## 金星大気の観測 これまで (風速以外)

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Figure 4 Atmospheric temperature and pressure variations Venera 4–10 landers compared with the model distribution adopted in 1975 as the COSPAR Reference Model atmosphere. The model's surface position corresponds to planetary radius  $R_Q = 6050$  km.

Component	Relative abundance (by volume)	Technique	References
CO2	0.97±0.03	Gas-analyzer Venera 4, 5, and 6	Vinogradov et al. 1968, 1970a, b
N <sub>2</sub> (+ noble gases)	< 0.02	Venera 4, 5, and 6	Vinogradov et al. 1968, 1970a,b
	( 10 <sup>-3</sup> -10 <sup>-2</sup>	Venera 4, 5, and 6	Vinogradov et al. 1968, 1970a, b
	10-6-10-5	Ground-based spectroscopy	Schorn et al. 1969, Fink et al. 1972, Barker 1975
H <sub>2</sub> O	3 × 10 <sup>-4</sup>	Ratio of H <sub>2</sub> O and CO <sub>2</sub> bands intensities, Venera 9 and 10.	Moroz et al. 1976
	$< 2 \times 10^{-3}$	Radio astronomy	Kuzmin & Marov 1974, Smirnova & Kuzmin 1974
со	5 × 10 <sup>-5</sup>	Ground-based spectroscopy	Kuiper et al. 1967, Connes et al. 1968, Young, L. 1972
нсі	$4 \times 10^{-7}$	Ground-based spectroscopy	Connes et al. 1967, Young, L. 1972
HF	10-*	Ground-based spectroscopy	Connes et al. 1967, Young, L. 1972
	10-4-10-3	Gas-analyzer, Venera 8.	Surkov et al. 1973
NH3	<10-7	Ground-based spectroscopy	Kuiper et al. 1968-69
	<10-3	Radio astronomy	Kuzmin & Marov 1974, Smirnova & Kuzmin 1974

#### able 4 Chemical composition of the Venus atmosphere

Marov (1978)

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Horizontal Linear Polarization

# Polarization of sunlight reflected by Venus





Refractive index = 1.44  $\rightarrow$  consistent with H<sub>2</sub>SO<sub>4</sub>-H<sub>2</sub>O solution

Effective radius ~ 1  $\mu$ m

Hansen & Hovenier (1974)



FIG. 7. Observations and theoretical computations of the polarization of sunlight reflected by Venus at  $\lambda = 0.99 \ \mu$ m. The observations were made with an intermediate bandwidth filter, the X's being obtained by Coffeen and Gehrels (1969) in 1959–67 and by Coffeen (cf. Dollfus and Coffeen, 1970) from 1967 to March 1969, and the O's being obtained by Coffeen (cf. Dollfus and Coffeen, 1970) in May-July, 1969. The theoretical curves are for spherical particles having the size distribution (8) with b = 0.07. The different theoretical curves are for various refractive indices, the effective particle radius being selected in each case to yield closest agreement with the observations for all wavelengths.

## **Microphysical properties**

- $H_2SO_4-H_2O$  droplets with radii r < 5  $\mu$ m
- Smallest mode (including sub-cloud haze) might be condensation nuclei whose composition is unknown.
- Size distribution is variable.





Fig. 11. (LCPS). LCPS optical diagram.

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Knollenberg & Hunten (1980)





Gnedykh et al., 1987

## Three-layered structure – Different dynamical and chemical regimes at different altitudes?



#### Sulfur-rich atmosphere: origin of H<sub>2</sub>SO<sub>4</sub>?



Figure 24. The  $SO_2$  mixing ratio vertical profile retrieved for ISAV 2 (data points) is compared to that determined for ISAV 1. There is a large difference of structure above 40 km, while the profiles are nearly identical below 40 km. A peak of 210 ppm is observed at 43 km in the ISAV 2 data.

Bertaux et al. (1996)

## H<sub>2</sub>SO<sub>4</sub> vapor profiles



#### Traditional 1D model (Krasnopolsky & Pollack 1994; James et al. 1997)



### Ground-based observations of cloudrelated gaseous species



Pollack et al., Icarus 103, 1, 1993



Pollack et al., Icarus 103, 1, 1993



Fig. 9. Important chemical pathways for sulfur species.

Mills et al. (2007)

Zhang et al. (2012)

## SO, SO<sub>2</sub> profiles above cloud observed by Venus Express solar occultations (Belyaev et al. 2011)

![](_page_14_Figure_1.jpeg)

 Enhancement at high altitudes cannot be explained by traditional photochemical models.

#### (Zhang et al. 2012) Artificial $H_2SO_4$ source added above 90 km:

![](_page_15_Figure_1.jpeg)

Transport of cloud particles to the upper atmosphere by winds ? → Open question

**Fig. 8.** Same as Fig. 2, for the sulfur oxides. The  $SO_2$  and SO observations with errorbars are from the Belyaev et al. (2012). The temperature at 100 km is 165–170 K for the observations. The OCS measurement (0.3–9 ppb with the mean value of 3 ppb) is from Krasnopolsky (2010).

![](_page_16_Picture_0.jpeg)

Pioneer Venus 紫外画像(365 nm)

雲の模様の源

![](_page_17_Figure_1.jpeg)

![](_page_17_Picture_2.jpeg)

- 遠紫外(<320nm)の吸収はほぼSO<sub>2</sub>
- 近紫外(>320nm)の吸収物質の候補: S<sub>3</sub>, S<sub>4</sub>, S<sub>8</sub>, Cl<sub>2</sub>, FeCl<sub>3</sub>, SCl<sub>2</sub>, C<sub>5</sub>O<sub>5</sub>H<sub>2</sub>, (NH<sub>4</sub>)<sub>2</sub>S<sub>2</sub>O<sub>5</sub>, NOHSO<sub>4</sub>, S<sub>2</sub>O (Mills et al. 2007)

![](_page_18_Figure_0.jpeg)

Fig. 2 Observational range in volume absorption coefficient (shaded area) and volume extinction coefficient multiplied by (1 - g)(---) as a result of direct computation from the UV photometer data. Also shown are the extinction coefficients multiplied by (1-g) = 0.2 (after ref. 13) for mode-1 only (---) and total  $(\cdots)$ .

#### Ekonomov et al. (1984)

![](_page_18_Figure_3.jpeg)

FIG. 11. Global-mean solar heating rates for the nominal, "high-UV absorber" case (—) are compared to results obtained for the alternate "low-UV absorber" distribution (. . .) described in the text. The total amount of solar flux absorbed by these two distributions is identical. Other curves shown here are those obtained for the extreme cases where the entire planet is covered by a bright (——) or dark (---) UV feature.

Crisp (1986)

### Other vertical transport processes

![](_page_19_Picture_1.jpeg)

Y-feature - Kelvin wave?

![](_page_19_Picture_3.jpeg)

'Cells' near sub-solar region

## Convection cells near subsolar?

![](_page_20_Figure_1.jpeg)

### Radio wave generation by lightning discharges ?

![](_page_21_Figure_1.jpeg)

Figure 1. Waveform of 8-ms samples of the magnetic field at periapsis on 1 July 2006, bandpass filtered from 42 to 60 Hz and rotated into principal axis coordinates.

Russell et al. (2008)

![](_page_21_Figure_4.jpeg)

![](_page_22_Figure_0.jpeg)

![](_page_22_Picture_1.jpeg)

0.9

![](_page_22_Picture_2.jpeg)

![](_page_23_Figure_0.jpeg)

![](_page_24_Figure_0.jpeg)

Figure 11. A cartoon schematically illustrating the metallic frost model of Brackett et al. (1995). Volatile metal halides and chalcogenides (i.e., sulfides, selenides, and tellurides), are erupted into the atmosphere of Venus. Some vapor may immediately form aerosol hazes in the near-surface atmosphere of Venus, which may be the cause of the low-altitude hazes observed by two Pioneer Venus entry probes (Ragent and Blamont 1980). Some vapor may immediately snow out and condense out over the entire surface of the planet, and other more volatile species may remain in the near-surface atmosphere. See text for details.

## 大気-地殻化学相互作用

#### 炭酸塩による大気バッファ仮説 CaSiO<sub>3</sub> + CO<sub>2</sub> $\leftarrow \rightarrow$ CaCO<sub>3</sub> + SiO<sub>2</sub>

![](_page_25_Figure_2.jpeg)

## Variability of SO<sub>2</sub> above clouds

![](_page_26_Figure_1.jpeg)

![](_page_27_Picture_1.jpeg)

Long-term variations of the UV contrast on Venus observed by the Venus Monitoring Camera on board Venus Express

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![](_page_27_Figure_4.jpeg)