

RETRIEVAL OF UPPER HAZE AEROSOL PROPERTIES FROM SPICAV-UV AND -IR DATA

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UPPER HAZE STUDY OVERVIEW

- Submicron particles ($r \leq 0.3 \mu\text{m}$)
 - Polarimetry from Pioneer Venus Orbiter (12 years of observations)
 - Kawabata et al., 1980; Esposito and Travis, 1982; Sato et al., 1996; Braak et al., 2002
 - Day-side limb scans from Pioneer Venus Orbiter
 - Lane and Opstbaum, 1983 (12 orbits analyzed)
 - Day-side limb scans from Venera 9, 10
 - Krasnopol'sky, 1983 (2 orbits)
- 1- μm particles
 - Solar occultations from SPICAV/SOIR onboard VEx
 - Montmessin et al., 2008 (joint analysis of 3 channels for data set of 3 orbits)
 - Wilquet et al., 2009 (independent analysis of 3 channels for data set of 3 orbits)
 - Luginin et al., 2016 (single SPICAV IR channel analysis of full 8 year dataset)
 - Nightside limb observations performed by VIRTIS onboard Vex
 - De Kok et al., 2011 (4 observations)

SPICAV-UV AND IR SPECTROMETERS

	UV	IR
Spectral range	118–320 nm	0.65–1.05 μm
Resolution	1.5 nm	7.8 cm ⁻¹

Single pixel FOV = 40"×40"
Off-axis parabolic mirror + concave
UV grating + CCD matrix with 384×288
exposed elements
1 measurement – 5 bins containing 1, 2, 4,
8, 16, or 32 lines in spatial direction

Combination of AOTF and 2
photodiodes
Consecutive scan of a spectrum
FOV = 0.07° at solar occultation
mode



METHODOLOGY OF SOLAR OCCULTATION

$$1) \quad T(\lambda) = \frac{I_A(\lambda)}{I_0(\lambda)}$$
$$2) \quad \tau_\lambda = -\ln(T(\lambda))$$

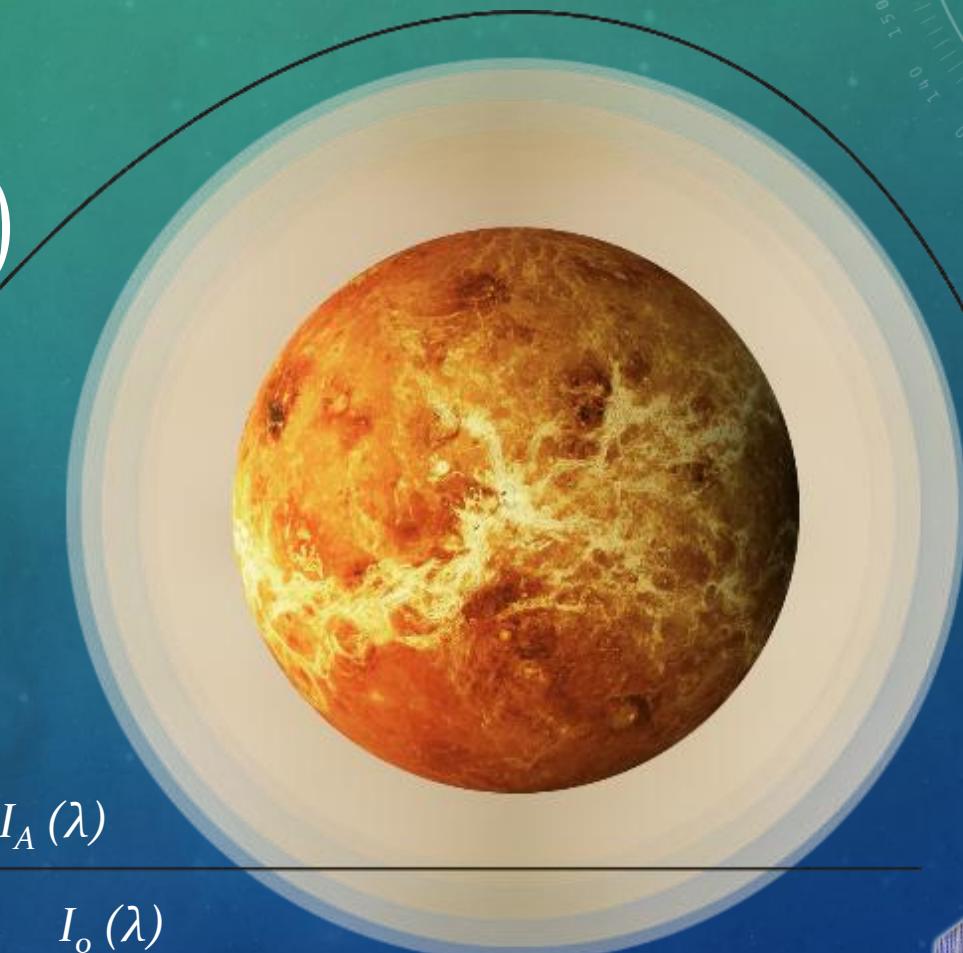
3) “Onion peeling”

$$\downarrow k_{ext}(\lambda)$$



$$I_A(\lambda)$$

$$I_o(\lambda)$$



RETRIEVAL OF AEROSOL EXTINCTION

UV

Belyaev et al. 2017. Night side distribution of SO₂ content in Venus' upper mesosphere. *Icarus*, 294, 58–71.

Fit of the transmission spectra in 180–300 nm

- CO₂, SO₂, SO absorption
- CO₂ Rayleigh scattering
- Aerosol parameterization: “α-model” τ_0 and α are free parameters

$$\tau_{aer}(\lambda) = \tau_0 \times (\lambda_0/\lambda)^\alpha$$

$$\lambda_0 = 235 \text{ nm}$$

- ‘Onion peeling’ method with Tikhonov’s regularization to retrieve aerosol extinction
- 6 UV wavelengths to maintain the “density” of IR wavelengths
- 200, 220, 236, 260, 280, 300 nm
- Interpolation to IR altitude mesh

IR

Luginin et al. 2016. Aerosol properties in the upper haze of Venus from SPICAV IR data. *Icarus*, 277, 154–170.

- Transmission at 10 wavelengths outside of gaseous absorption bands
- ‘Onion peeling’ method to retrieved extinction
- Subtraction of CO₂ Rayleigh scattering gives aerosol extinction
- 649.4, 756.6, 852.5, 982.4, 1101.1, 1159.6, 1197.3, 1273.4, 1323.0, 1553.7 nm

LATITUDE COVERAGE OF OBSERVATIONS

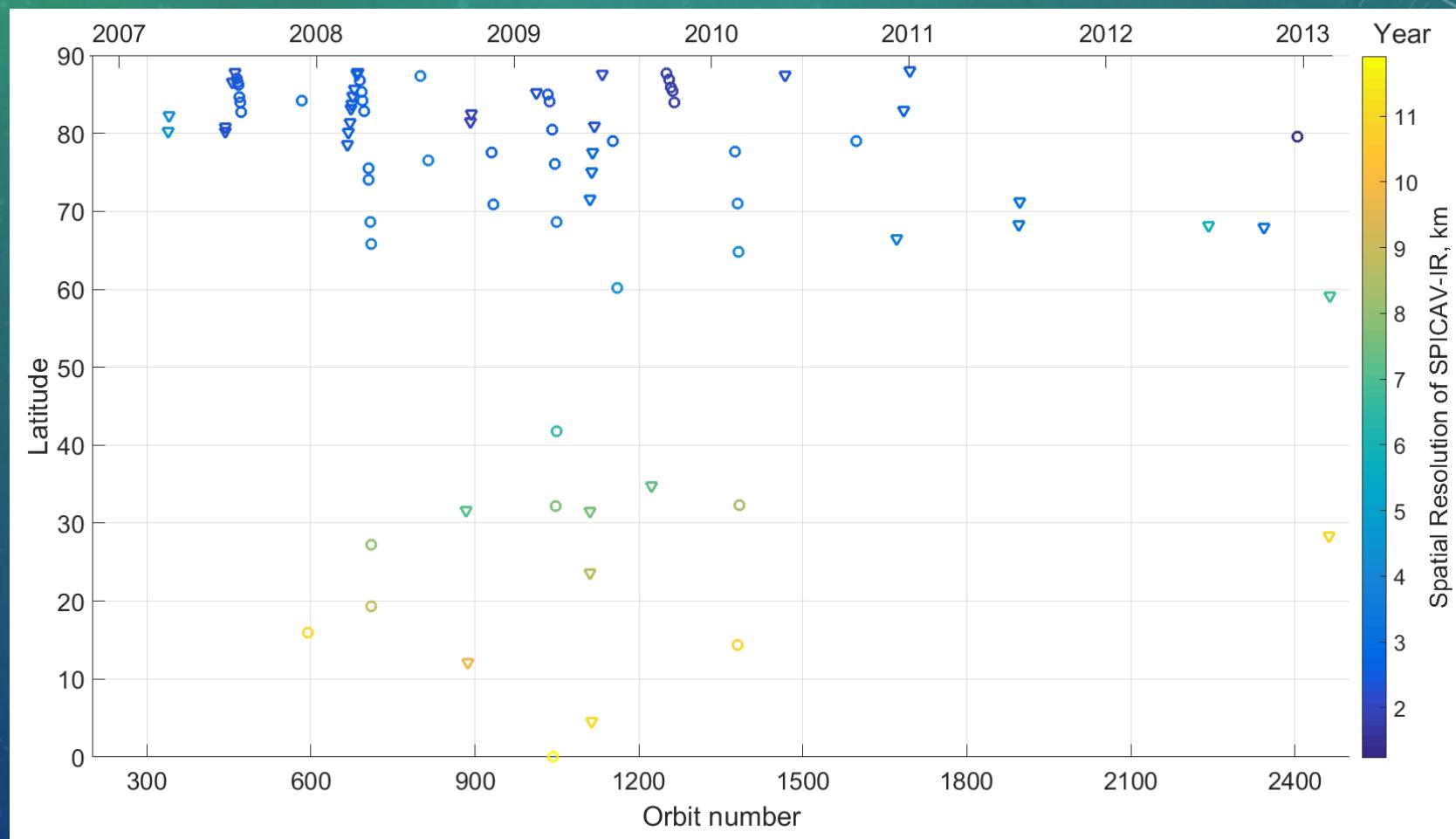
UV

402 solar occultation sessions

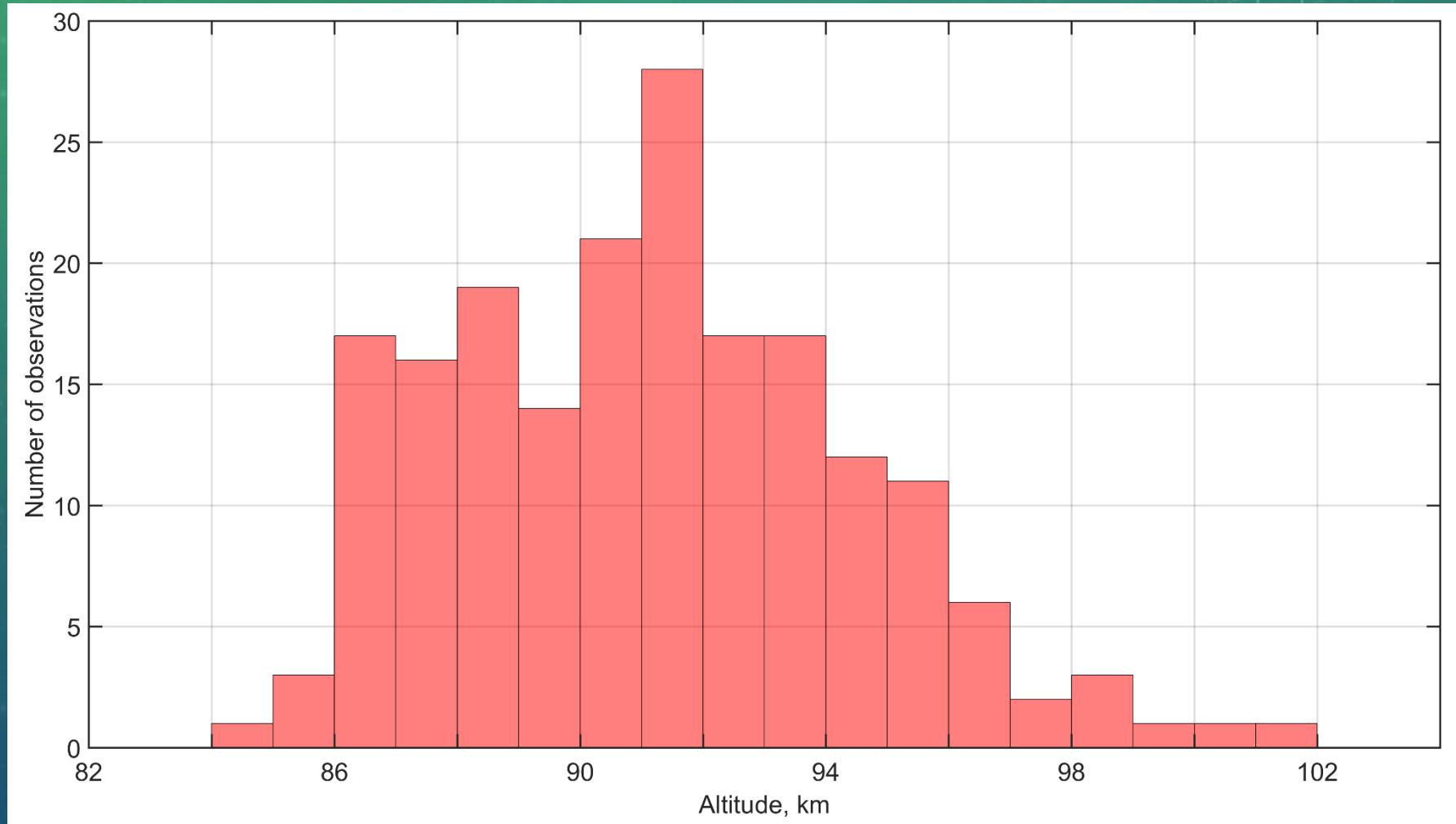
IR

222 solar occultation sessions

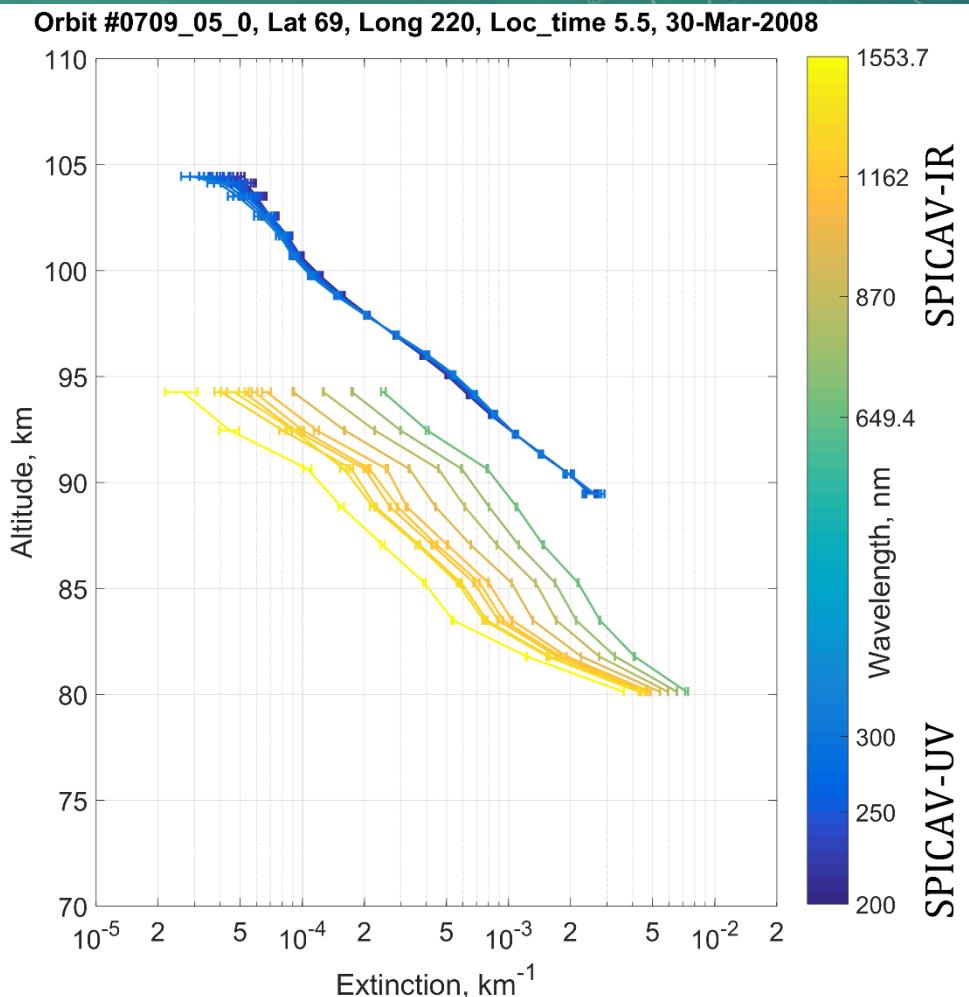
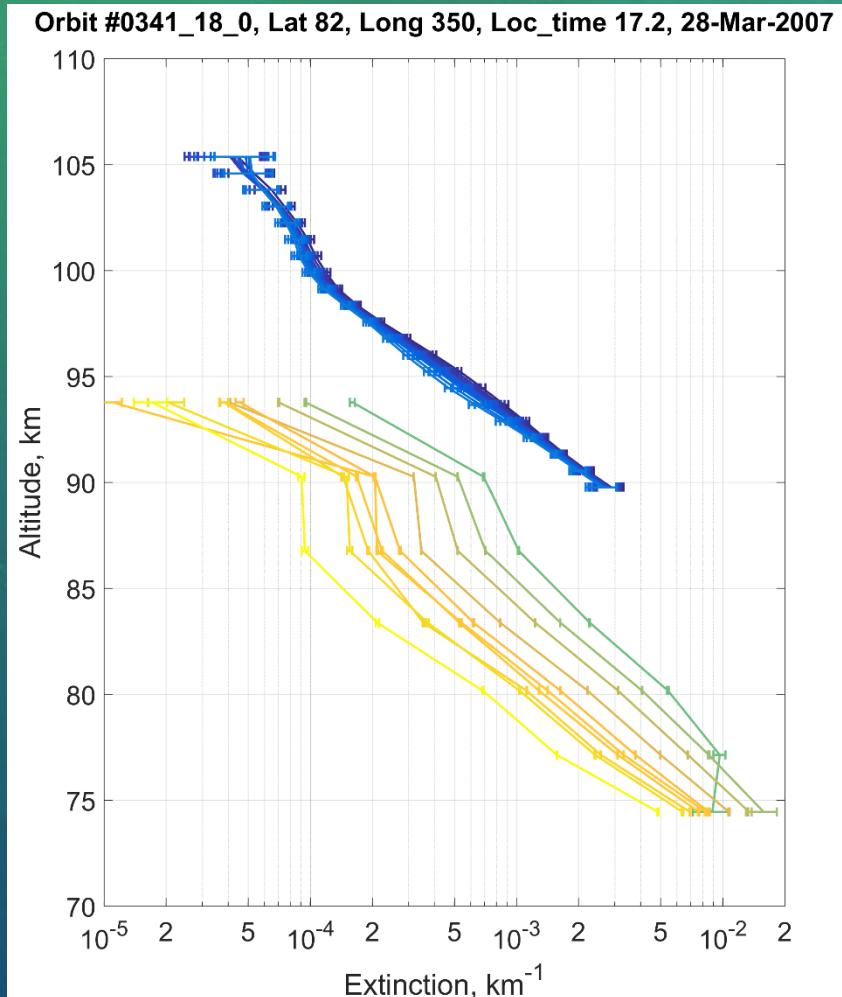
84 solar occultation sessions
186 altitudes



ALTITUDE COVERAGE

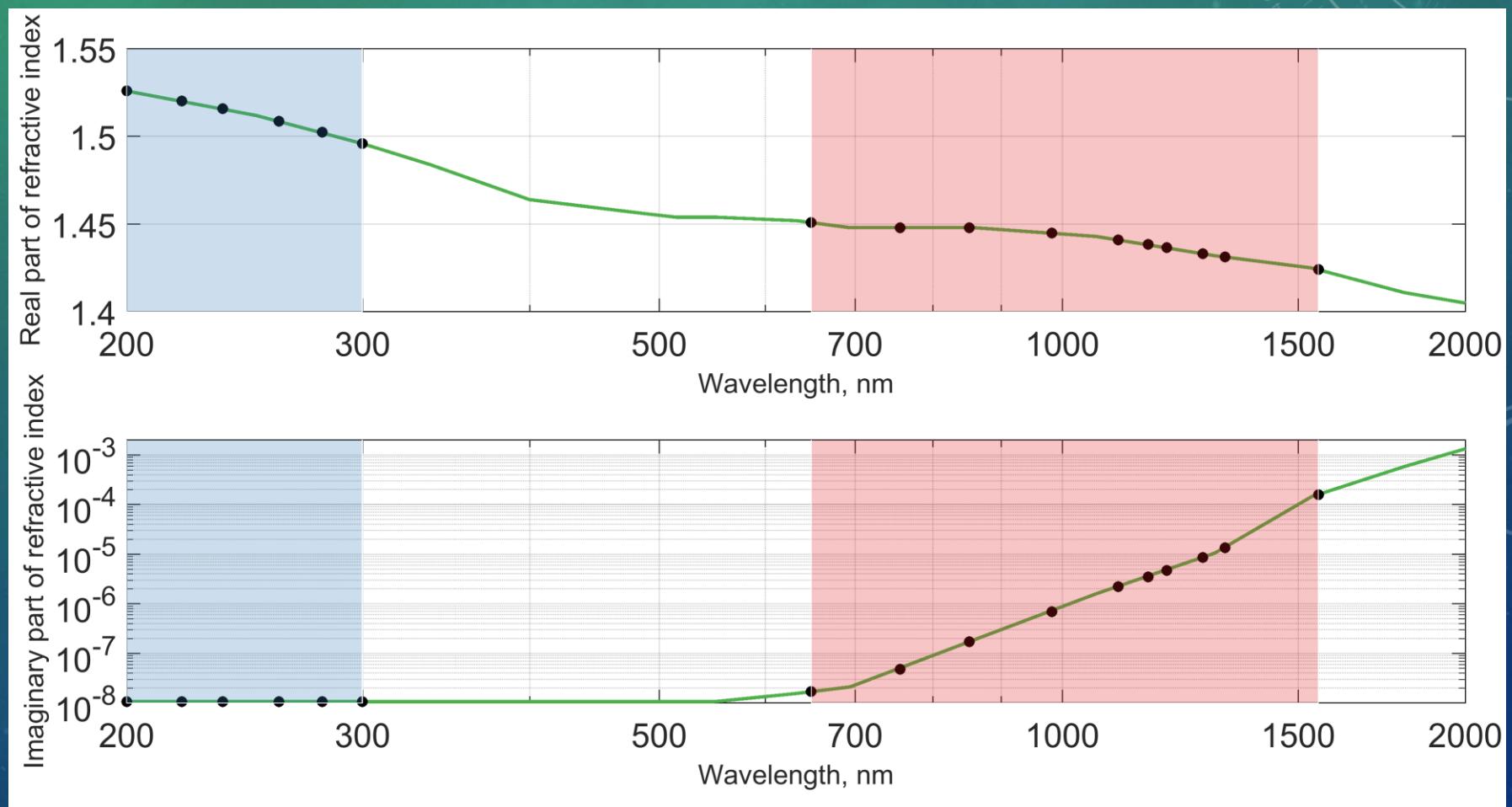


JOINT EXTINCTION PROFILES



H_2SO_4 refractive index

Real and Imaginary indices of refraction for 75 % solution of sulfuric acid at 215 K were taken from Hummel et al. (1988), A new background stratospheric aerosol model for use in atmospheric radiation models, AFGL-TR-88-0166, Air Force Geophys. Lab., Hanscom Air Force Base, Mass



PARTICLE SIZE AND NUMBER DENSITY DETERMINATION

Levenberg-Marquardt fitting procedure to find distribution parameters

$$\chi^2 = \frac{1}{N-k} \sum_1^N \frac{(k_{\text{exp}} - k_{\text{mod}})^2}{\delta_{\text{exp}}^2}$$

Unimodal lognormal distribution

r_g is fitted
 ν_{eff} is set to 0.15

$$n(r) = C \times r^{-1} \exp\left(-\frac{(\ln r - \ln r_g)^2}{2 \ln^2 \sigma_g}\right)$$
$$r_{\text{eff}} = r_g \exp\left(\frac{5}{2} \ln^2 \sigma_g\right)$$
$$\nu_{\text{eff}} = \exp(\ln^2 \sigma_g) - 1$$

Bimodal lognormal distribution

$$n(r) = \frac{C}{r} \left\{ \exp\left[-\frac{(\ln r - \ln r_{g1})^2}{2 \ln^2 \sigma_{g1}}\right] + \gamma \exp\left[-\frac{(\ln r - \ln r_{g2})^2}{2 \ln^2 \sigma_{g2}}\right] \right\}$$

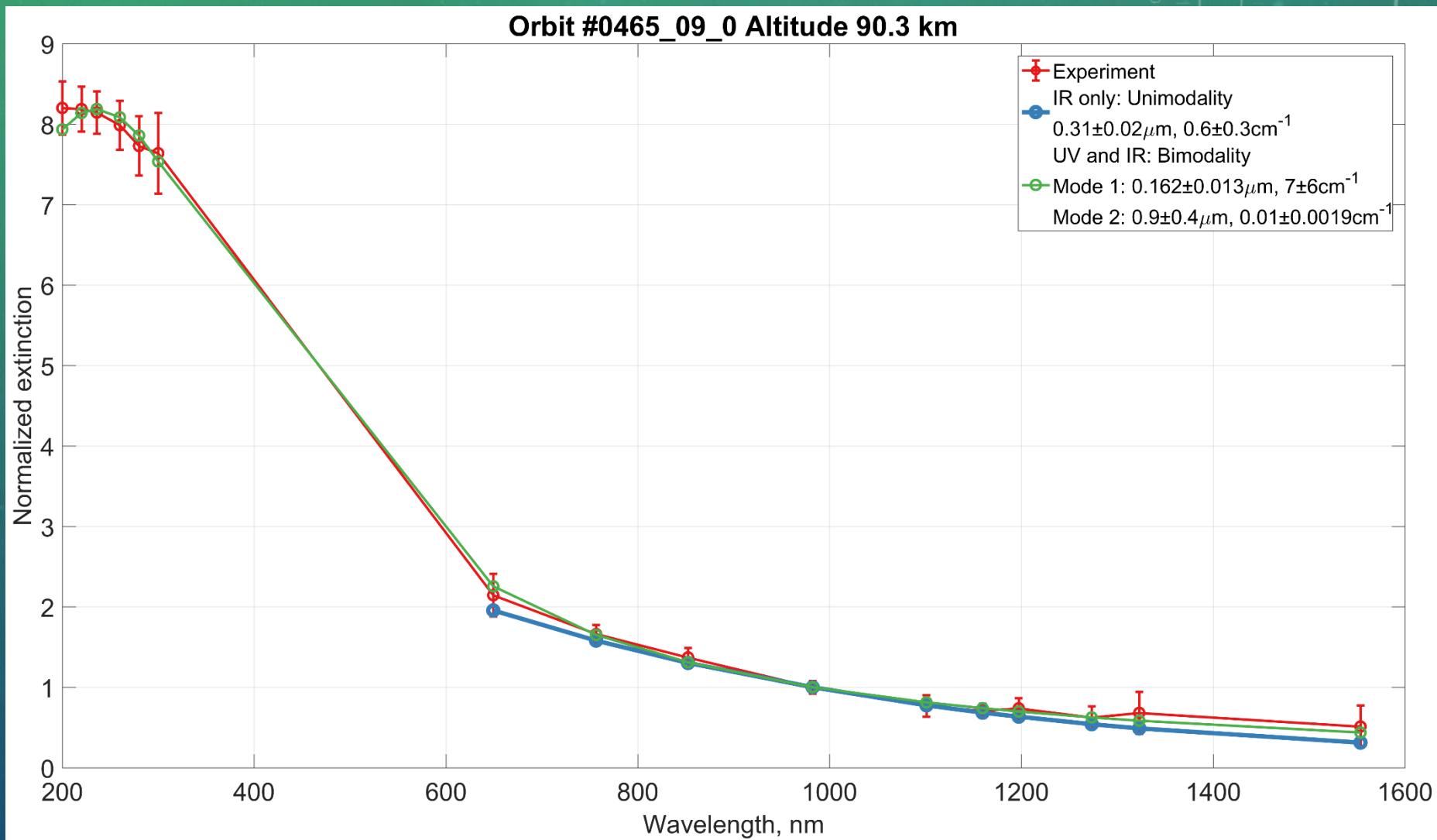
r_{g1} , r_{g2} and γ are fitted

$\nu_{\text{eff}1} = 0.15$, $\nu_{\text{eff}2} = 0.1$

$$N(z) = \frac{k_{\text{ext}}(z)}{\int_0^\infty \sigma_{\text{ext}}(\lambda, z) n(r, z) dr}$$

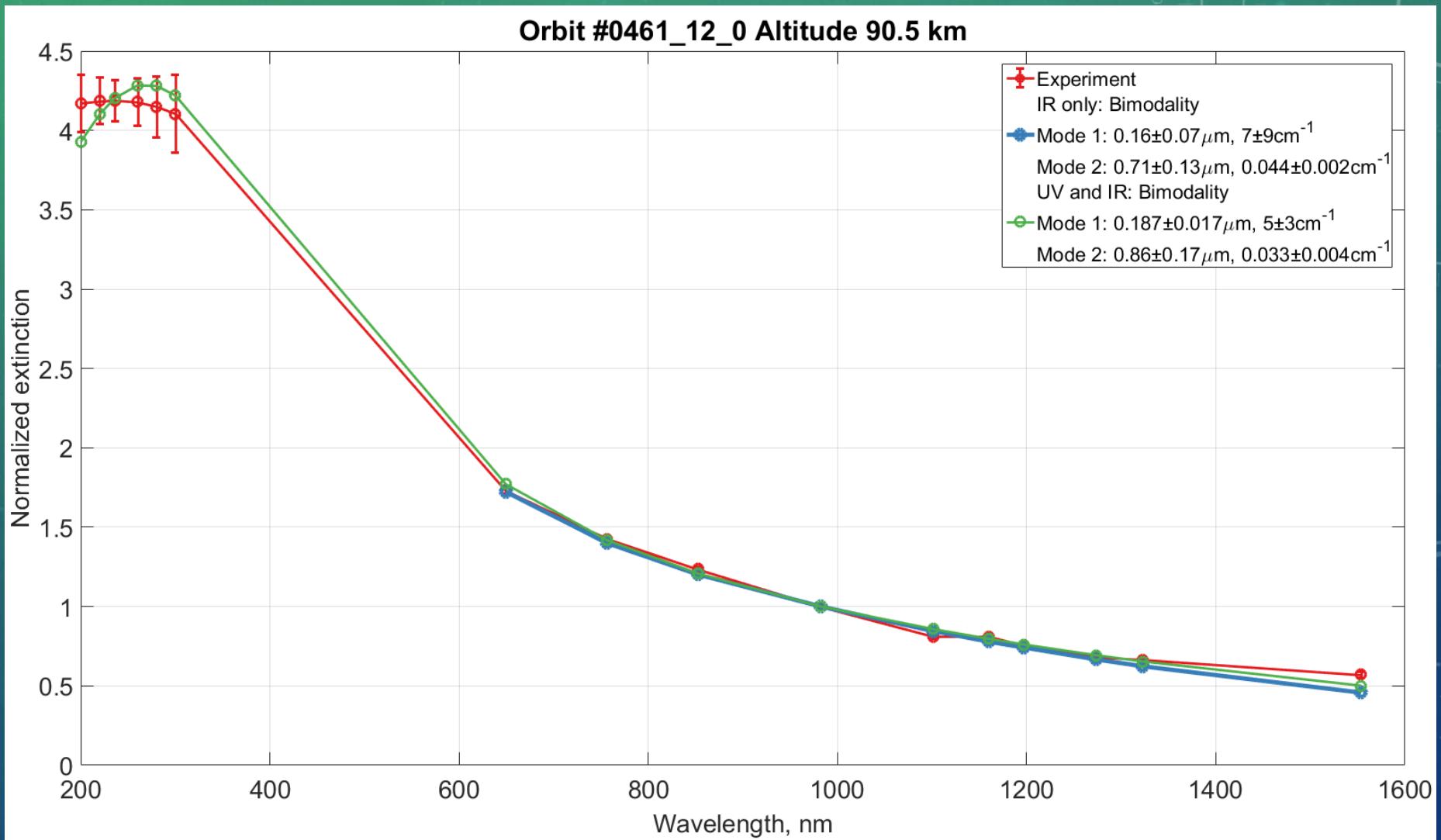
RESULTS:

unimodal distribution → bimodal distribution



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bimodal distribution → bimodal distribution

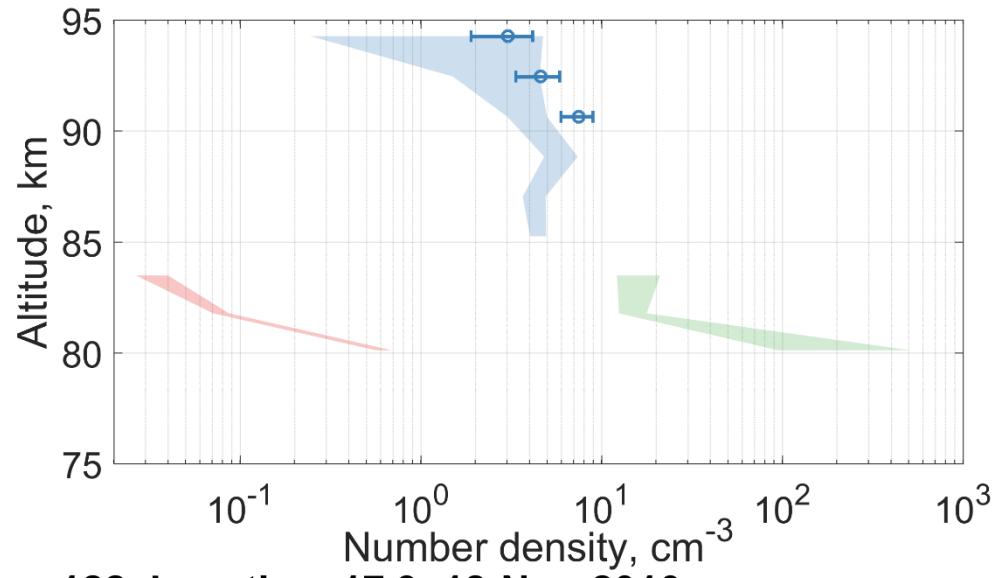
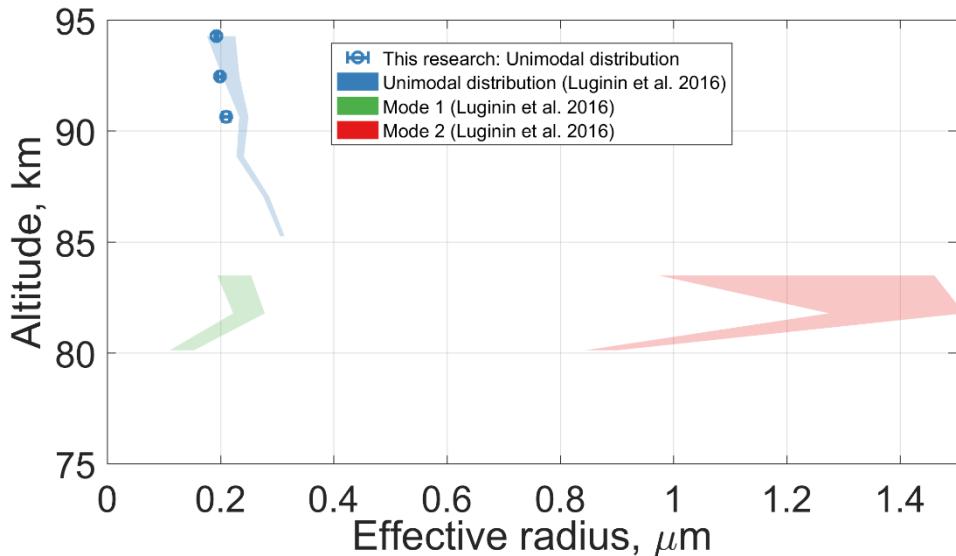




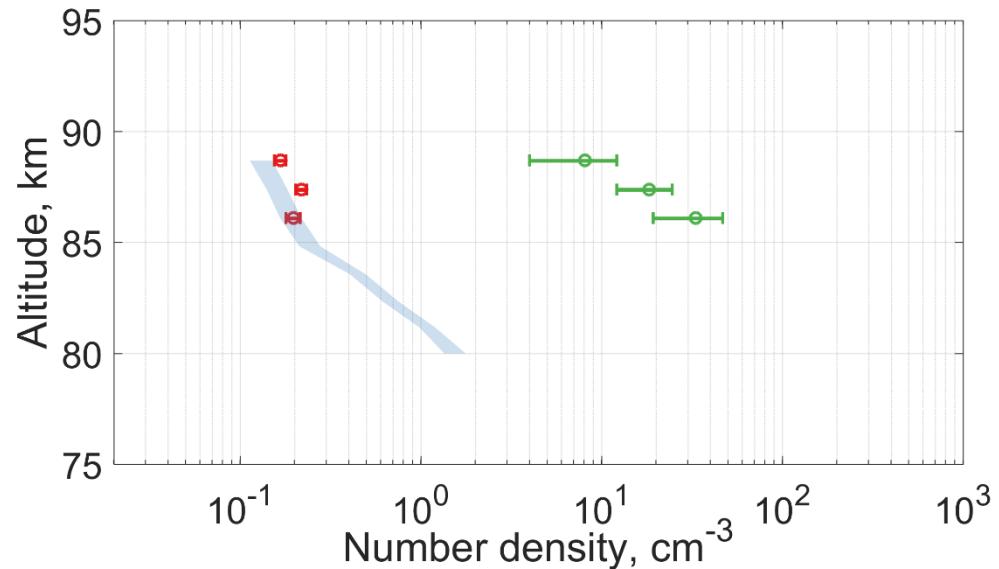
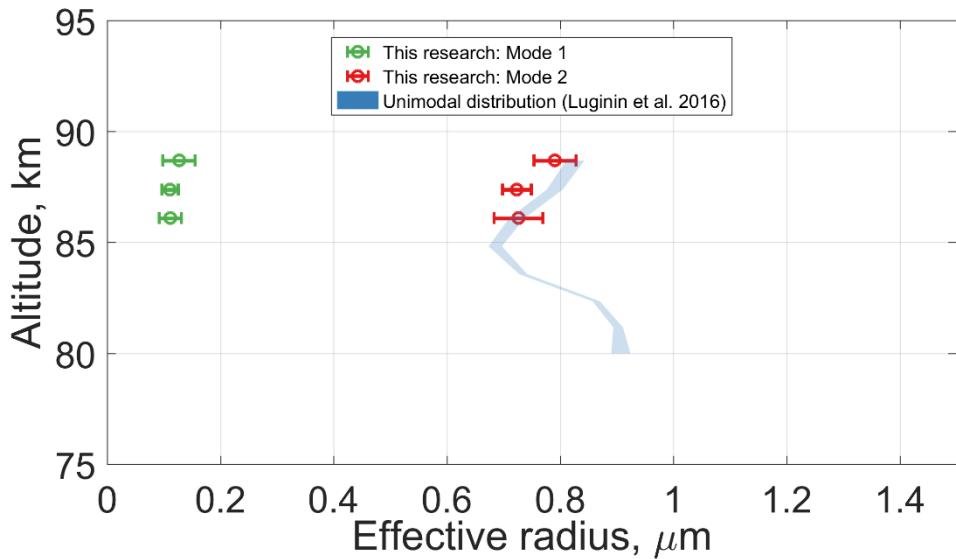
RESULTS:

comparison of IR only and joint UV and IR data

Orbit #0709_05_0, Lat 69, Long 220, Loc_time 5.5, 30-Mar-2008

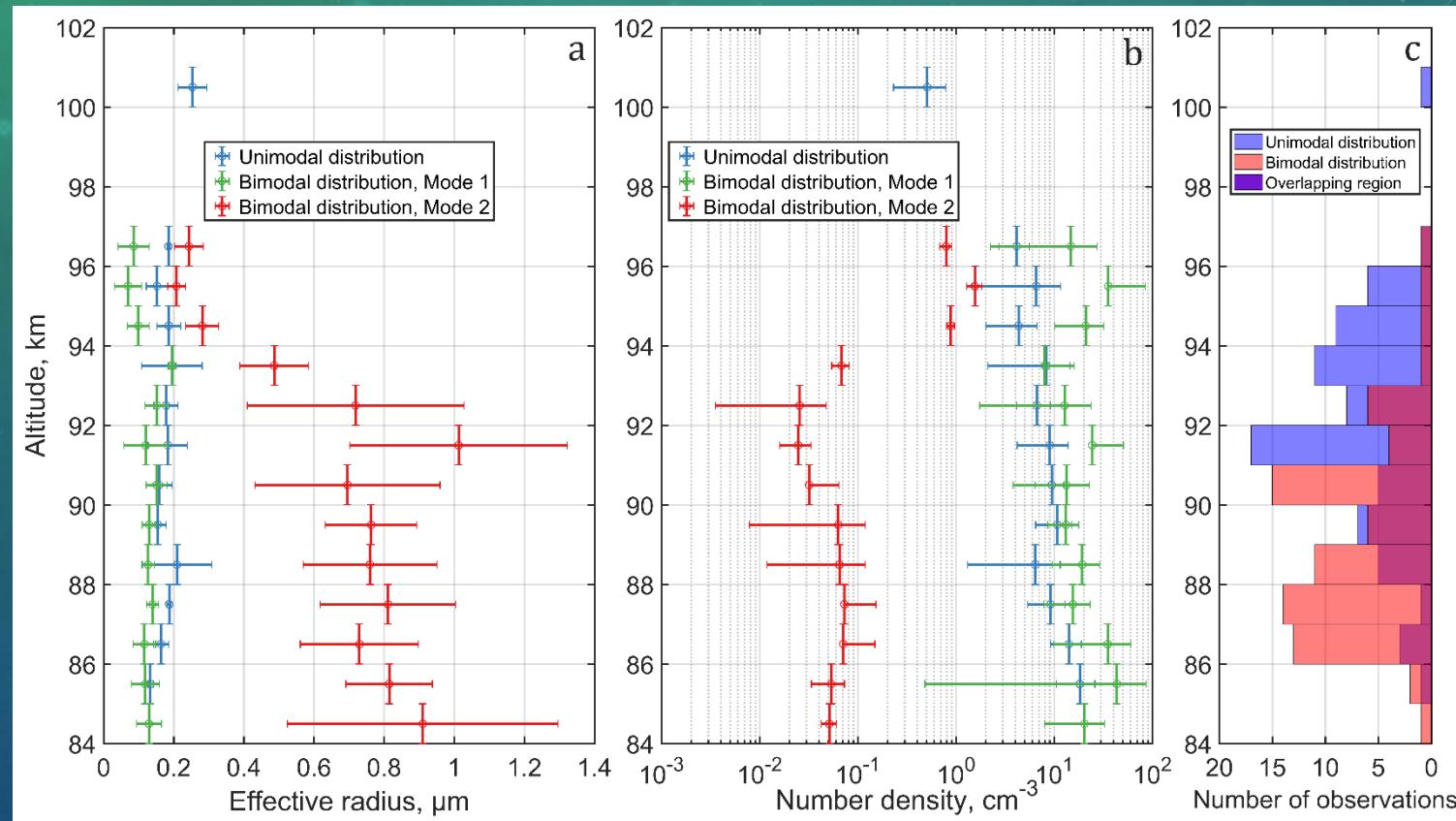


Orbit #1672_06_2, Lat 67, Long 122, Loc_time 17.9, 18-Nov-2010

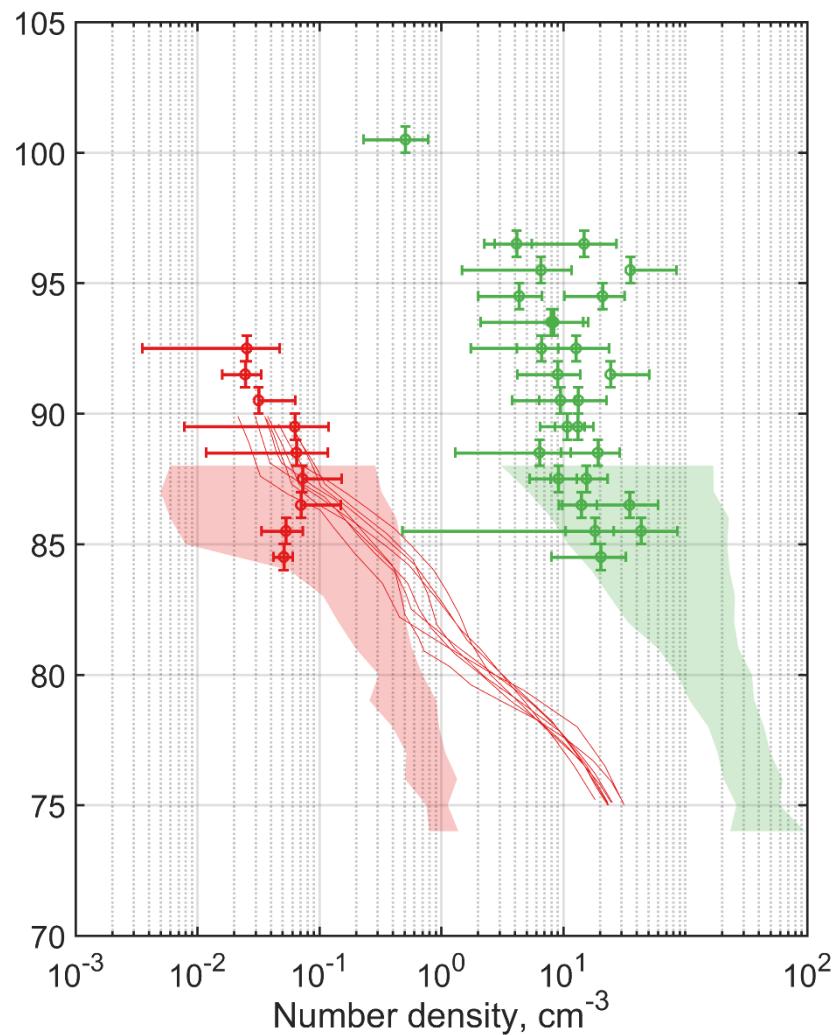
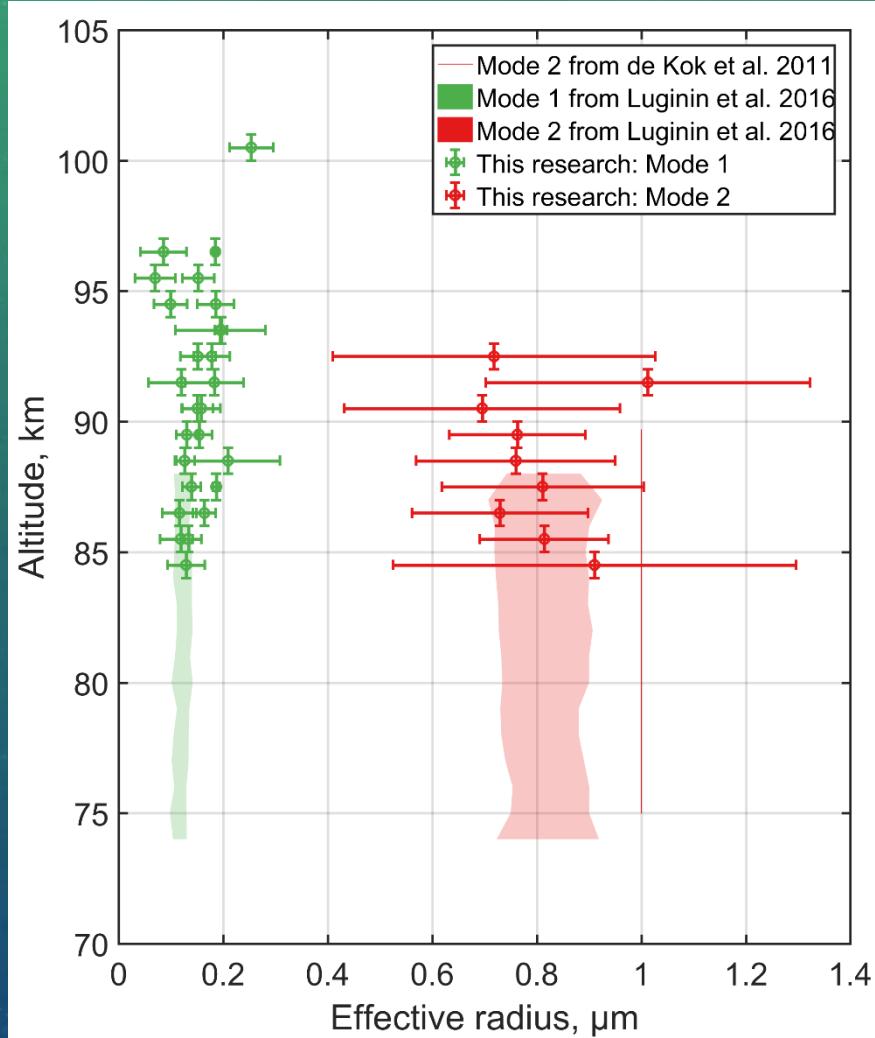


MEAN PROFILES

- Unimodal case: $0.18 \pm 0.05 \mu\text{m}$ and $8 \pm 5 \text{ cm}^{-3}$
- Bimodal case averaged in 86–92 km altitude range
- Mode 1: $0.13 \pm 0.03 \mu\text{m}$ and $20 \pm 14 \text{ cm}^{-3}$
- Mode 2: $0.8 \pm 0.2 \mu\text{m}$ and $0.05 \pm 0.04 \text{ cm}^{-3}$



COMPARING WITH PREVIOUS RESULTS



MAIN RESULTS

- We prove that SPICAV-IR is capable to detect bimodal size distribution of aerosols
- We confirm results obtained from single SPICAV-IR channel. Luginin et al. 2016. Aerosol properties in the upper haze of Venus from SPICAV IR data. Icarus, 277, 154–170.
- Upper haze extends up to 95+ km at high latitudes
- Mode 2 particles are observed up to 91 km at high latitudes