# A numerical study on convection of a condensing CO<sub>2</sub> atmosphere under an early Mars like condition

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### Warm climate and CO<sub>2</sub> ice cloud in Early Mars

- Major atmospheric component, CO<sub>2</sub>, condenses, and the ice clouds are widely distributed.
- In Early Martian climate study, <u>stratiform cloud</u> is assumed.
  The nature of convective CO<sub>2</sub> cloud is not studied well.



# Condensation convection of CO<sub>2</sub> and cloud type (Colaprete et al. 2003)

- Whether supersaturation is allowed or not is key point.
  - No: Stratiform cloud

Yes: Convective cloud



 However, circulation features and cloud distribution due to condesation convection of CO<sub>2</sub> are not examined.

- Results of Colaprete et al. (2003) is based on 1D cloud model.

# The purpose of this study

- We investigate circulation features and cloud distribution associated with condensation convection of CO<sub>2</sub> under Early Mars condition
  - Use 2D cloud resolving numerical model, implemented CO<sub>2</sub>
    ice cloud physics considering supersaturation
  - Perform long-term numerical simulations until statistical equilibrium states are reached.
- In this presentation...
  - We show dependencies of critical saturation ratio(Scr) on the circulation features and cloud distribution.
    - Scr is the saturation ratio for the onset of condensation.

# Outline of numerical model

- Dynamics: 2D Quasi-compressible fluid
  - Pure CO<sub>2</sub> atmosphere and its condensation is considered
- CO<sub>2</sub> ice cloud microphysics (Tobie et al. 2003)
  - Diffusional growth, gravitational settling of cloud particles
  - Number density of condensation nuclei is constant
- Radiation
  - Only IR is considered, given as a horizontal uniform cooling
- Surface flux (Louis, 1979)
  - Estimated from bulk formula with fixed surface temperature
- Subgrid scale turbulence (Klemp and Wilhelmson, 1978)
  - Turbulent flux is calculated by using urbulent kinetic energy.

# Governing equations

• Quasi-compressible equations(Klemp and Wilhelmson, 1978) and conservation equation for CO2 ice

Momentum equation:	$\frac{\partial \mathbf{u}}{\partial t} = -\mathbf{u} \bullet \nabla \mathbf{u} - C_p \overline{\theta} \nabla \Pi' + \mathbf{D}_{\mathbf{u}} + \frac{\theta'}{\overline{\theta}} \mathbf{g} - \frac{R}{p_0} \frac{\overline{\theta}}{\overline{\Pi}^{c_v/R}} \rho_s \mathbf{g}$							
Pressure equation:	$\frac{\partial \Pi'}{\partial t} = -\frac{\bar{c}_s^2}{C_p \bar{\rho} \bar{\theta}^2} \nabla \bullet \left(\bar{\rho} \bar{\theta} \mathbf{u}\right) + \frac{\bar{c}_s^2 L}{C_p^2 \bar{\rho} \bar{\theta}^2 \overline{\Pi}} M_{cond} - \frac{\bar{c}_s^2}{C_p \bar{\rho} \bar{\theta}} M_{cond}$							
Thermodynamic equation:	$\frac{\partial \theta'}{\partial t} = -\mathbf{u} \bullet \nabla \theta' - w \frac{\partial \overline{\theta}}{\partial z}$	$+\frac{1}{\overline{\Pi}}\left(\frac{LM_{cond}}{C_p\overline{\rho}}+Q_{dis}-\right)$	+Q	$D_{rad} + D_{\theta}$				
Conservation equation for CO2	$\frac{\partial \rho_s}{\partial t} = -\nabla \bullet \left( \rho_s \mathbf{u} \right) + M_{fall} + M_{cond} + D_{\rho_s}$			$\mathbf{u} = (u, w)$ : Velocity, $\theta$ : Potential temperature, $\Pi$ : The Exner function, $\rho$ : Density of vapor, $\alpha$ : Density of cloud $T$ : Temperature				
ice:		Condensation		$c_s$ : Sound speed,				
		/evaporation		$C_p$ : Specific heat at constant pressure,				
		term		L: Latent heat, $M_{cond}$ : Condensation rate,				
			-	$Q_{dis}$ : Dissipative heating rate,				
				$Q_{rad}$ : Radiative heating rate,				
				$\mathbf{D}_{\mathbf{u}} = (D_u, D_w), D_{\theta}, D_{\rho_s}$ : Turbulent diffusion term				
				$\mathbf{g} = (0, g)$ : Gravitational accerelation				

# CO<sub>2</sub> ice cloud microphysics

• Condensation/Evaporation rate (Tobie et al., 2003)

$$M_{cond} = \frac{4\pi r \rho N_* k R \theta^2 \Pi^2}{L^2} (S-1)$$

- For 1<S<Scr, condensation does not occur if cloud density is less than a threshold value
- We assume the value of threshold as 1.0<sup>-6</sup> kg/m<sup>3</sup>
- Gravitational settling rate

$$M_{fall} = \frac{\partial}{\partial z} (\rho_s V_{term})$$
$$V_{term} = \left(1 + \frac{4}{3}K_n\right) \frac{2r^2 g\rho_I}{9\eta}$$

 $S = \frac{p}{p_*}: \text{Saturation ratio}$   $p: \text{Pressure} \quad p_*: \text{Saturation vapor pressure}$   $\rho: \text{Gas density} \quad \eta: \text{viscosity coefficien t}$   $N^*: \text{Number density of condensation nuclei}$   $r: \text{Radius of cloud particle} \quad \rho_I: \text{CO2 ice density}$  Kn: Knudsen number for cloud particle  $k: \text{Thermal diffusion coefficien t} \quad R: \text{Gas constant}$ 

# Setup of experiments

- Domain size:
  - 100km in horizontal direction (grid spacing: 500m)
  - 80km in vertical direction (grid spacing: 400m)
- Initial temperature and cooling profile
  - Based on Kasting (1991)

- Critical saturation ratio (Scr): 1.0, 1.35 (Glandorf et al., 2002)
- Number density of condensation nuclei (N\*): 5.0x10<sup>4</sup>, 5.0x10<sup>6</sup>, 5.0x10<sup>8</sup> /kg
  (Tobie et al., 2003: Forget et al., 2013)



# Results

### Parameter dependence: time evolution of total cloud mass



# Parameter dependence: time evolotion of total kinetic energy

Scr = 1.0



#### Circulation features and cloud distribution: <u>Scr=1.0</u> (100 days)



#### Circulation features and cloud distribution: <u>Scr=1.35</u> (143 days)



#### Dependency of number density of condensation nuclei: Scr=1.35



### Estimation of radius of cloud particles and cloud optical depth

$S_{cr}$	$N_{*} (kg^{-1})$	radius of cloud particles( $\mu$ m)			optical depth			
		all	non cond.	cond.	$\operatorname{all}$	non cond.	cond.	
1.0	$5.0 \times 10^4$	35			5			
1.0	$5.0 \times 10^6$	15	—	—	70	—	—	
1.0	$5.0 \times 10^8$	5	_	—	1500			
1.35	$5.0 \times 10^4$	5	5	30	1	0.5	10	
1.35	$5.0 \times 10^6$	2	0.1	10	15	0.01	100	
1.35	$5.0 \times 10^8$	5	—	—	1500	—		

- Suitable combination of cloud particle radius and optical depth for "scattering green house effect" does not occur.
  - The estimation of optical depth is following to Petty (2006)
  - The scattering green house effect is effective when cloud particle radius is 10 μm and its optical depth is about 10.

# **Concluding remarks**

- Circulation features and cloud distributions associated with condensation convection of CO<sub>2</sub> vary greatly with the values of Scr.
- Scr=1.0: No supersaturation case
  - Condensation occurs continuously
  - <u>Stratiform clouds</u> is generated by gravity wave
- Scr=1.35: Supersaturation case
  - Non-condensation and condensation periods occur alternately
  - In condensation periods, <u>convective clouds</u> forms
- Note: In any cases, suitable combination of cloud particle radius and optical depth for scattering green house effect does not occur.

This study is submitted to Journal of the Atmospheric Sciences.

# Appendices

#### Horizontal mean cloud distribution



#### Dependency of number density of condensation nuclei:



### **Total cloud mass**



### Total kinetic energy



### Total kinetic enregy and cloud mass



### Vertical profiles of cloud mass density in no supersaturation case



- Vertical profiles of cloud mass density is determined by
  - In the Uppper part: balances between radiative cooling and condensation heating, and condensation rate and gravitational setteling (red dot line)
  - In the lower part: conservation of liquid water static energy (blue dush line)