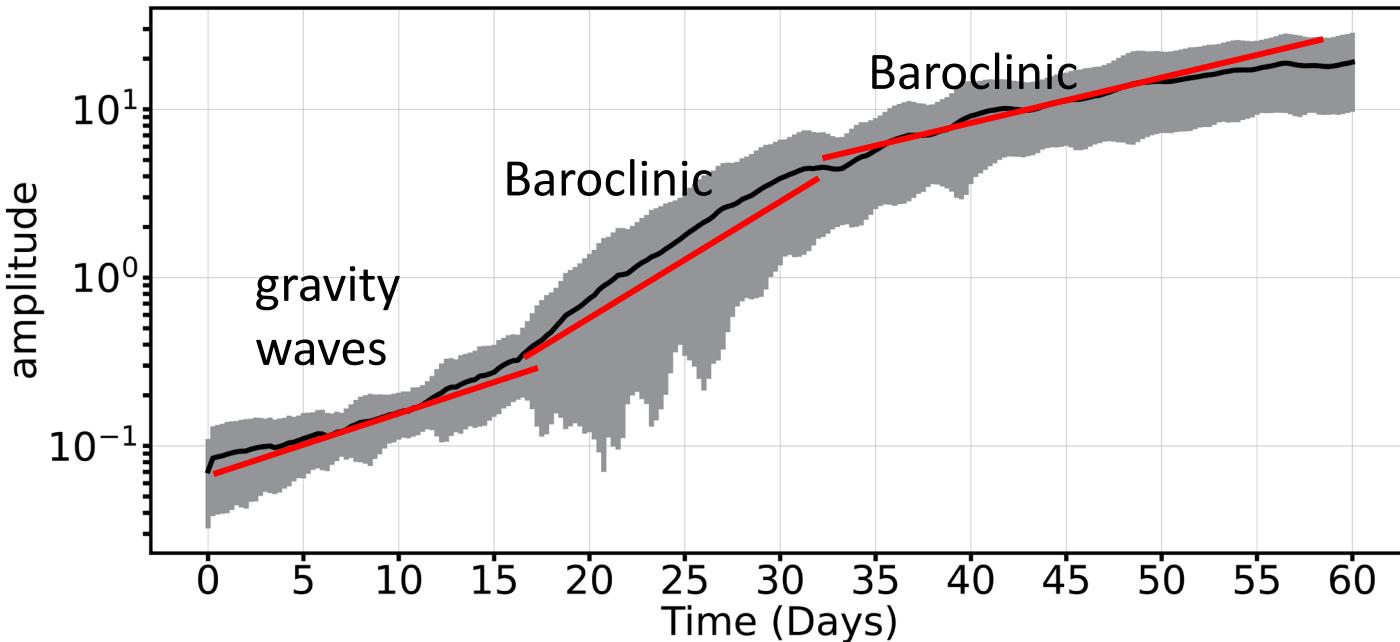
# Predictability



# Predictability

54 km,  
30-60N

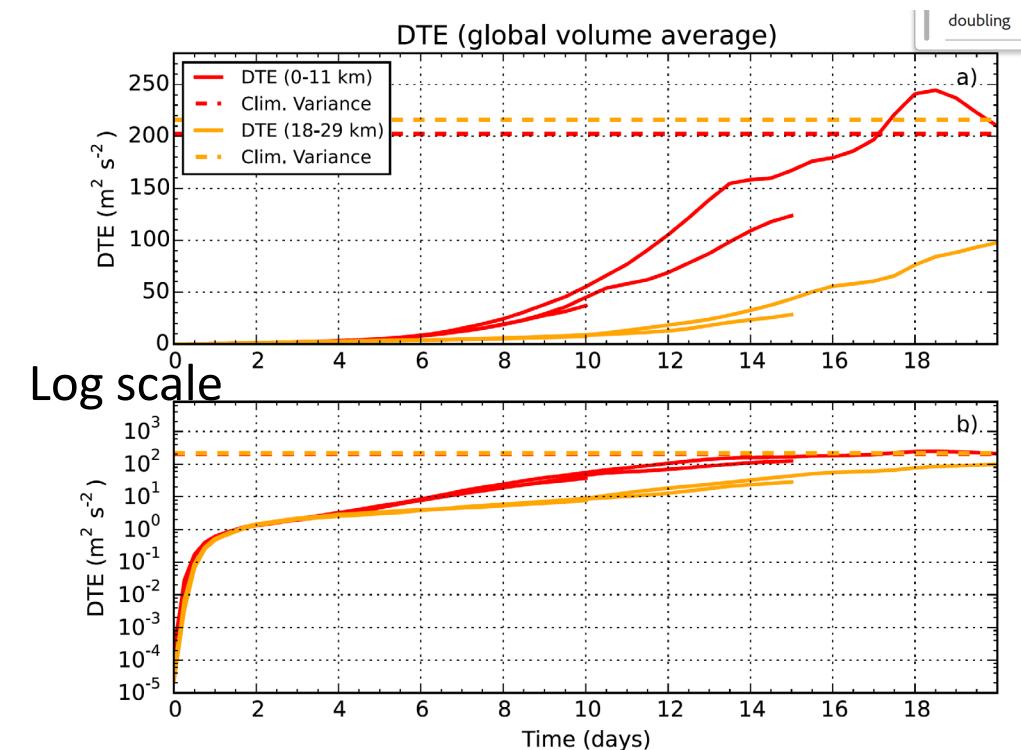
Predictability: > 1 month



days	daily growth rate	error doubling time (days)	e folding time (days)
0-to-17	0.11	6.6	9.5
17-to-32	0.15	4.6	<b>6.6</b>
32-to-60	0.05	13.8	19.9

Earth atmosphere

Predictability: 2-3 weeks



Judy  
2018

## Summary

- Identical twin experiments show that perturbations at 40-70 (cloud layer) and 70-100 km (above cloud layer) altitudes are **slow-growing** and **fast-growing** modes, respectively.
- Breeding cycle experiments generate perturbations associated with **barotropic**, **baroclinic**, **Rossby-Kelvin** instabilities in the lower layer, and the **generation and breaking of the gravity waves** in the upper layer.
- BVs show that **thermal tide** enhances the baroclinic instability
- The intrinsic **predictability** related to the baroclinic instability is >1 month, longer than that for the Earth's atmosphere of ~2 weeks.

## Future works

- study **predictability** in different regions, different model resolution, different models
- use **Bred vector energy equation** to identify the origin of the instability
- use data assimilation to **estimate the model parameters**: solar heating parameters, etc.

Liang, J., N. Sugimoto, and T. Miyoshi, 2024: Analyzing the Instabilities in the Venus Atmosphere Using Bred Vectors. *Journal of Geophysical Research: Planets*, in press





## Definition of the norm of BV

$$E = \left\{ \sum \sum \sum \left[ \left( \delta u^2 + \delta v^2 + \frac{c_p}{T_r} \delta T^2 \right) p_s d\sigma \cos \theta \right] \right\} / 10^8,$$

- where  $\delta u$  ( $m s^{-1}$ ),  $\delta v$  ( $m s^{-1}$ ), and  $\delta T$  (K) are the zonal wind, meridional wind, and temperature in the perturbations, respectively.
- $dL$  (m) is the longitudinal distance of a model grid at the equator.
- $dy$  (m) is the latitudinal distance of a model grid.
- $g$  is the gravitational force constant.
- $c_p$  is the specific heat at constant pressure with a value of  $1000 \text{ J kg}^{-1} \text{ K}^{-1}$ .
- $T_r$  is the reference temperature with a value of 737.15 K.
- $p_s$  is the surface pressure (Pa) in the control run.
- $d\sigma$  is the difference between two sigma levels.
- $\theta$  is the Latitude.
- $\sum \sum \sum$  is the horizontal and vertical summation.
- The first two terms are the kinetic energy of the perturbation while the third term is related to the potential energy of the perturbation.