

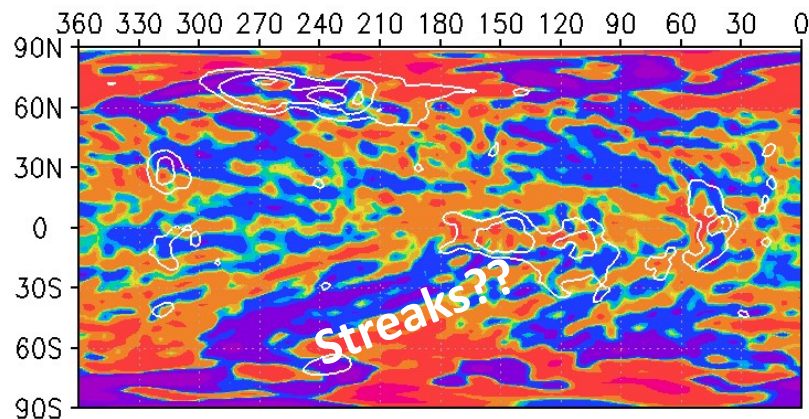
# Atmospheric simulations using Venus AORI general circulation models

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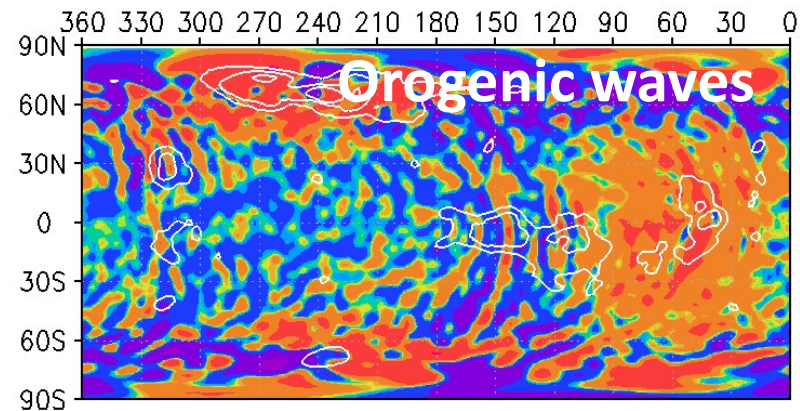
#AORI: *Atmosphere and ocean Research Institute, Univ. Tokyo.*

*This work is supported by cooperative research project for climate system research of AORI.*

## T63L52 with radiative transfer and topography



Snapshot of vertical flow (mm/s)  
at 50 km ~ cloud base



Snapshot of vertical flow (mm/s)  
at 60 km ~ upper clouds

# Overview of Venus AORI GCM

## 1. Venus GCM with simplified radiation

### Venus GCM from initially motionless state

Yamamoto & Takahashi (2003JAS) : SR by zonal mean heating (T10)

Yamamoto & Takahashi (2003GRL) : SR by zonal mean heating (T21), fast Kelvin wave

Yamamoto & Takahashi (2004GRL, 2006JAS, 2009JGR) : SR by 3D heating, thermal tide, &

Yamamoto & Takahashi (2007EPS) : SR by 3D heating with topography Kelvin-Rossby waves

Yamamoto & Takahashi (2007GRL, 2008AA) : Effects of seasonal variation & planetary rot.

### Venus middle atmospheric GCM

Yamamoto & Takahashi (2012Icarus) : Dynamics of Kelvin wave

Yamamoto & Takahashi (2015PSS) : Polar vortex in the presence of thermal tide

Yamamoto & Takahashi (2018Icarus) : Interannual zonal-flow variation

### Sensitive issues for dynamical core in extremely long-term time integrations

Yamamoto & Takahashi (2016, 2018JGR) : Effects of planetary rot. & horizontal res. (T21to106)

Our recent work : Sensitivity to horizontal subgrid-scale (SGS) diffusion (T42)

## 2. Venus GCM with radiative transfer

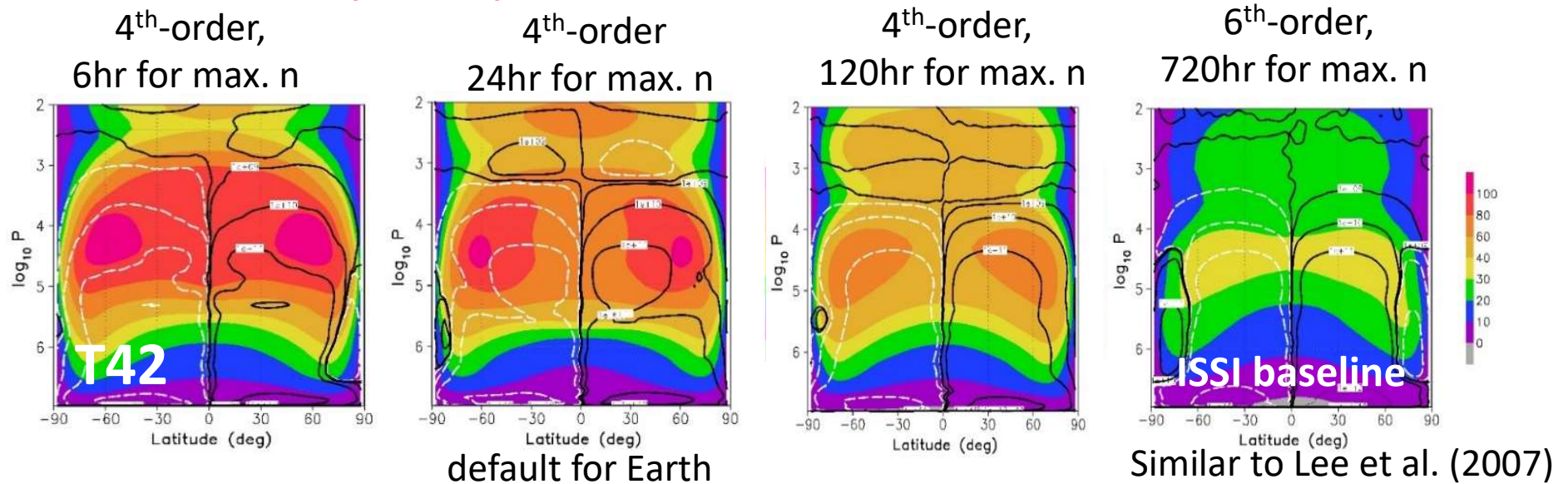
Ikeda (2011PhD) : developments of radiation code and gravity wave parameterization

Kuroda, U. Tohoku group et al. : application to cloud model

Yamamoto et al (2019Icarus) : T21 with topography and radiative code

Our recent work : T63 with topography and radiative code

# Sensitivity of dynamical core to horiz. SGS diffusion



➔ SR depends on the horiz. SGS diffusion in extremely long-term experiments from initially motionless for ISSI baseline run

## VGCM with topography and radiative code

< Purpose >

To accurately simulate thermal tides, stationary waves and general circulation by a realistic Venus GCM, along with fine and large cloud features.

➔ The results shows following slides of our presentation

# Model (Ikeda 2011 PhD)

1. Dynamical core : CCSR/NIES/FRCGC AGCM ver.5.6  
T21L52, T63L52 (Numaguti et al. 1997)
2. Radiative transfer : Discrete-Ordinate/Adding, 2-stream
  - IR and Solar: 28 bands (Nakajima et al. 2000)
  - Gas absorption
    - Correlated k-distribution method
    - CO<sub>2</sub>: CDSD-1000, sub-Lorentz lineshape(Fukabori et al., 1986)
    - H<sub>2</sub>O, SO<sub>2</sub>, OCS, CO: HITRAN2004, Voigt lineshape
    - Collision-induced absorption by CO<sub>2</sub> (Moskalencko et al., 1979)
  - Cloud
    - 75% H<sub>2</sub>SO<sub>4</sub> (refractive index: Palmer and Williams, 1975)
    - log-normal 3 mode (Pollack et al., 1980)
    - Unknown UV absorber (Crisp, 1986)
3. Gravity-wave momentum parameterization is not used.
4. Magellan topography height data (Ford & Pettengill 1992)

*#The details of model setup in Yamamoto et al. (2019 Icarus, 321, 232-250).*



# Model setup

## Temperature

Initial: VIRA (Seiff et al. 1985)

## Zonal wind

### *Nudging run until equilibrium*

0—10Vdays: U is nudged to  $U_{ref}$  with a time constant of 3 E<sub>days</sub>.

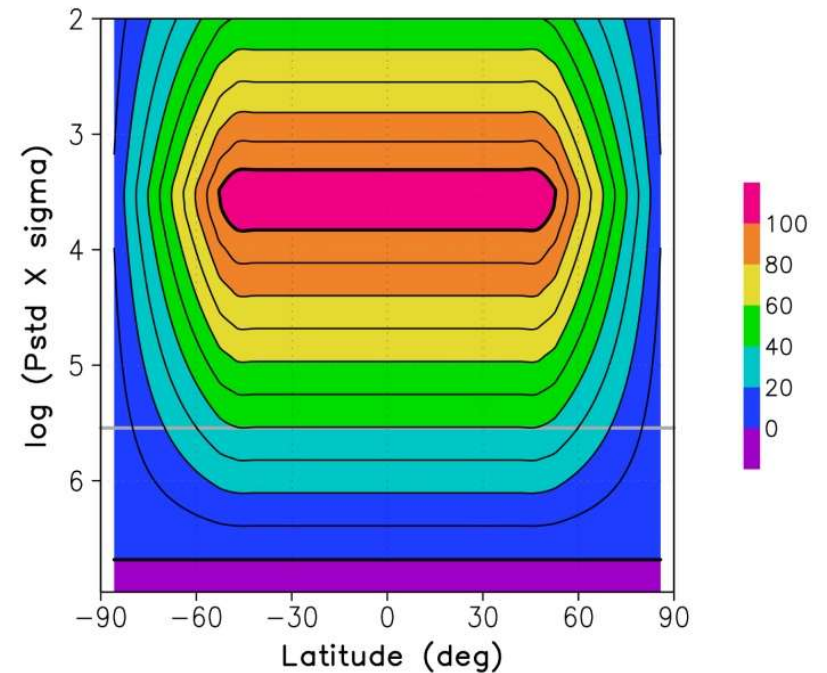
10—90Vdays: Zonal-mean component of U below 40 km is nudged to  $U_{ref}$  with a time const. of 3 E<sub>days</sub>.



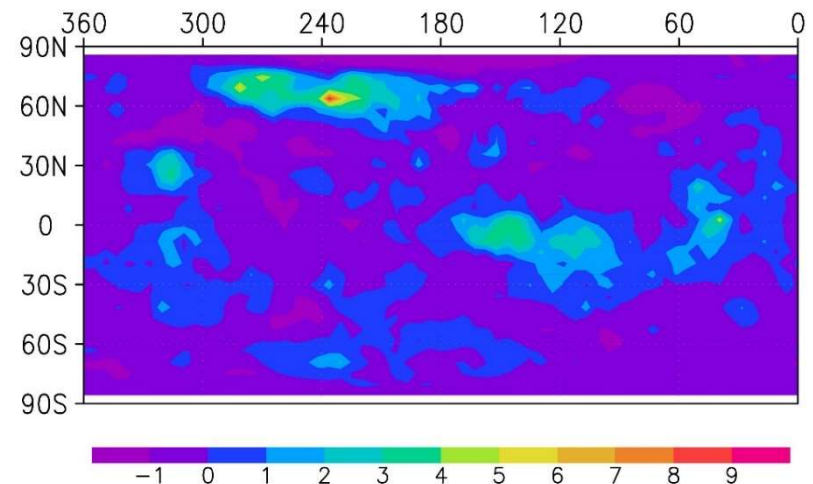
### *Nudging-free run after equilibrium*

90—93Vdays  $\Rightarrow$  Data analysis

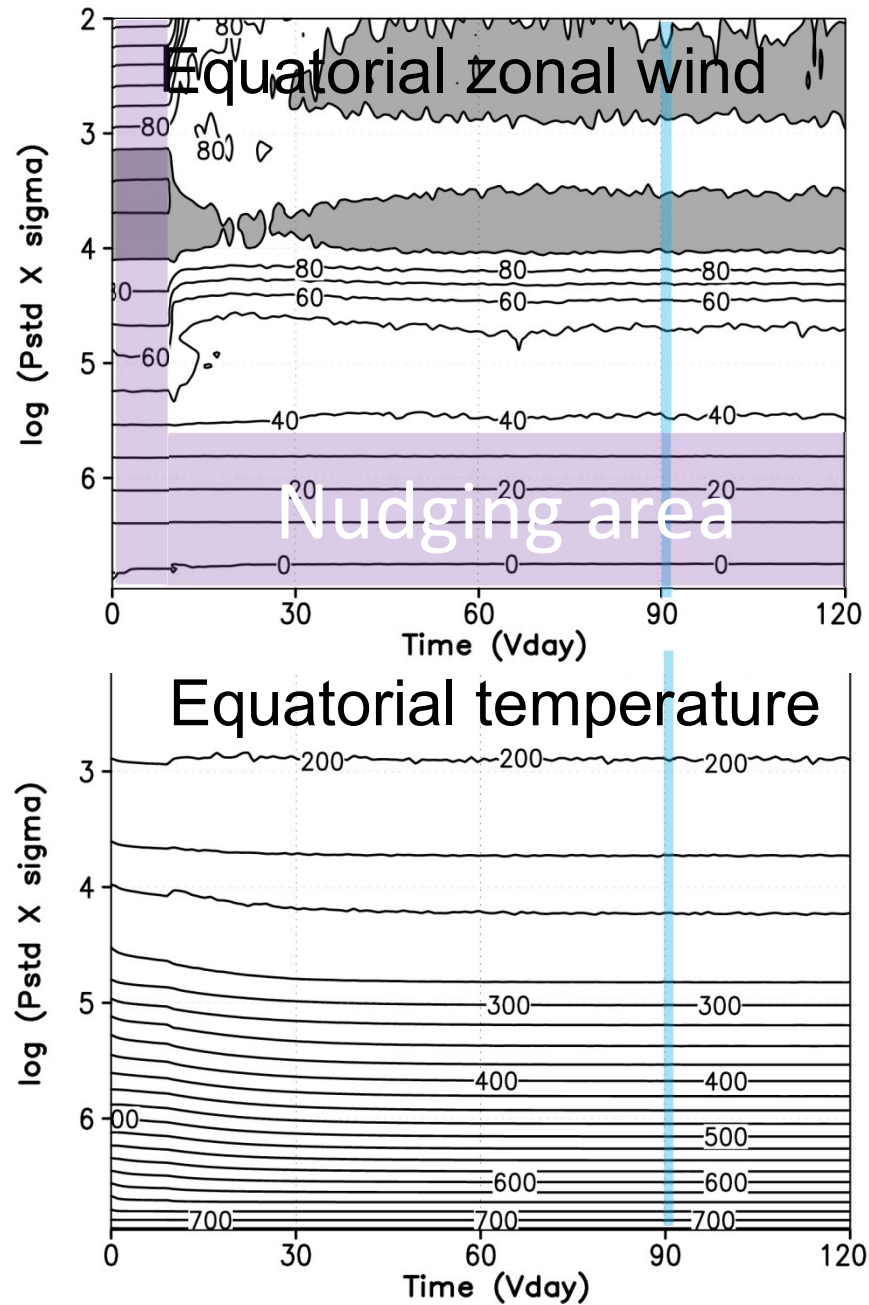
## Reference zonal flow $U_{ref}$



## Topography



## T21 Nudging run

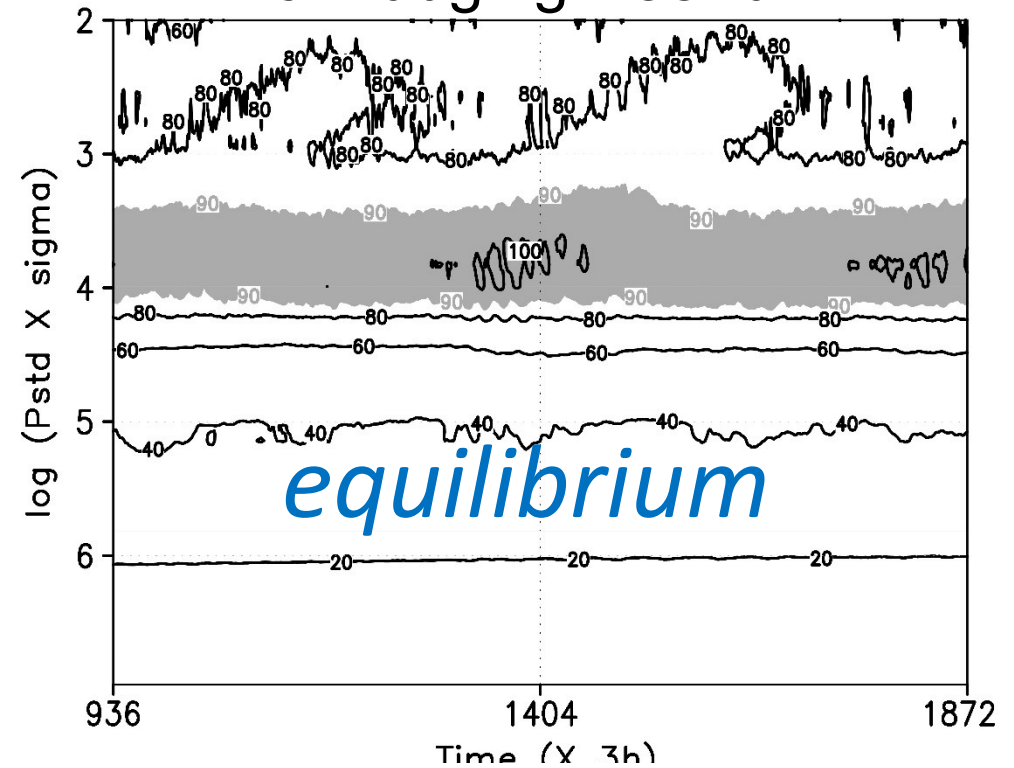


## T63 Nudging-free run

We analyzed the nudging-free model output

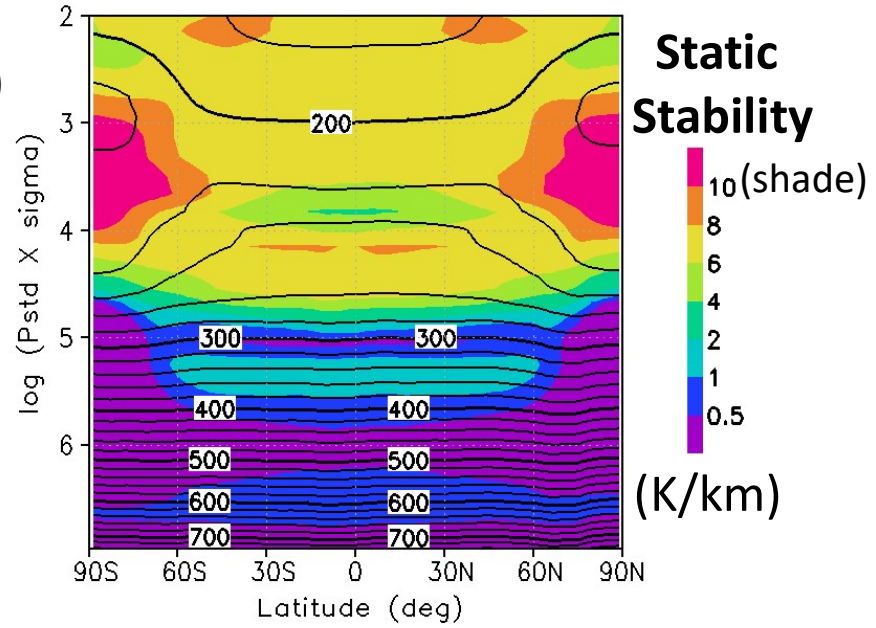
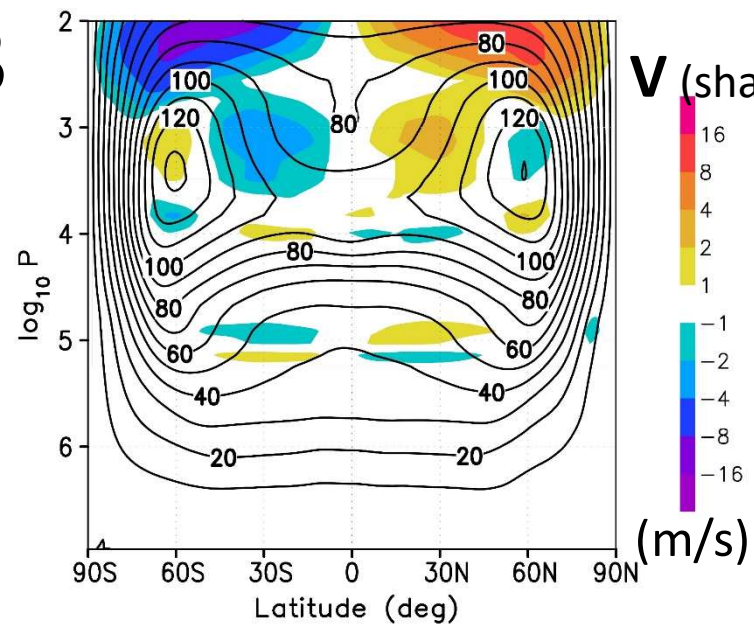


Zonal-mean equatorial flow  
of nudging-free run

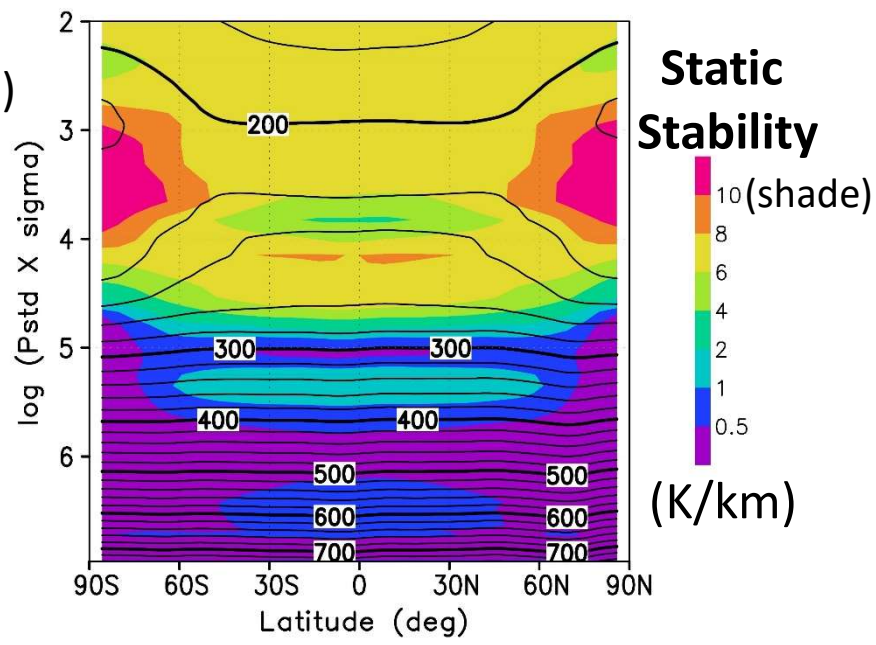
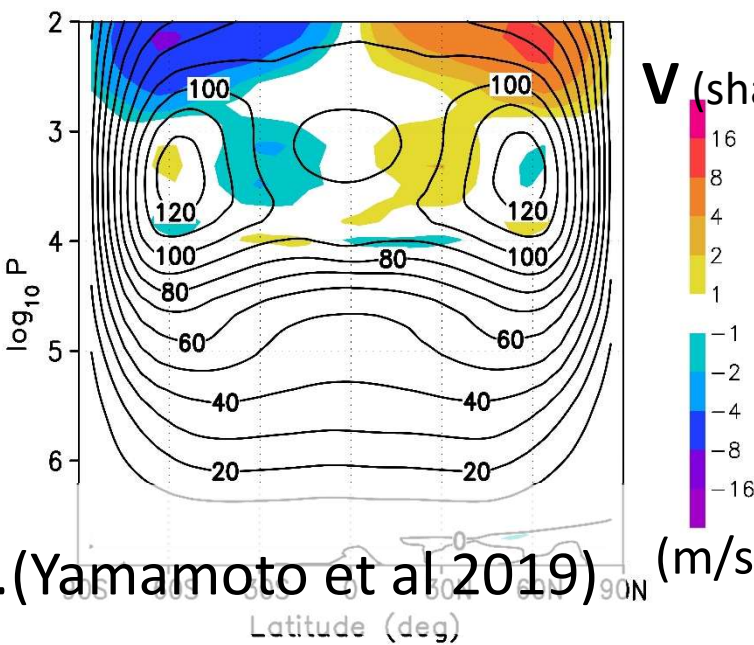


# Zonal wind (contour) Temperature (contour)

T63



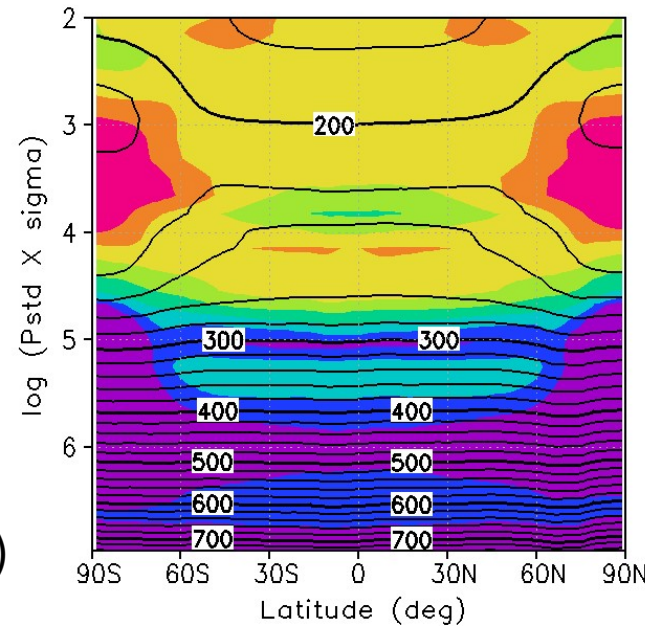
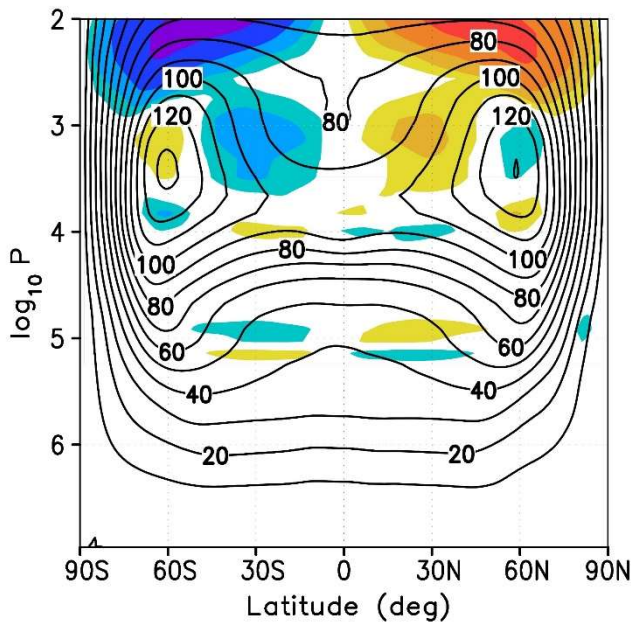
T21 (Yamamoto et al 2019)





# wind, temperature & radiative heating (bottom)

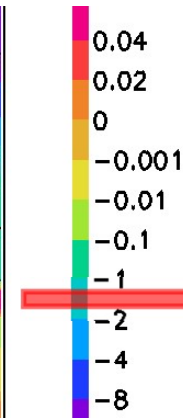
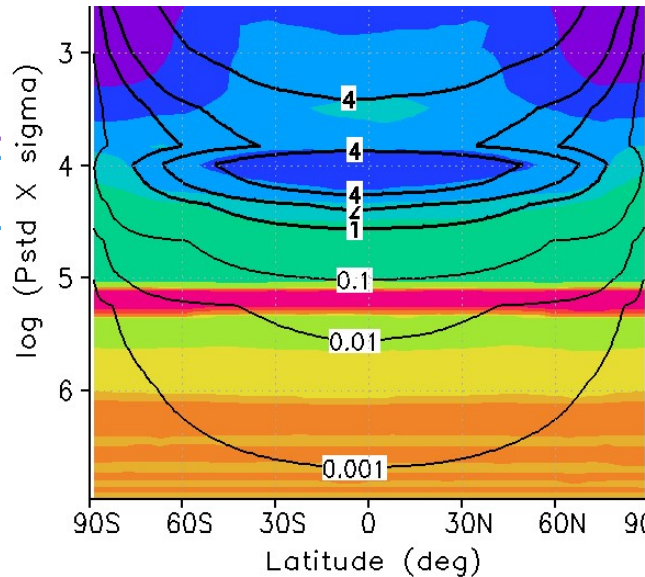
T63



Equatorial heating  
Polar cooling  
E-P Radiative *imbalance*

Local radiative *balance*

## Short-wave heating (contour, K/d)



IR heating  
at the cloud  
bottom

Long-wave  
heating (shade, K/d)



# wind, temperature & radiative heating (bottom)

T63

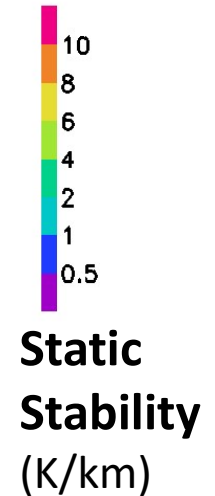
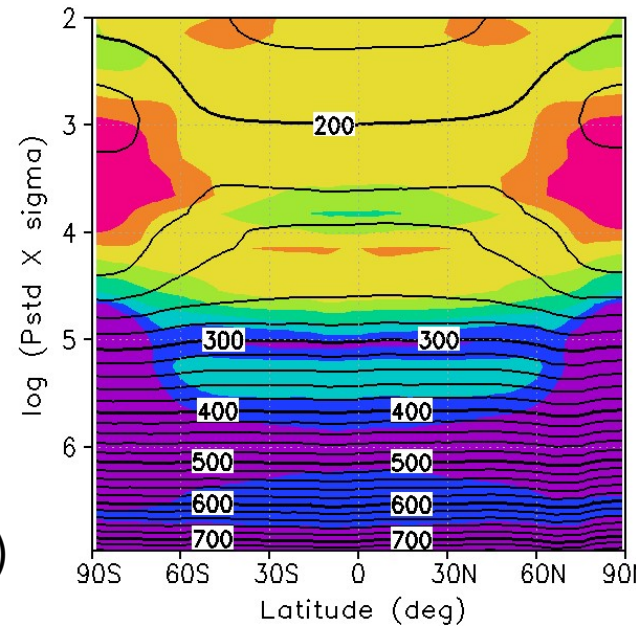
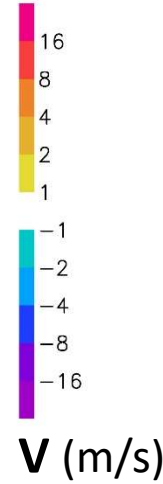
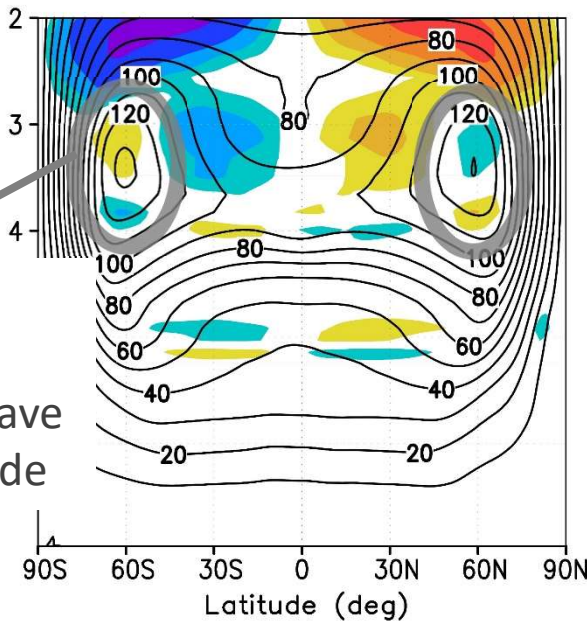
Ferrel cell induced by baroclinic wave & thermal tide

Strong Hadley cell

E-P Radiative imbalance

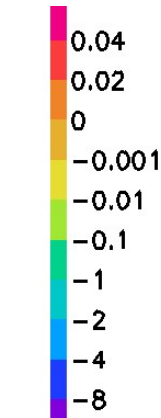
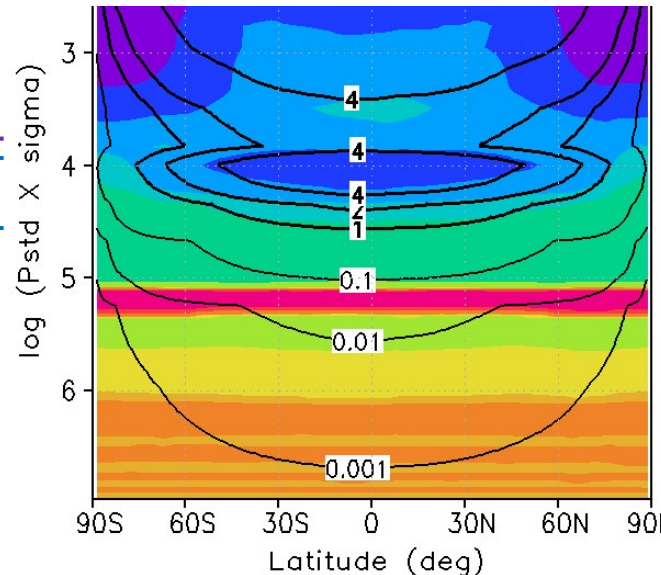
Local radiative balance

Weak meridional flow



Equatorial heating  
Polar cooling

Short-wave heating (contour, K/d)



Long-wave heating (shade, K/d)

# wind, temperature & radiative heating (bottom)

T63

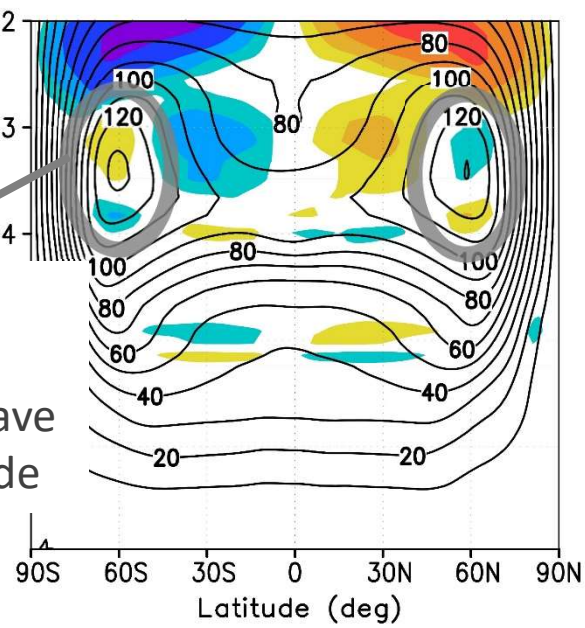
Ferrel cell induced by baroclinic wave & thermal tide

Strong Hadley cell

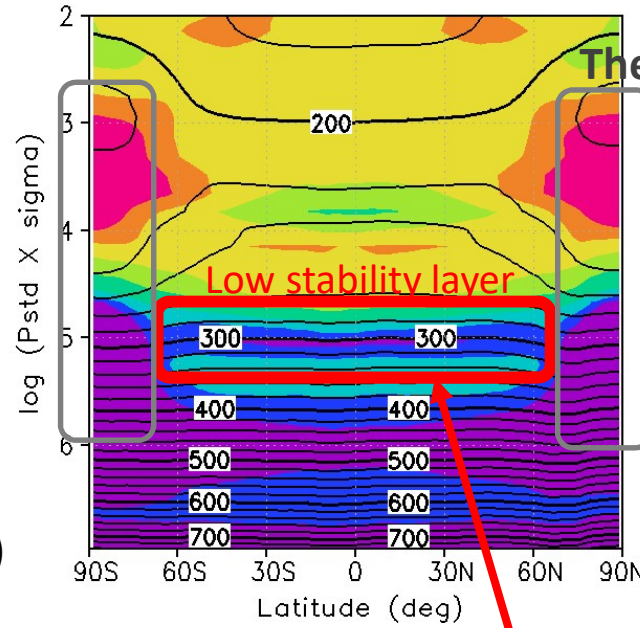
E-P Radiative imbalance

Local radiative balance

Weak Return flow



V (m/s)

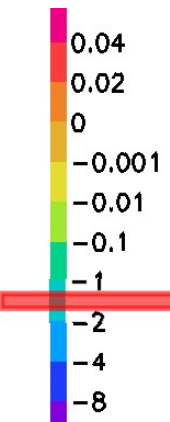
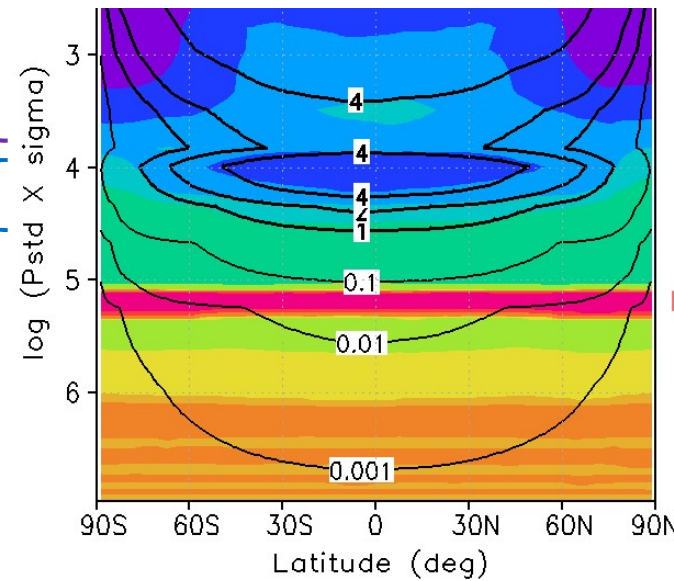


Thermal wind relation

Static Stability (K/km)

Equatorial heating  
Polar cooling

Short-wave heating (contour, K/d)

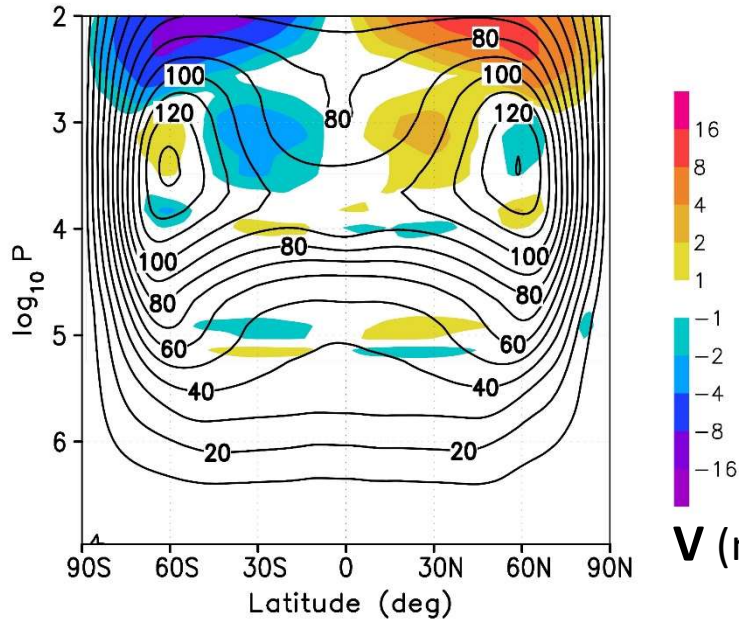


Long-wave heating (shade, K/d)

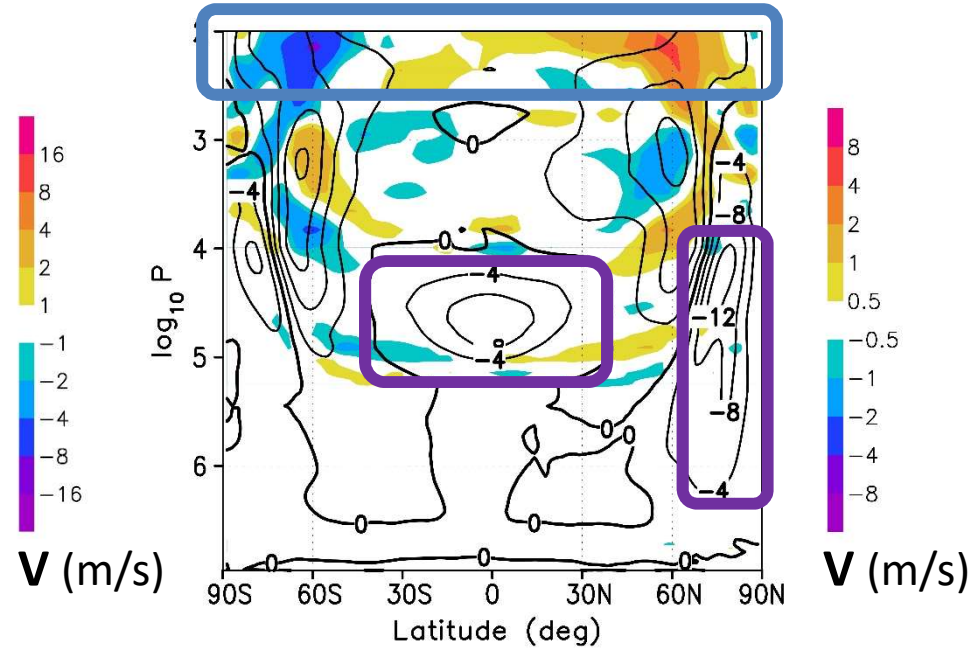
IR heating at the cloud bottom

# U & V with topo

T63



# topo-flat



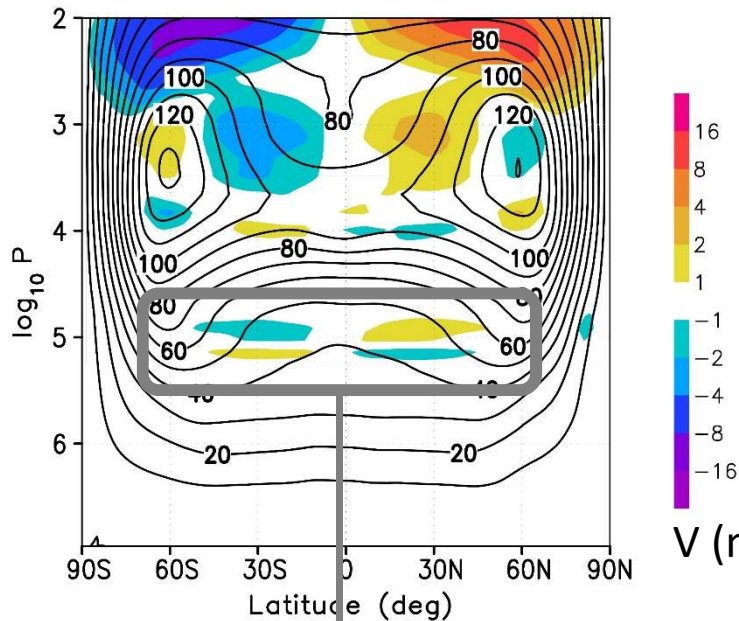
Topographically induced  
Zonal wind decrease

Poleward wind enhancement



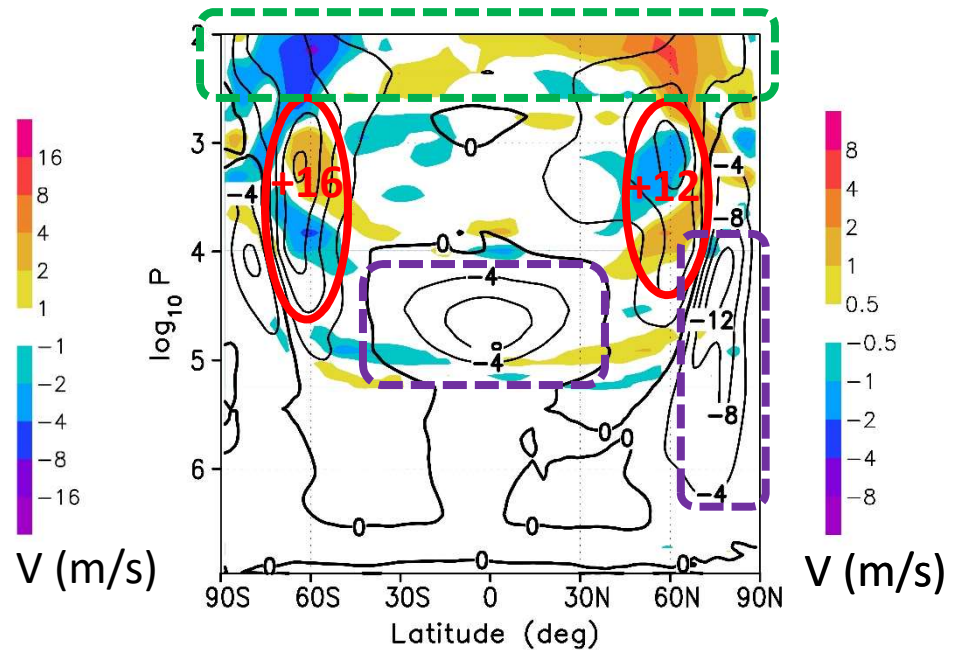
# U & V with topo

T63



a thin Hadley circulation

# topo-flat



Topographically induced  
Zonal wind decrease

Poleward wind enhancement



Enhancements of  
high-latitude jet  
and Ferrel circulation

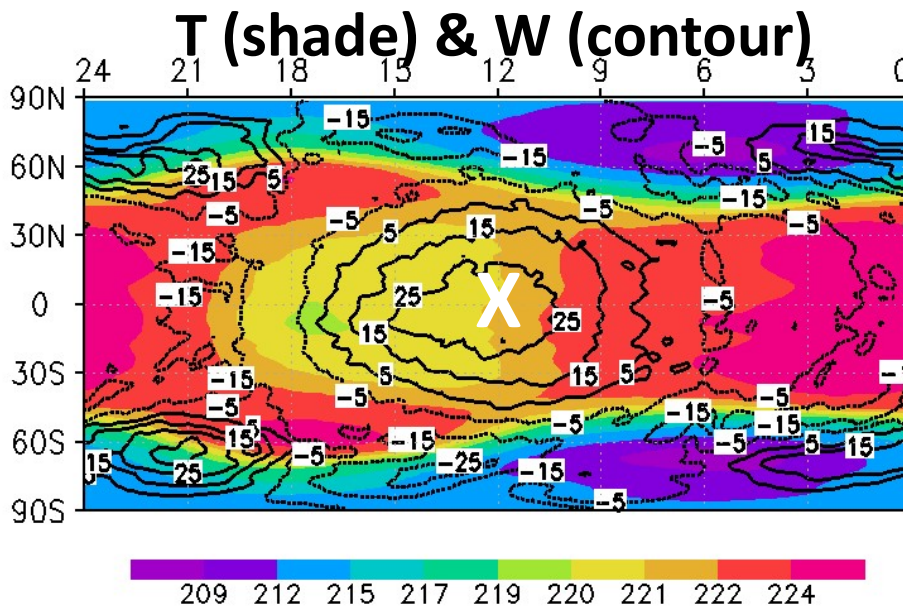
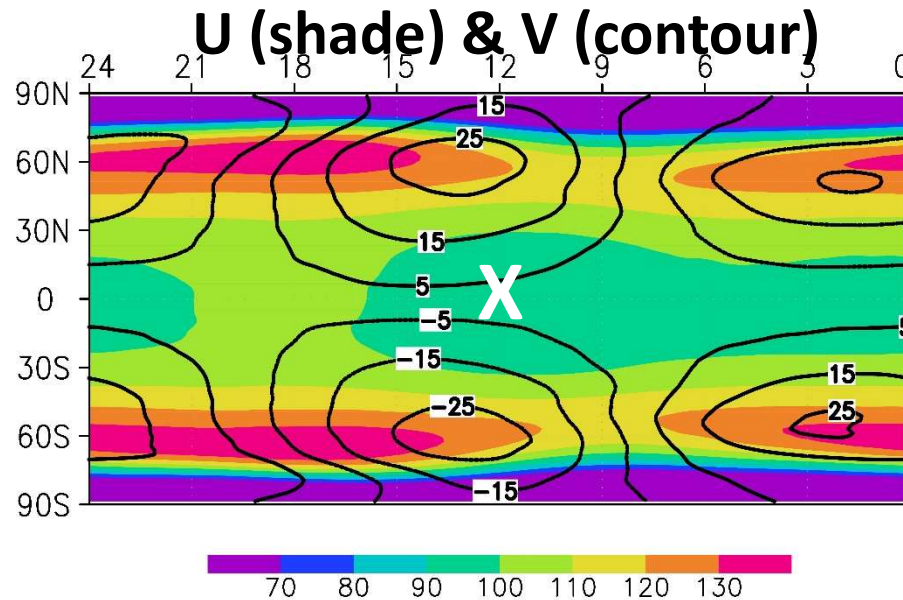
< Note >

These topographical effects are seen  
in the high-resolution (T63) model

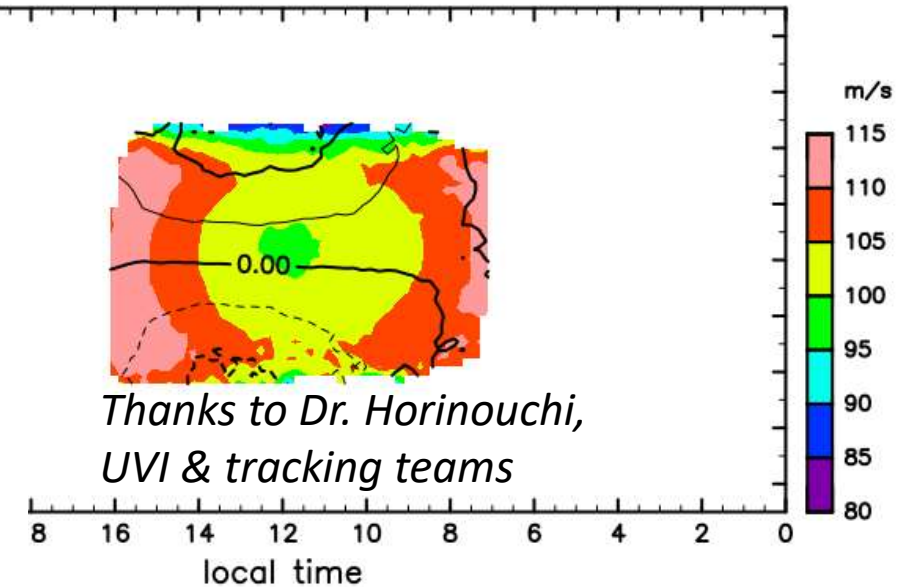


# Solar fixed circulation (solar day average)

around the cloud Top (~69.5 km)



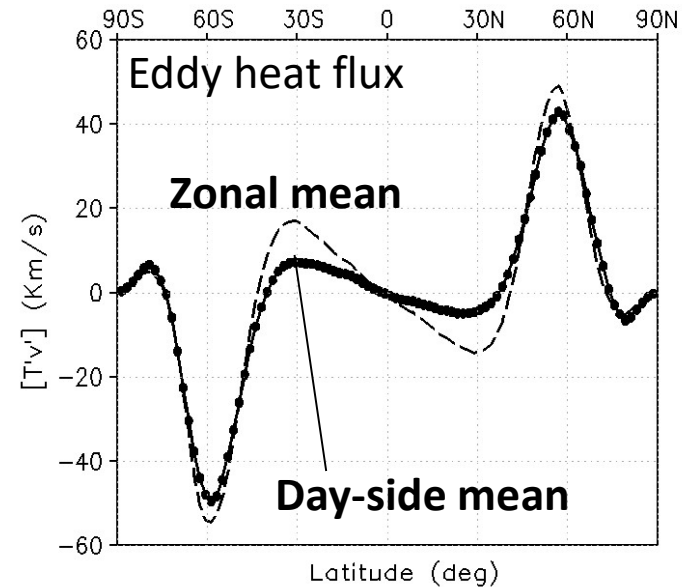
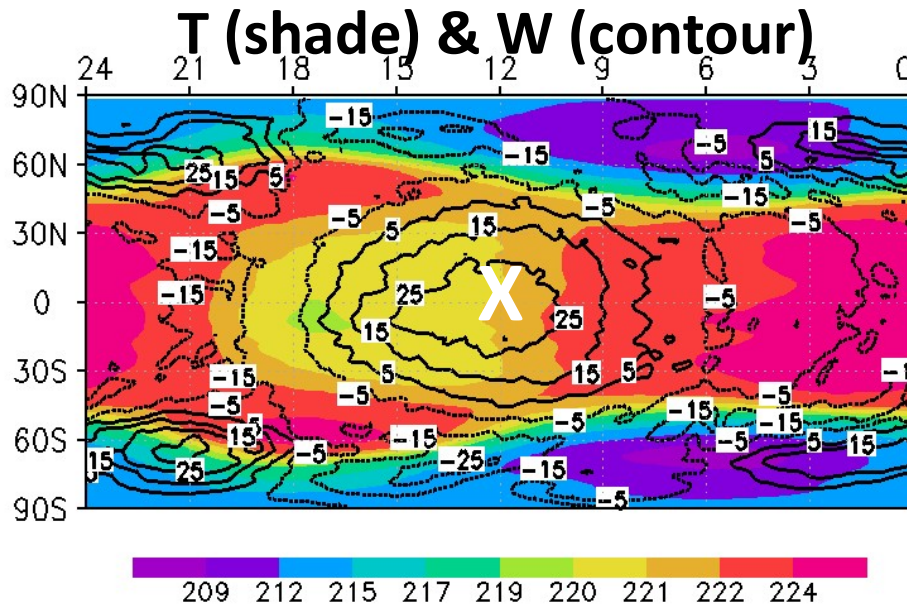
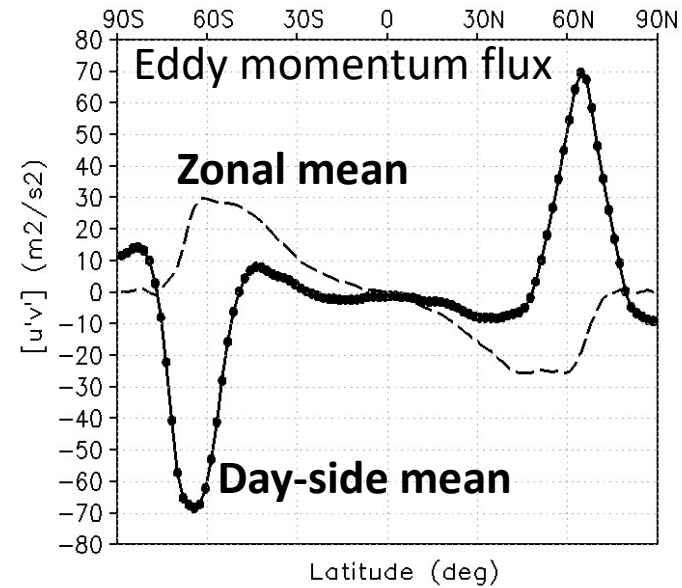
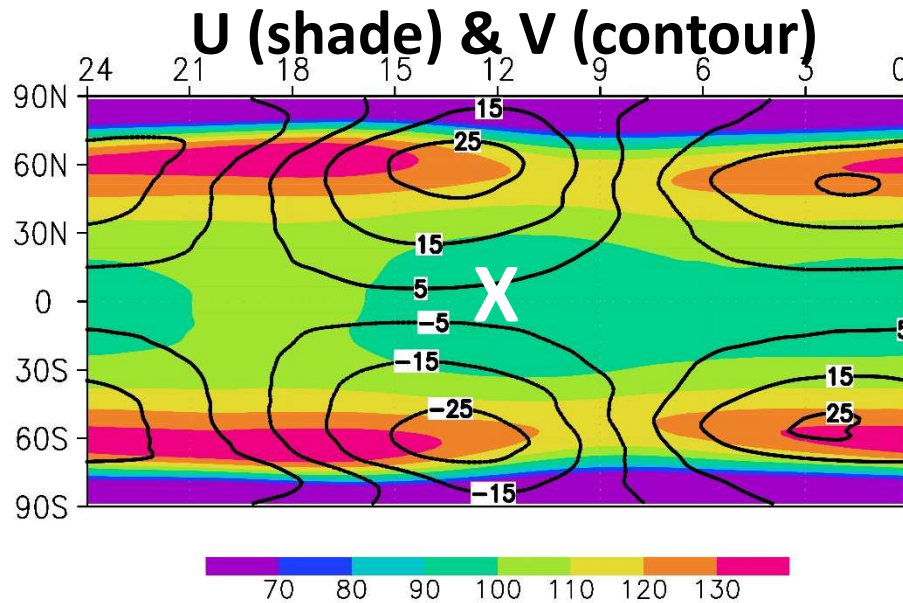
Akatsuki 201512–201703



Thanks to Dr. Horinouchi,  
UVI & tracking teams

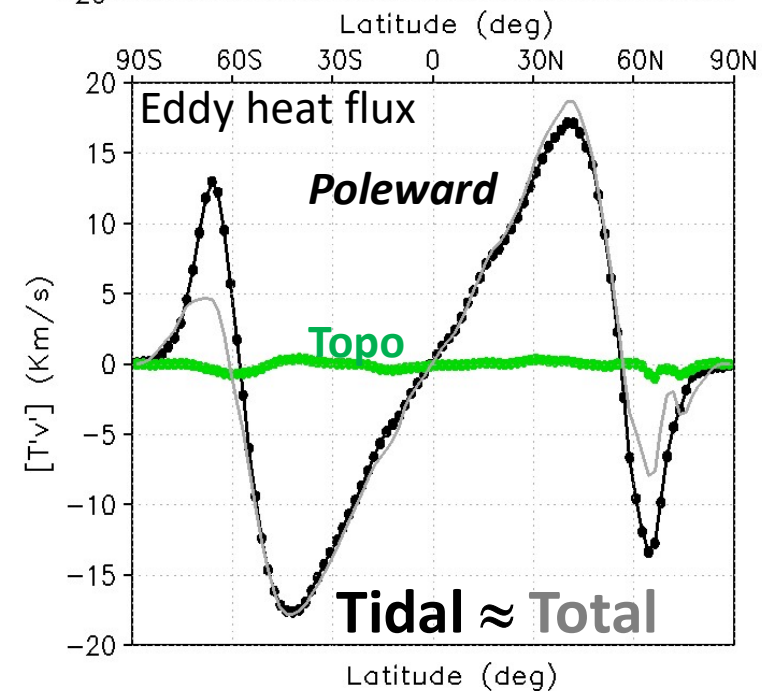
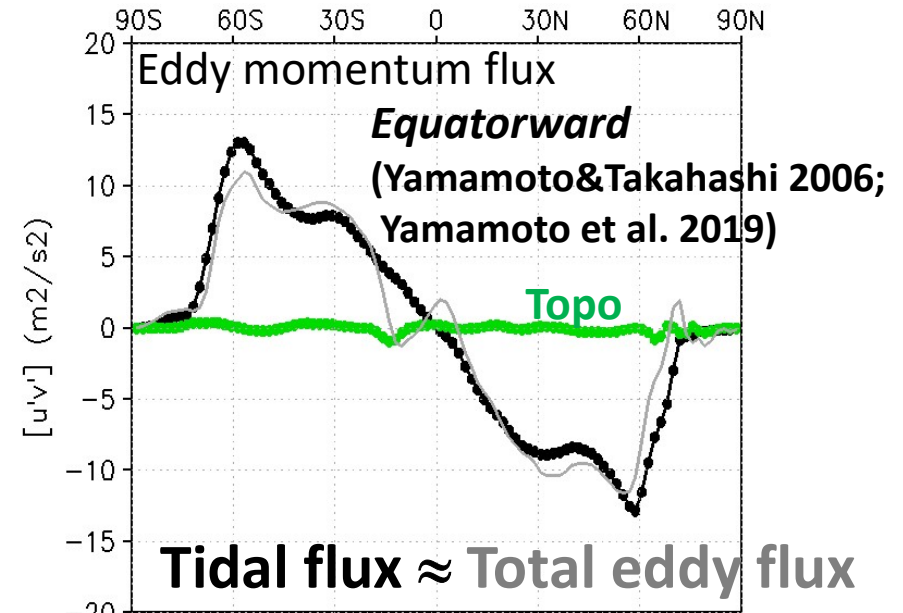
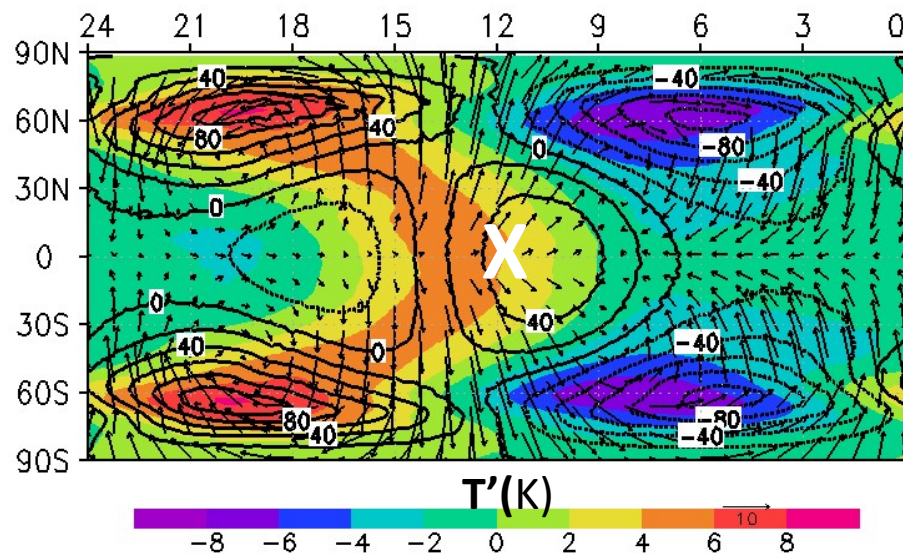
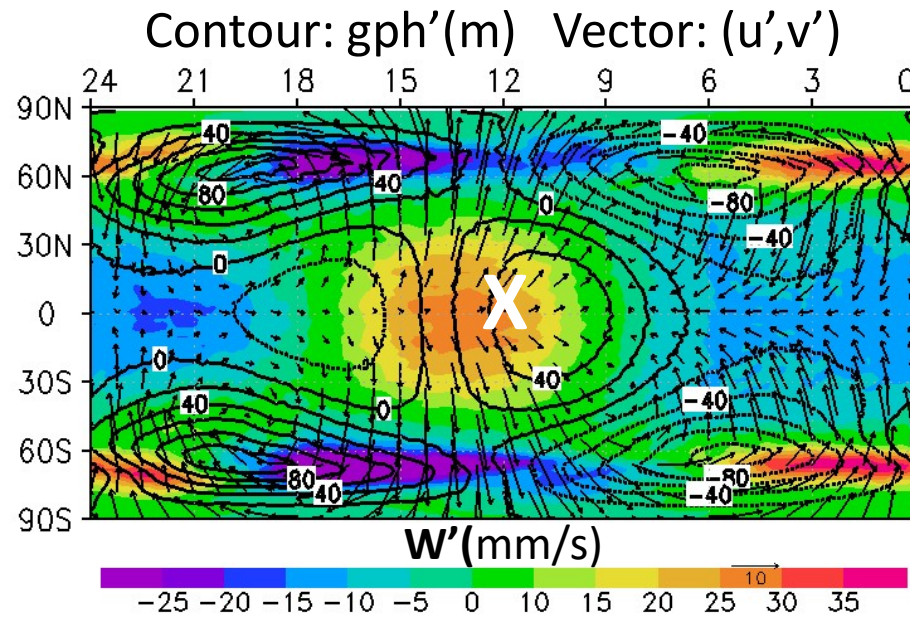
# Solar fixed circulation (solar day average)

around the cloud Top (~69.5 km)



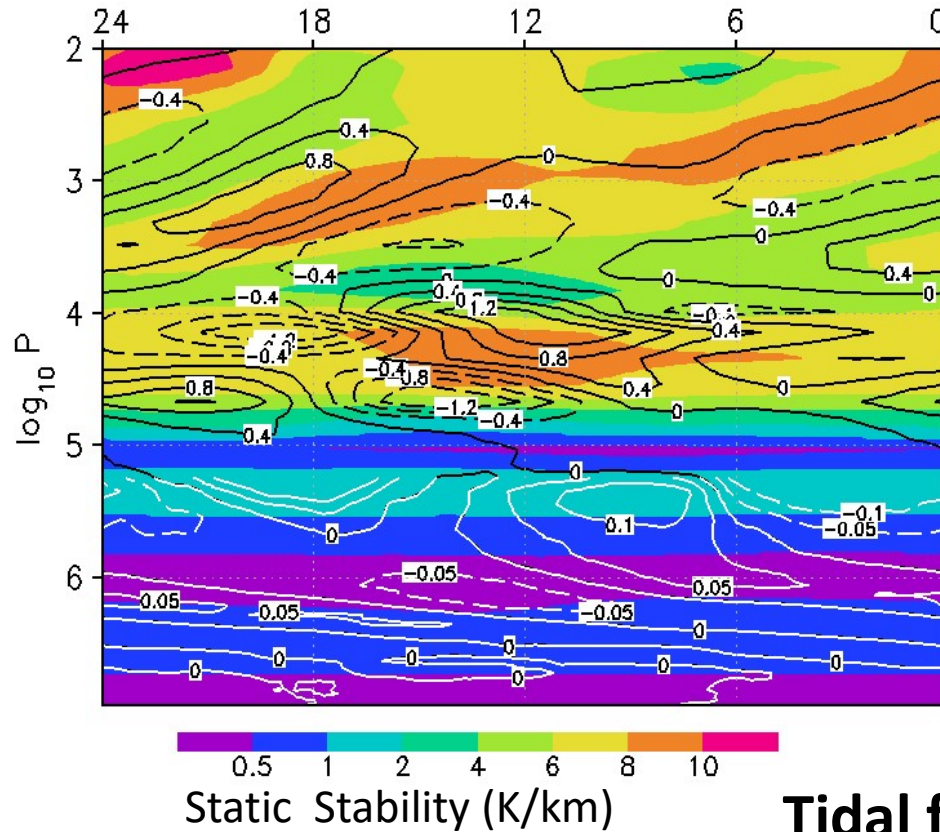


# Thermal tide at cloud heating max. (~65.5 km)

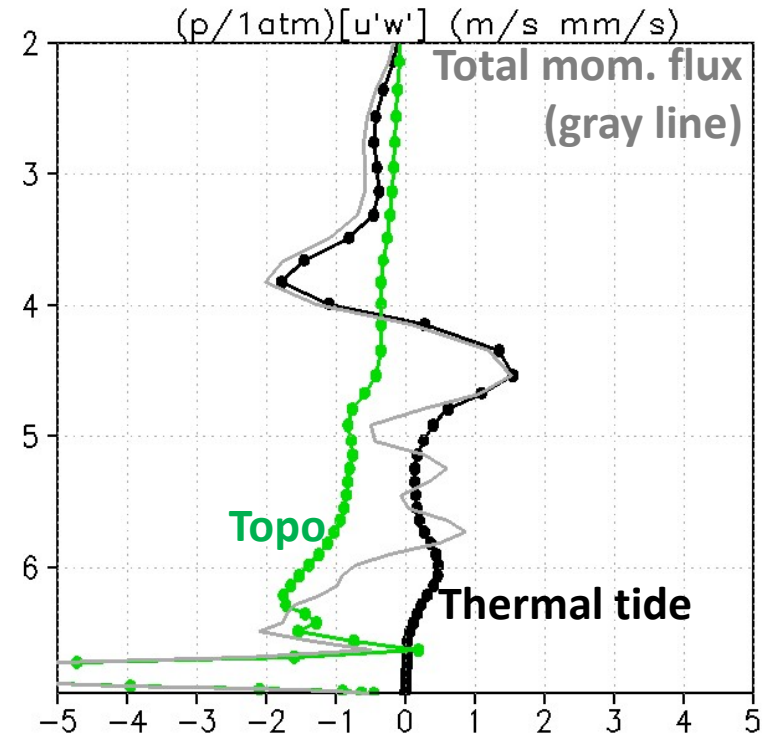


# Thermal tide near the equator (~0 deg lat.)

P-weighted  $T'$  (contour)



Eddy vertical mom. flux



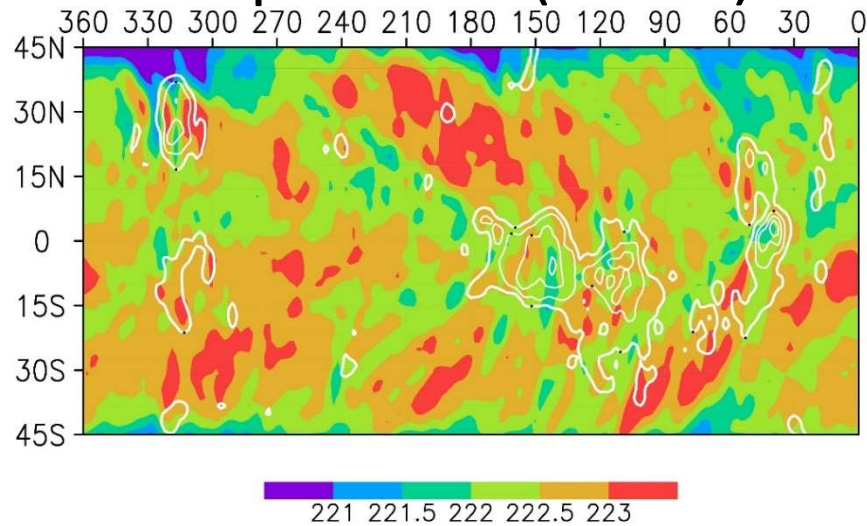
**Tidal flux  $\approx$  Total eddy flux**  
*in the middle atmosphere*

**Thermal Tide  $\ll$  Topograph. wave**  
*in the lower atmosphere*

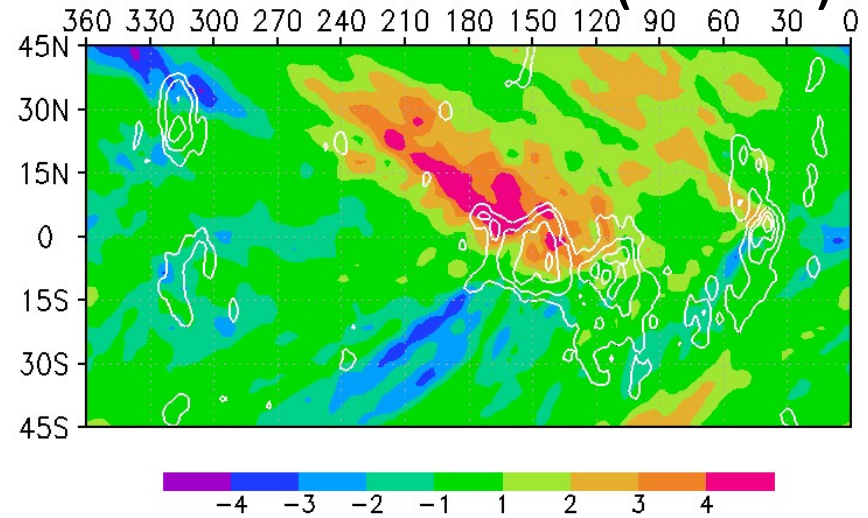


# Stationary structure (solar-day mean) around 69.5 km

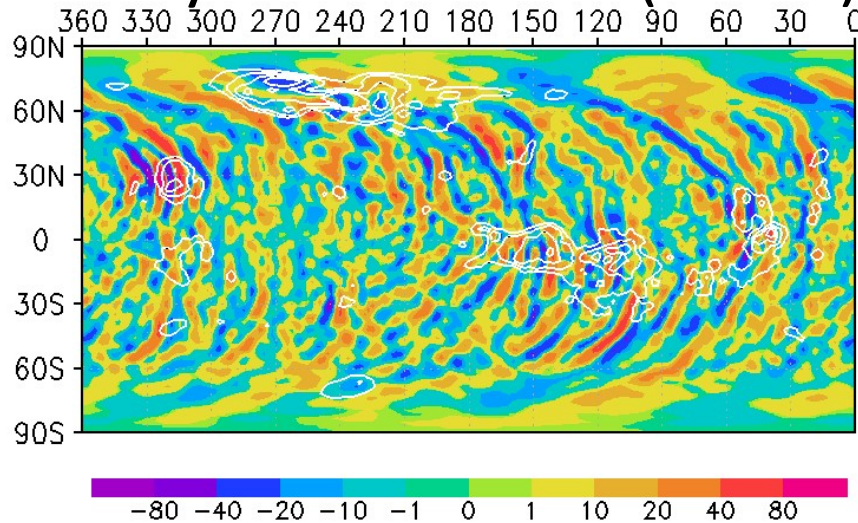
## Temperature (shade)



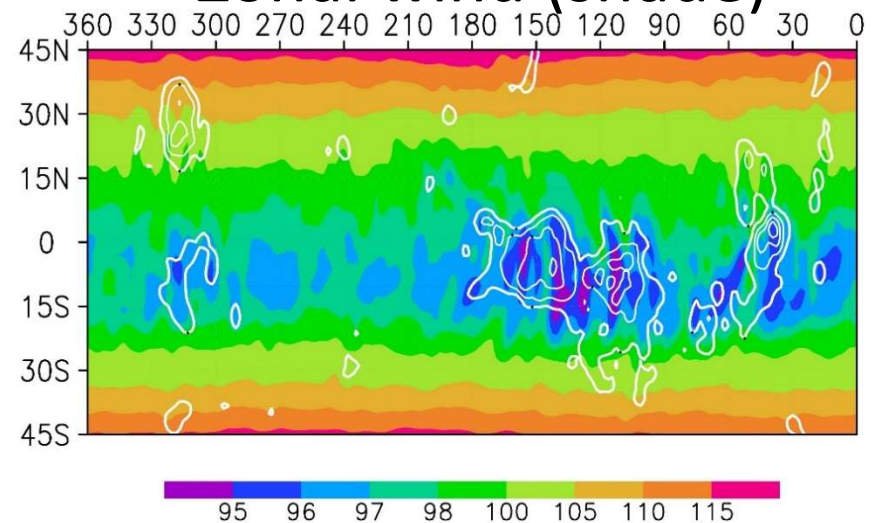
## Meridional wind (shade)



## Eddy vertical wind (shade)

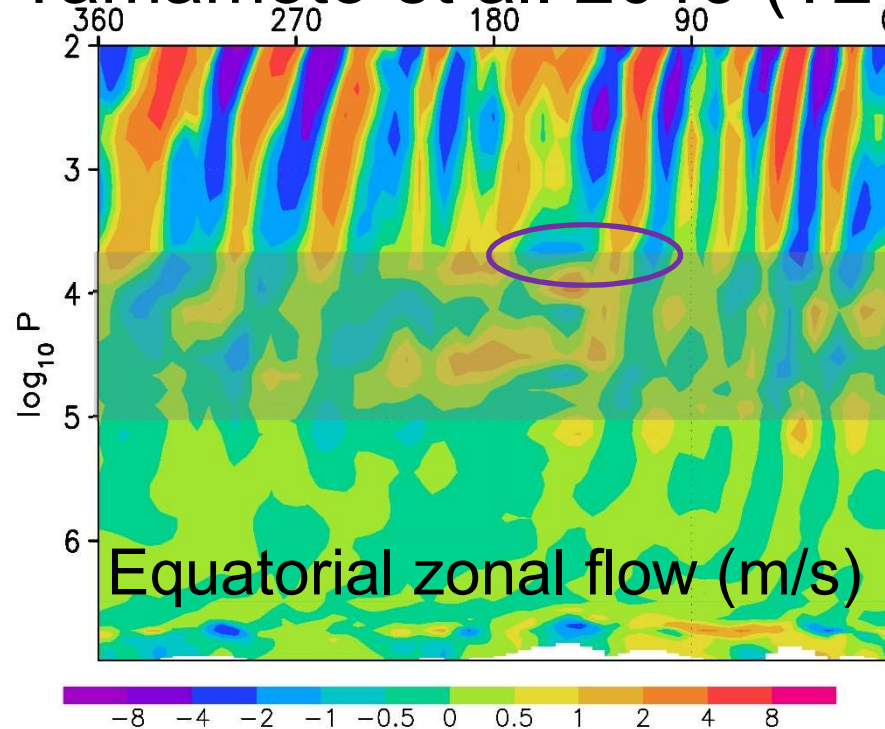


## Zonal wind (shade)

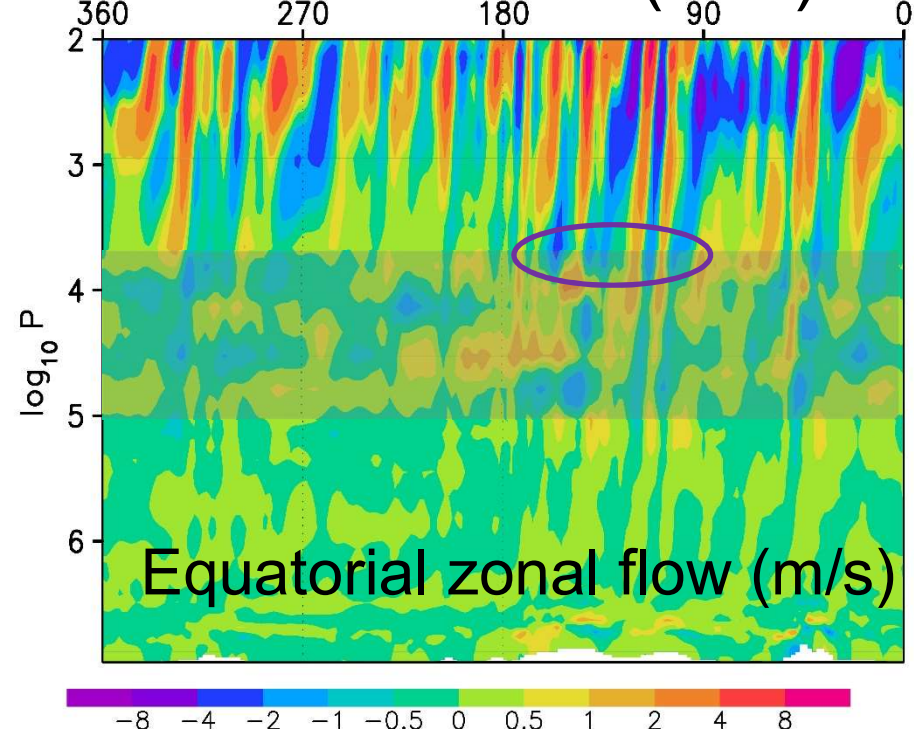


# Topographically stationary wave

Yamamoto et al. 2019 (T21)



Present work (T63)



The phase of stationary wave slightly tilts vertically.

The negative phase of the eddy zonal flow forms the local wind decrease at the cloud top over the Aphrodite terra similar to Bertaux et al. (2016).

# Summary(1)

- The UV tracked **horizontal flow around the subsolar point** is reproduced .
  - **Multi-layered and polar static stabilities** are reproduced.
  - **Thermal tides produce *equatorward* momentum flux** (Yamamoto & Takahashi 2006), along with the vertical flux.
  - The **zonal-wind decreases at 69 km over Aphrodite Terra.**
  - **Day-side mean eddy horizontal fluxes are quite different from the zonal means** owing to meridional flow of the tides.
- ⇒ These results are the same as the T21 simulation of Yamamoto et al. (2019 Icarus, 321, 232-250).



# Summary(2)

Further investigation shows the following results.

- (T63) **Topography weakens cloud-level SR** over the Aphrodite Terra and Maxwell Mt., while it **enhances upper-level poleward flow, zonal jet core and Ferrel circulation.**
- (T63) Topography induces **thin Hadley cell around the cloud base.**
- (T21,T63) **Thermal tides are dominant momentum transporters in the middle atmosphere, whereas topographical waves are dominant transporters in the lower atmosphere.**

## Future works

- Large- and small-scale cloud features
- Energy and angular momentum budgets