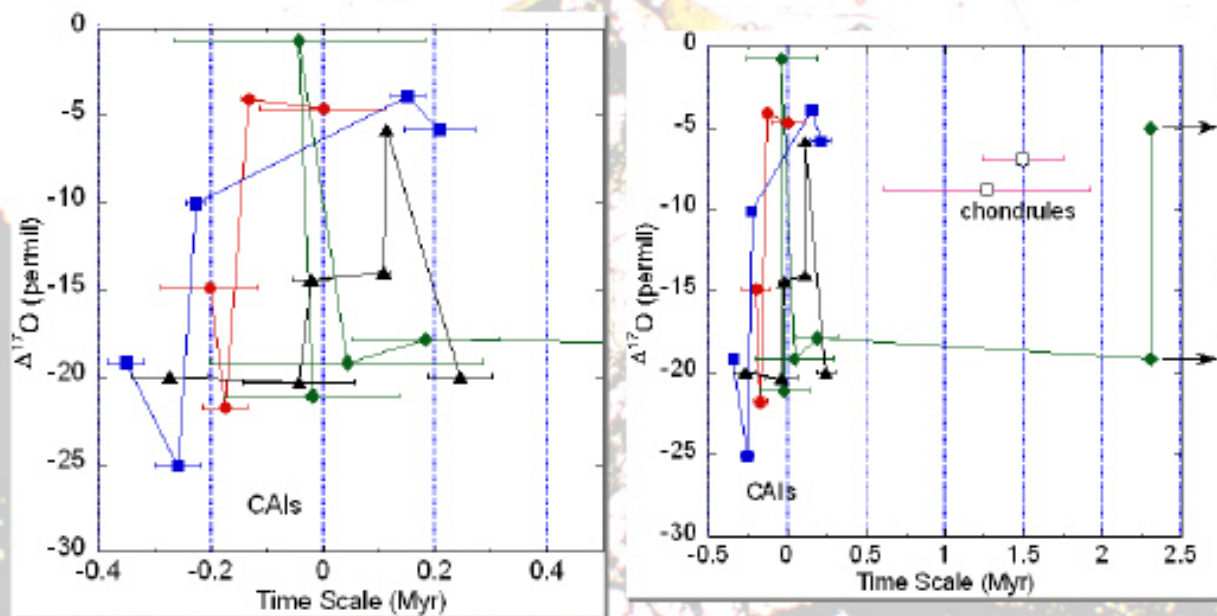


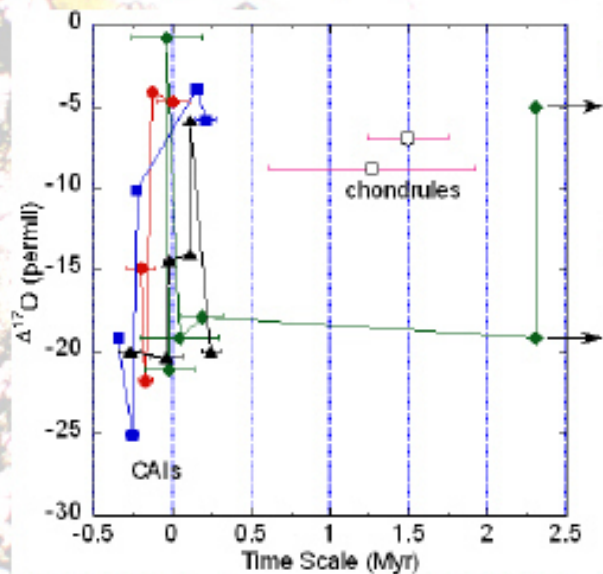
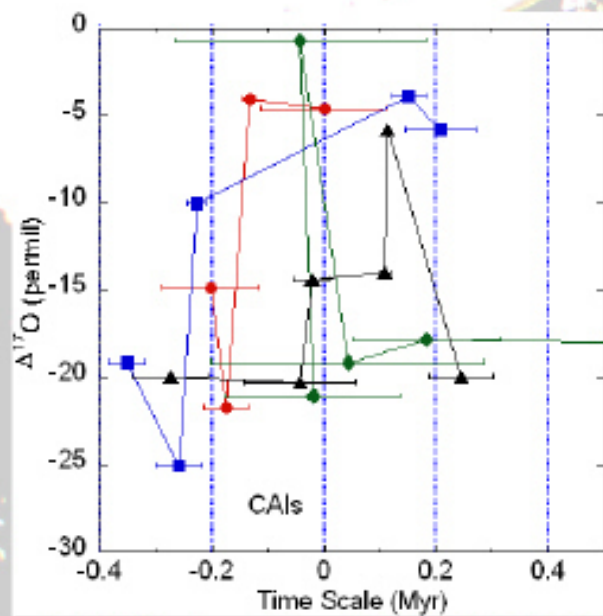
# 原始太陽系円盤ガスの酸素同位体変動



- 1個のCAIの形成期間（一瞬～数百万年）
- 1個のCAI周辺の円盤ガスは10万年より短い間に異なる酸素同位体比をもつものに切り替わる。
- 酸素同位体比の切り替わり現象は数百万年間起きている。

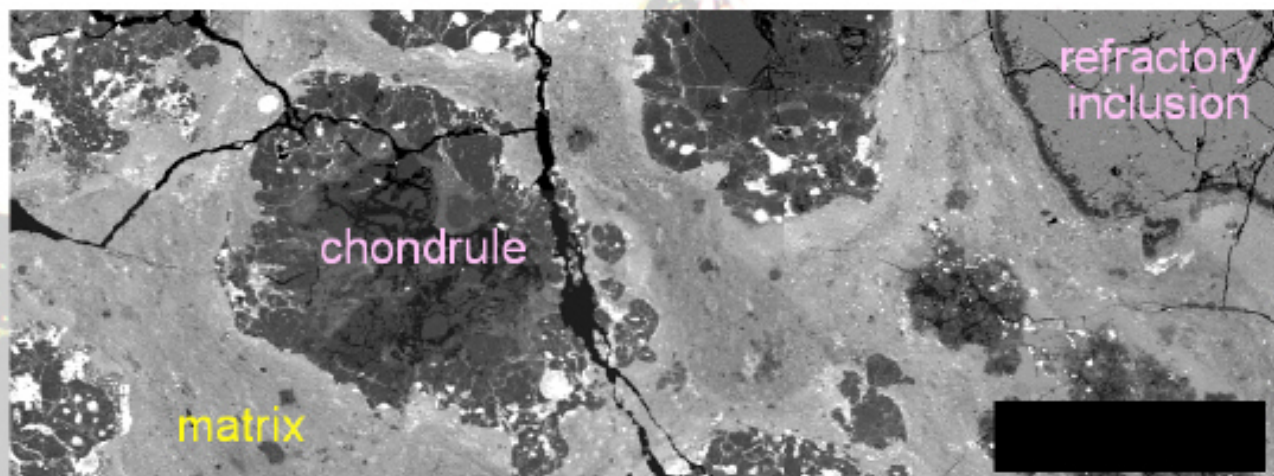
# 原始太陽系円盤ガスの酸素同位体変動

巻出健太郎, 吉武美和, 伊藤正一



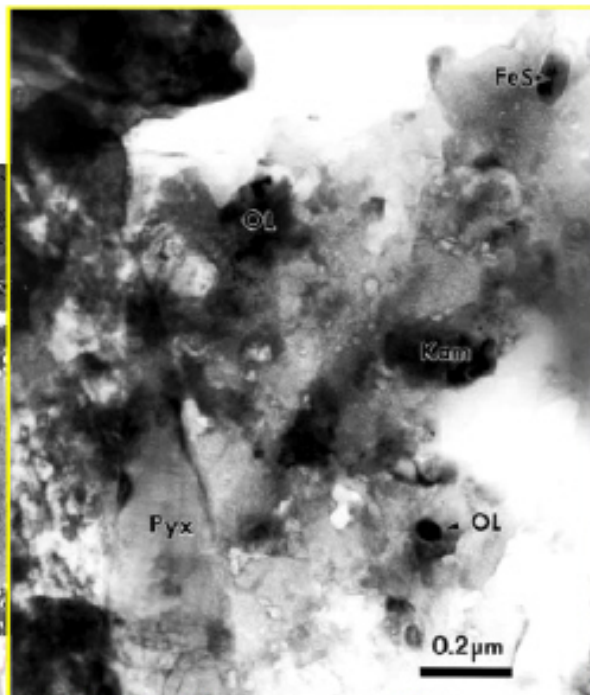
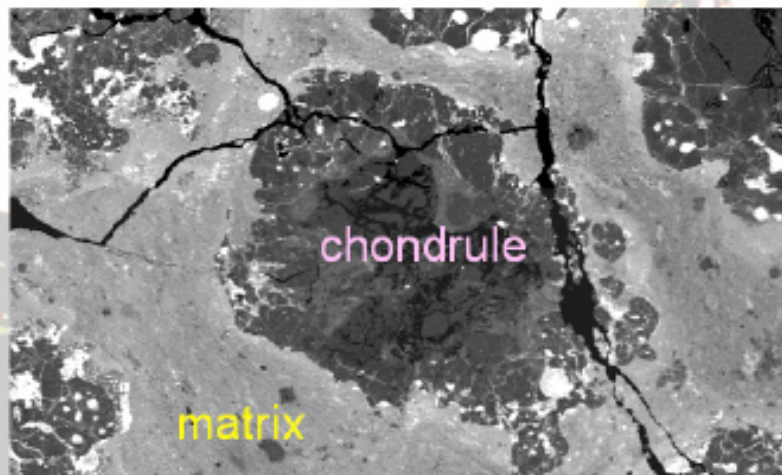
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## Mineralogy of matrix



- Disequilibrium mixtures
  - silicates, oxides, metals, sulfides, carbonates, phyllosilicates, hydrous silicates, amorphous materials
- Grain size: 10 nm–10  $\mu\text{m}$

# Mineralogy of matrix



ALHA 77307 (Brearley 1993)

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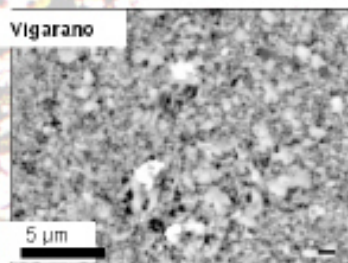
# Controversial origin of matrix material



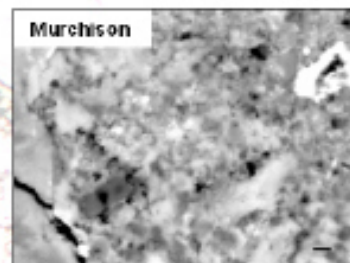
- Prior to the solar system
  - presolar materials (eg. Zinner 2003)
- In the solar system
  - nebular origin
    - condensates
    - chondrule/refractory inclusion related dusts
  - chondrite parent-body origin
    - accreted material is aqueously altered (eg. Brearley 2003)
- O-isotopes are a useful parameter for understanding genesis of matrix materials

# Alteration on chondrite parent-body

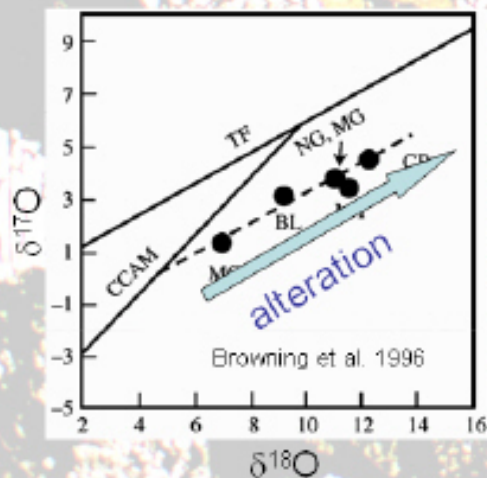
- Presence of phyllosilicates indicates aqueous alteration (Brearley 2003)



- slightly altered
- minor phyllosilicates

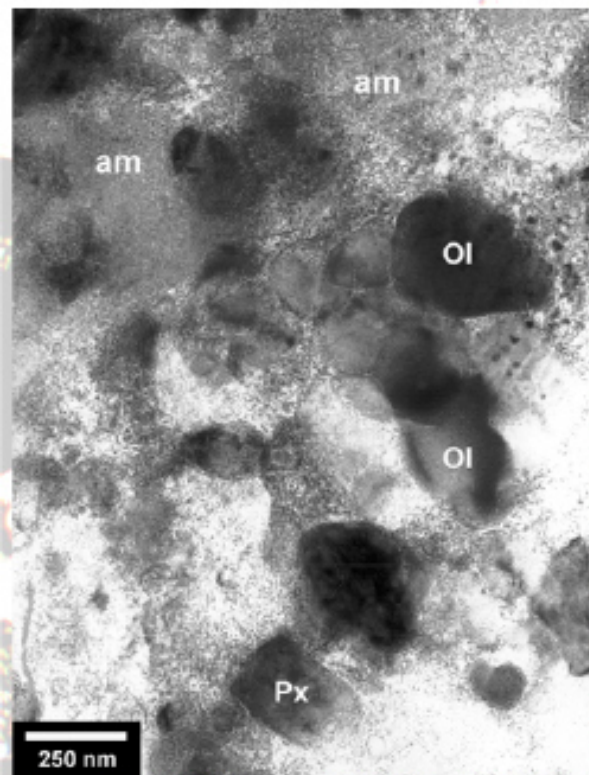


- pervasive alteration
- abundant phyllosilicates



- Correlation between O-isotopes and degree of alteration (Browning et al. 1996)
- Unaltered matrix should be examined to see O-isotope of nebular features

# Matrix of primitive meteorites



Acfer 094 (Greshake et al. 2004)

component	abundance in matrix
amorphous silicates	~40 vol%
crystalline silicates	~50 vol%
sulfides	~ 5 vol%

Greshake et al. 1997

- lack of phyllosilicates
- origin of amorphous silicates
  - condensation / irradiation
  - unlikely to be produced or reprocessed on parent body
- minimal alteration and metamorphism on parent body

# Phases in the matrix

chondrite	matrix phases
CI	<u>serpentine</u> , <u>saponite</u> , <u>ferrihydrite</u> , magnetite, Ca–Mg carbonate, pyrrhotite
CM	<u>serpentine</u> , <u>tochilinite</u> , pyrrhotite, amorphous phase, calcite
CR2	ol, <u>serpentine</u> , <u>saponite</u> , FeS, pentlandite, pyrrhotite, calcite
CV3 oxidized	fayalitic ol, Ca–Fe px, <u>nepheline</u> , pentlandite
CO >3.1	fayalitic ol, <u>phyllosilicates</u> , ferric oxide
<b>CV3 reduced</b>	fayalitic ol, low-Ca px, low-Ni metal, FeS
<b>CO 3.0</b>	amorphous silicate, Fo <sub>30-98</sub> , low-Ca px, Fe, Ni metal, magnetite, sulfides
<b>Acfcr 094</b>	amorphous silicate, forsterite, en, pyrrhotite, ferrihydrite
<b>Adelaide</b>	amorphous silicate, fayalitic ol, en, pentlandite, magnetite



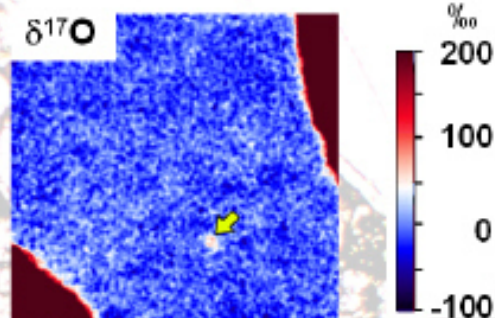
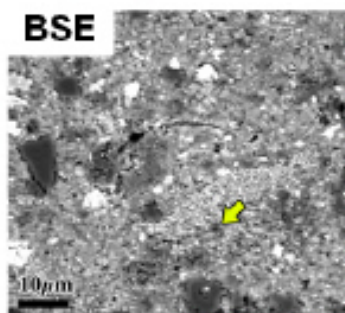
# Recent studies of matrix



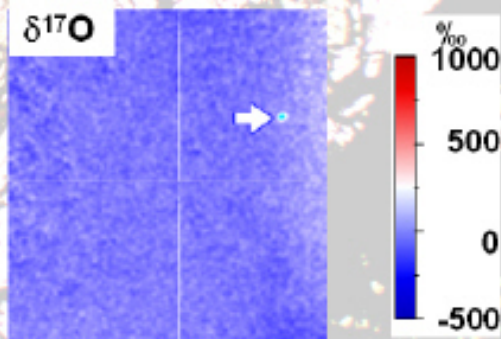
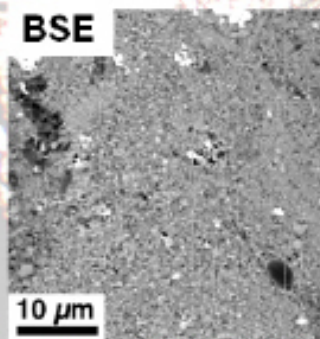
- New techniques allow isotope mapping
- Analyses of primitive matrix
  - **Acfer 094** (Nagashima et al., 2004; Mostefaoui and Hoppe, 2004; Stadermann et al., 2005; Bland et al., 2005)
  - **ALHA77307** (Kobayashi et al., 2005; Nguyen et al., 2005)
  - **Vigarano** (Kunihiro et al., 2005; Nagashima et al., in prep.)
  - **Semarkona, Bishunpur** (Mostefaoui et al., 2004)

# Presolar silicates in matrix

Acfer 094  
(Nagashima et al. 2004)



Vigarano  
(Nagashima et al. unpub.)



- Presolar silicates found in matrix (10-100 ppm)
- Grain size: < 1 μm

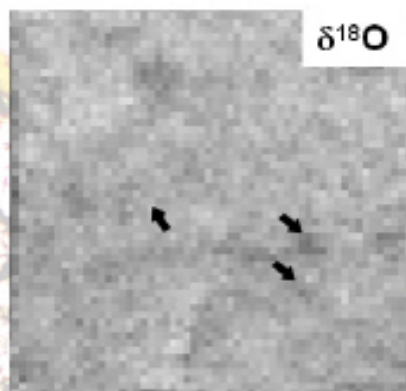
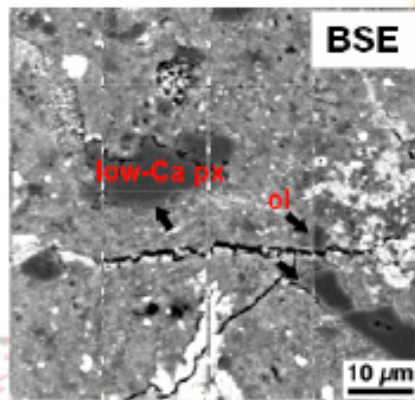
## Preservation of sub-micron presolar silicates in matrices

- Aqueous alteration on these meteorites did not equilibrate O-isotopes of sub-micron grains
- Matrices of primitive meteorites retain O-isotope distribution of the nebular dust

# Homogeneity in matrix

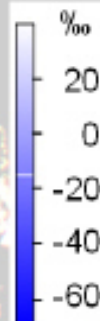
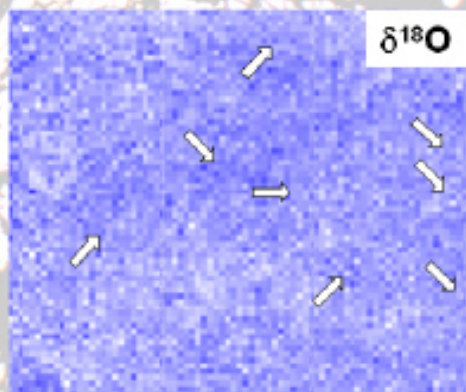
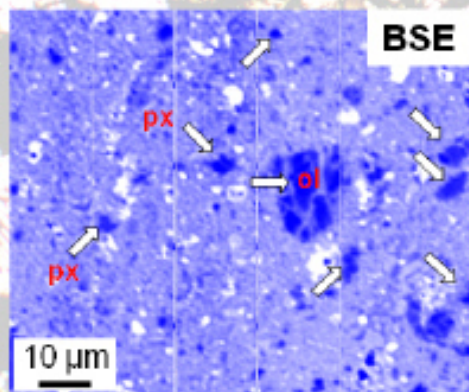
Acfer 094

(Nagashima et al. unpub.)



Vigarano

(Kunihiro et al. 2005)



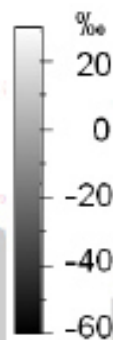
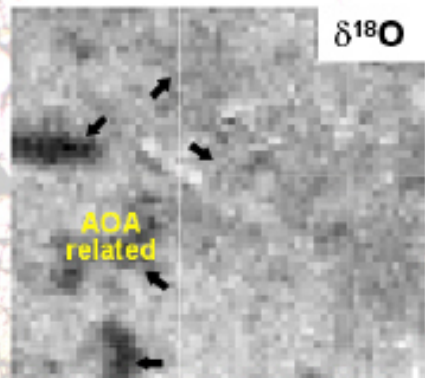
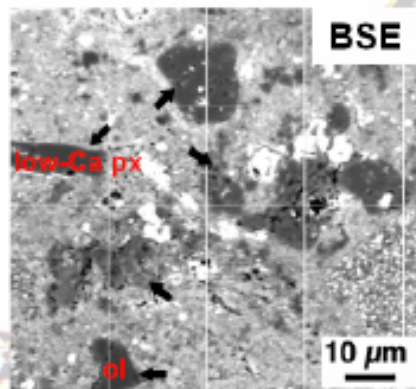
## Origin of the homogeneity

- Vast majority of matrix is:
  - $^{16}\text{O}$ -poor
  - homogeneous ( $\pm 5\text{‰}$ )
- Most of the matrix is not of presolar origin
- Matrix inherited its O-isotope homogeneity from the solar nebula

# Matrix with refractory components

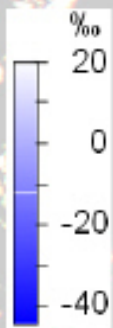
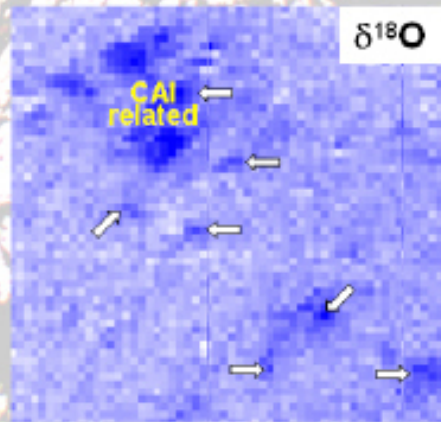
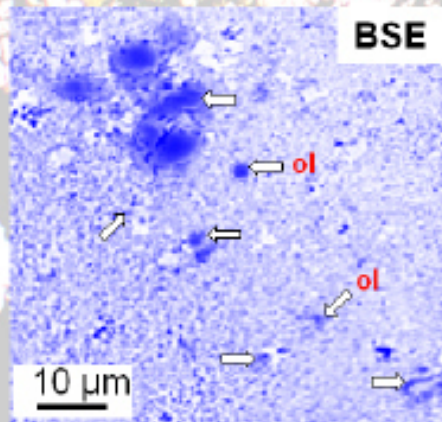
Acfer 094

(Nagashima et al. unpub.)



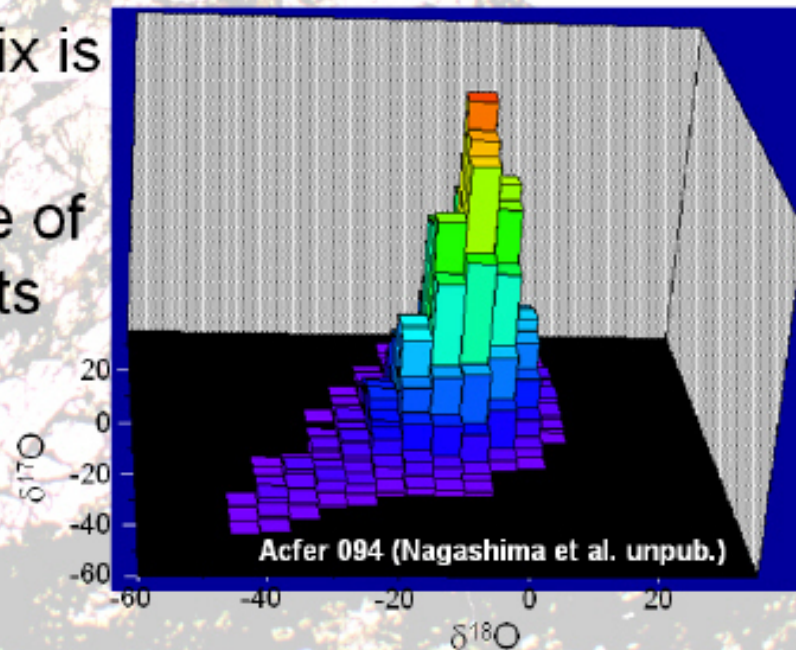
Vigarano

(Kunihiro et al. 2005)



# Quantification of matrix components

- Distribution is consistent with other CC components
- Only 1–5% of matrix is enriched in  $^{16}\text{O}$
- Mechanical mixture of nebular components



# Nebular dust

- Constituents of nebular dust

<b>presolar dust</b>	<b>&gt; 100 ppm</b>
<b><math>^{16}\text{O}</math>-rich solar dust</b>	<b>1-5 vol. %</b>
<b><math>^{16}\text{O}</math>-poor solar dust</b>	<b>&gt; 90 vol. %</b>

- Nebula mechanisms reprocessed most presolar dust generating solar O-isotope characteristics
- Most solar dust is  $^{16}\text{O}$ -poor



## Origin of $^{16}\text{O}$ -poor dust

- O-isotopic equilibration with  $^{16}\text{O}$ -poor gas
- The equilibration was completed before parent-body accretion
  - gas-solid equilibration in the nebula
  - condensation from  $^{16}\text{O}$ -poor gas
- Chondrule-forming region is a possible place to provide both heat source and  $^{16}\text{O}$ -poor reservoir

# Summary

- O-isotopes of primitive matrix
  - heterogeneity
    - presolar silicates
    - refractory-inclusion related
  - most of matrix is  $^{16}\text{O}$ -poor and homogeneous ( $\pm 5\%$ )
- Most of nebular dusts seem to be equilibrated with  $^{16}\text{O}$ -poor gas
- Nebular dust is a mixture of grains from:
  - presolar origin
  - refractory-inclusion-forming region
  - chondrule-forming region